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Temporal variations of PM concentrations, and its association with AOD and meteorology observed in Nanjing during the autumn and winter seasons of 2014–2017

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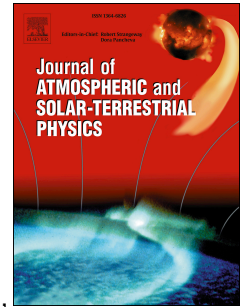
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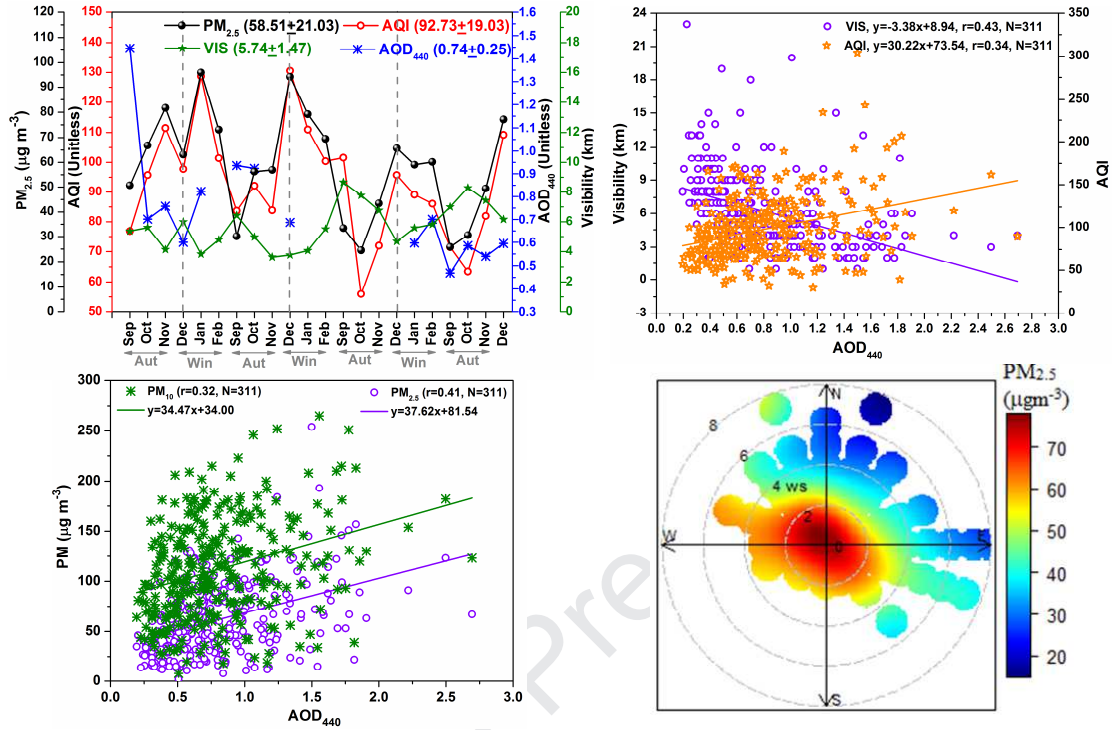
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Graphical Abstract



1 **Temporal variations of PM concentrations, and its association with**
2 **AOD and meteorology observed in Nanjing during the autumn and**
3 **winter seasons of 2014-2017**
4

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34 **ABSTRACT**

35 The present study aims to investigate temporal evolutions of particulate matter (PM)
36 concentrations and its association with the meteorology and aerosol optical depth (AOD) during
37 autumn and winter of 2014-2017 at an urban city, Nanjing in the Yangtze River Delta, East
38 China. The seasonal mean $PM_{2.5}$ (PM_{10}) was found maximum and minimum with $81.2 \pm 41.5 \mu g$
39 m^{-3} ($135.6 \pm 57.1 \mu g m^{-3}$) and $33.7 \pm 19.1 \mu g m^{-3}$ ($65.8 \pm 34.5 \mu g m^{-3}$) during winter and autumn
40 seasons, respectively. Furthermore, the mean ratio of $PM_{2.5}/PM_{10}$ was around ~ 0.57 for the entire
41 study period, with a lower contribution (0.53) in autumn and higher (0.60) in winter. However,
42 the seasonal mean AOD_{440} , precipitable water vapor content, Ångström exponent ($AE_{440-870}$) was
43 found maximum with 0.97 ± 0.31 , 1.58 ± 0.80 cm, and 1.14 ± 0.23 during autumn, and a
44 minimum of 0.62 ± 0.34 , 0.60 ± 0.27 cm, and 1.29 ± 0.19 in winter, respectively. The potential
45 source contribution function (PSCF) and concentration weighted trajectory (CWT) models
46 revealed considerable long-distance transport of $PM_{2.5}$ from north and northwest China. Besides,
47 the concentration bivariate probability function (CBPF) revealed a significant contribution of
48 $PM_{2.5}$ occurred when the winds blown from southerly and northwesterly directions. The
49 relationship between $PM_{2.5}$ and meteorology found that $PM_{2.5}$ concentration had a positive
50 relationship with AQI, while negative correlations with the major meteorological parameters. A
51 notable spatial heterogeneities and trends were observed in AOD and AE, with negative
52 correlations (-0.5 to 0) in winter over East China.

53

54 **Keywords:** Particulate matter; Meteorology; Aerosol optical depth; PSCF and CWT models.

55

56

57 1. Introduction

58 The importance of atmospheric aerosols such as particulate matter (PM) and cloud
59 condensation nuclei (CCN) for the climate is well known and is an inevitable topic in
60 Atmospheric Physics and Chemistry (Kalluri et al., 2016). They consist of solid or liquid
61 particles suspended in the atmosphere with various sizes (1 nm-1 μm), and play a crucial role in
62 weather and climate change. Aerosols (natural and anthropogenic) are often mixed, and can also
63 be described by their size, e.g., PM_{10} are particles with aerodynamic diameters less than 10 μm ,
64 while $\text{PM}_{2.5}$ particles with aerodynamic diameters less than 2.5 μm . They affect the climate
65 directly by scattering and absorbing solar radiation; and thereby, altering radiation budget at the
66 top, surface and within the atmosphere, which in turn influences the atmospheric heating rate
67 (Charlson et al., 1992). However, the long-term climatic effects lead to changes in aerosols
68 optical properties, and microphysical properties of clouds including their lifetime, formation and
69 precipitation (Twomey, 1997); and thereby, indirectly change terrestrial radiation. Further, many
70 studies have shown that human health can be affected by long-term exposure to fine PM
71 concentrations, particularly from $\text{PM}_{2.5}$ is associated with various diseases such as respiratory
72 tract infections, asthma, and lung diseases, that eventually leads to visibility degradation with
73 their increased concentration in the atmosphere (Pope et al., 2006).

74 With rapid economic development, urbanization, and industrialization, China has become
75 one of the most polluted regions in Asia during recent decades. Since 2013, the China National
76 Environmental Monitoring Center (CNEMC) has established numerous ground-based
77 observation stations to measure several air pollutant concentrations, which are of great
78 significance for the assessment of air quality in China (Hou et al., 2019). By the end of 2013, a
79 national air quality monitoring network with more than 1500 stations in China is put into the

80 operation, which provides air quality for the pollutants such as PM_{2.5}, PM₁₀, SO₂, NO₂, CO, O₃ to
81 the public. The earth's environment is constantly evolving, owing to the increasing human
82 population, and their corresponding activities leading to source emission of air pollutants. For
83 example industrial activities, land use and combustion of fossil fuels emit greenhouse gases and
84 aerosols, resulting in the change of atmospheric composition (Crowley et al., 2000) greatly due
85 to their day by day variations. Subsequently, these factors badly affect air quality, human health,
86 climate and environment (Tiwari et al., 2013; Zhang and Cao, 2015; Hou et al., 2019).

87 People from their fields studied the PM_{2.5} and PM₁₀ issues from various perspectives. A lot of
88 publications were involved in the spatial distribution and seasonal/diurnal changes of PM and
89 their association with aerosol optical depth (AOD) and key meteorological variables. For
90 example, Han et al. (2014) have shown that urbanization played a considerable part in PM_{2.5}
91 concentrations that eventually leads to a significant influence on public health and air quality.
92 Though much progress has been made, large discrepancies and knowledge gaps still exist among
93 PM_{2.5} and PM₁₀. However, several previous authors were noticed maximum concentrations in
94 their recent studies during the autumn and winter seasons in China. Zhang and Cao (2015)
95 noticed increased levels of PM_{2.5} concentration during autumn and winter seasons over East
96 China due to enhanced contribution from open biomass burning. Li et al. (2017) found higher
97 values of PM_{2.5} in autumn (85.5 $\mu\text{g m}^{-3}$) and winter (97.2 $\mu\text{g m}^{-3}$) compared to other seasons in
98 an urban region of Northeast China. Similar results were reported by Shao et al. (2017) in
99 Nanjing for the PM_{2.5} observations during 2013-2015. Further, the recent investigation by Wang
100 et al. (2017) on PM_{2.5} and PM₁₀ noticed high concentration in winter with the samples collected
101 at an urban site of Zhengzhou, could be associated with an increase in air pollutants (due to the
102 continuous usage of heating systems) and stable meteorological conditions. A recent study by

103 Hou et al. (2019) revealed that surface $PM_{2.5}$ concentrations were decreased during 2013-2018
104 over highly polluted areas due to a reduction in anthropogenic emissions, with stable values
105 found over the Pearl River Delta region.

106 Even though, the previous investigations were dedicated to a specific location based on the
107 large span of data covering all seasons and failed to present changes in PM concentrations to
108 understand the mechanisms involved in haze-fog phenomena during autumn and winter seasons.
109 Using the measured datasets provided by CNEMC, the present study aims to examine the
110 temporal variability of $PM_{2.5}$ and PM_{10} concentrations and source transport analysis from
111 multivariate statistical trajectory models such as potential source contribution function (PSCF)
112 and concentration weighted trajectory (CWT) in conjunction with the concentration bivariate
113 probability function (CBPF) analysis during autumn and winter of 2014-2017. This will allow
114 the authors to examine the role played by the atmospheric circulations in understanding the haze-
115 fog phenomena at an urban-industrial polluted city, Nanjing in the Yangtze River Delta (YRD)
116 region, East China. Further, the paper focused to reveal the relationship and causes of changes in
117 pollution between the aerosol optical properties and meteorological variables with the PM
118 concentrations over the region. At last, we derived and presented the clustering results obtained
119 from the identification and classification schemes to define different aerosol types, with a
120 perspective to give some brief analysis on spatial changes and trends in AOD using the satellite
121 data observed from the MODIS sensor.

122 **2. Data and meteorology**

123 **2.1. Experimental station**

124 The study area (Nanjing, 32.05° N, 118.78° E, 62 m above sea level) lies on the left bank
125 of the Yangtze River Delta (YRD) region and is located in the Jiangsu province, East China. The

126 city lies at an altitude of 83 m above sea level and is surrounded by several iron, steel and power
127 generation plants located within the proximity of ~5-7 km in the northeast of the measurement
128 site. Besides, the region suffers from both the heavy traffic and the agglomeration of heavy
129 industry. The climatological annual mean temperature and total precipitation in the region were
130 about 15-17° C and 1000-1200 mm, respectively, with a maximum occurrence during the
131 summer. However, the prevailing winds are southeasterly in summer and northeasterly in winter
132 derived from the Siberian anticyclone (Chen et al., 2018). More details about the layout of the
133 site, meteorology, and aerosol types and sources can be found elsewhere (Yu et al., 2016; Kang
134 et al., 2016a, b) and will be discussed in the following sections for the present database.

135 **2.2. Air pollutant data**

136 Real-time hourly monitoring data of PM_{2.5} and PM₁₀ mass concentrations in China between
137 September 2014 and February 2017 are available from the China air quality online monitoring
138 and analysis platform (<https://www.aqistudy.cn/>) and the Ministry of Environmental Protection
139 of China (<http://106.37.208.233:20035/>). All the measurement sites in Nanjing were equipped
140 with the Tapered Element Oscillating Microbalance (TEOM, RP1400 Model) instruments from
141 Thermo Scientific Company, USA to measure the main atmospheric air pollutant concentrations
142 at a resolution of 0.1 µg m⁻³ (Hou et al., 2019).

143 As a vertical integral, AOD not only reflects the impact of aerosols on the environment and
144 climate on the surface but also reflects the impact of aerosols at different heights of the boundary
145 layer as well, even in the upper and middle troposphere. In this study, the level 2 Collection 6.0
146 daily averaged AOD at 550 nm derived from the MODIS sensor onboard the Aqua satellite is
147 obtained from the NASA GIOVANNI (<https://giovanni.gsfc.nasa.gov/giovanni/>) for the period

148 during September 2014–February 2017. The downloaded MODIS AOD data are averaged into
149 the seasonal and annual AOD values and are used to supplement the measured pollutant data.

150 **2.3. Meteorological parameters**

151 Nanjing has a humid subtropical climate with hot summers and cold winters. Moderate to
152 low rates of air temperatures and precipitation were recorded during the autumn (from
153 September to November) and winter seasons (from December to January) throughout the study
154 period. The meteorological parameters include, air temperature (AT in °C), relative humidity
155 (RH in %), rainfall (RF in mm), wind speed (WS in m s^{-1}), and wind direction (WD in deg) to
156 some extent exert a strong influence on PM concentrations. The 24-h daily average
157 meteorological parameters for the period 2014-2017 were procured from
158 <http://wunderground.com>, a world-wide-web based meteorological variables for a geographical
159 area, and previously been used by other researchers (e.g., Kang et al., 2016a, b).

160 Fig. S1 of Supplementary Material (SM) shows the changes in air temperature, relative
161 humidity, and rainfall observed at Nanjing during autumn and winter for the period 2014-2017.
162 The monthly mean ambient air temperature found maximum in September 2016 (23 °C) and
163 minimum in January 2018 ($\text{AT} < 3$ °C) due to the frequent occurrence of haze and fog conditions
164 (Kang et al., 2016b) over the region. Whereas, the RH (relative humidity) was recorded highest
165 in October 2016 ($>85\%$) and lowest in December 2014 ($<59\%$) causing dry conditions attributed
166 to low water vapor and weak winds. It is obvious that both the parameters had depicted the
167 opposite trend in winter, with a similar variation in autumn. However, previous studies (Gao et
168 al., 2008, Yang et al., 2013) suggested that changes in meteorological conditions, regional
169 urbanization, and land use can cause the temperature to rise and RH to decrease, which result
170 mostly in more haze-fog conditions and less precipitation in China. It is also inferred from Fig.

171 S1 of SM that the precipitation rate reaches a maximum in the autumn of 2016 (> 195 mm), with
172 an increase of temperature. Moreover, it is revealed that the precipitation was found directly
173 related to the air temperature, with an increase of temperature results in an increase in the rate of
174 precipitation also, at the study site.

175 Fig. S1 of SM also presents the wind rose plots to analyze the relative frequency of wind
176 speed in autumn, winter, and autumn+winter at Nanjing for the period 2014-2017. The wind rose
177 shows the direction of wind blowing, while the color gives the magnitude of binned wind speed.
178 It is found that the direction of winds in autumn and winter seasons was dominated by the winds
179 coming from south and southeast directions. The strong (6 ms^{-1}) southeasterly winds bring air
180 masses, which may possess a large fraction of anthropogenic aerosols, significantly influence the
181 regional climate. Since the wind experienced at any given location is highly dependent on local
182 topography and meteorological factors.

183 3. Data processing methods

184 3.1. Spectral dependence of AOD and AE

185 A very brief description about the aerosol optical and physical parameters is given here
186 since several previous studies (e.g., Eck et al., 1999; Kaskaoutis et al., 2009; Kumar et al., 2013)
187 have described the physical significance, its use and derivation, which confirm to the applied
188 statistics and instrument algorithm. The spectrally-dependent AE parameter provides useful
189 information about the particle-size computed from the AOD using any pair of wavelengths given
190 by the Ångström law (Ångström, 1961);

$$191 \quad AOD_{\lambda} = \beta \lambda^{-AE} \quad (1)$$

192 where ' λ ' (in μm) is the wavelength. The Ångström turbidity coefficient (β) varied in the range
 193 0–1, and acts as an indicator of the number of aerosols present in the atmospheric column, being
 194 equal to AOD (λ) at $\lambda=1 \mu\text{m}$.

195 Also, the Volz's method can be applied to any pair of wavelengths in the spectral AOD for
 196 the computation of AE. It relies on the Junge power law and its value is a qualitative indicator of
 197 aerosol particle-size (Kumar et al., 2013) where $AE < 1.0$ and > 1.0 represents the dominance of
 198 coarse- and fine-mode, respectively in the aerosol size distribution. In the present study, AE was
 199 estimated in the range 440–870 nm from the following equation:

$$200 \quad AE = -\frac{d \ln AOD_{\lambda}}{d \ln \lambda} = \frac{\ln\left(\frac{AOD_{\lambda_1}}{AOD_{\lambda_2}}\right)}{\ln\left(\frac{\lambda_1}{\lambda_2}\right)} \quad (2)$$

201 where AOD_{λ_1} and AOD_{λ_2} represent AOD at two wavelengths λ_1 and λ_2 , respectively.

202 **3.2. Multivariate statistical trajectory models**

203 **3.2.1. HYSPLIT Trajectory analysis**

204 The analysis of the movement of air parcels can be easily quantified utilizing the trajectory
 205 model. This technique is much effective in identifying the transport carries source receptor from
 206 one place to the other following the forward and backward calculations. There are some
 207 limitations in the trajectories, which involves majorly grid resolution of input data. The
 208 reanalysis data, forecast data at various grid resolutions in both spatial and vertical, gives a
 209 deviation in the air mass trajectories. To defend this problem, there can be some other trajectory
 210 statistical analysis models such as PSCF and CWT (Wang et al., 2010) which are discussed in
 211 the upcoming sections. In the present study, air mass backward trajectories arriving at Nanjing
 212 originated at a height of 500 m above ground level ending at 12:00 h UTC with 120 h duration of

213 time. These trajectories were calculated every 6 h (00:00, 06:00, 12:00, and 18:00 h UTC) using
 214 the NOAA's version 4 of Hybrid Single-Particle Lagrangian Integrated Trajectory (HYSPLIT-4)
 215 model (Draxler and Rolph, 1997), which is a widely used model for analyzing and calculating
 216 the transport of air pollutants.

217 3.2.2. The PSCF model

218 The potential source areas can be identified by using the PSCF method, which combines the
 219 backward trajectory and defined values of air pollutants (Wang et al., 2009). The study field is
 220 divided into small equal grid cells (ij). The value of PSCF is expressed as:

$$221 \quad PSCF_{ij} = \frac{m_{ij}}{n_{ij}} \quad (3)$$

222 where i and j denote the latitude and longitude, respectively, n_{ij} represents the total number of
 223 endpoints passing through the ij cell, and m_{ij} is defined as the number of endpoints in the same
 224 cell associated with samples that are higher than the criterion value. The 75th percentile for each
 225 chemical species is selected as the criterion value for m_{ij} . The PSCF value represents a
 226 conditional probability describing the potential contribution of a grid cell of the high pollutant
 227 loadings at the receptor site. Cells with high PSCF values indicate areas of high potential
 228 contributions to the pollutant concentration at the receptor site, and the trajectories passing
 229 through these cells are the major transport pathways leading to the high pollutant loadings at the
 230 receptor site. To reduce the uncertainty in cells, a weighting function $w(n_{ij})$ should be multiplied
 231 with the PSCF value when n_{ij} is lower than three times an average number of trajectory
 232 endpoints (n_{mean}) in each cell (Zeng and Hopke, 1989). The weighted potential source
 233 contribution function (WPSCF) is described as follows:

$$234 \quad WPSCF_{ij} = \frac{m_{ij}}{n_{ij}} \times W(n_{ij}) \quad (4)$$

235 The weight function W_{ij} is defined as:

$$w_{ij} = \begin{cases} 1.00 & n_{ij} > 3 \cdot Avg \\ 0.70 & Avg < n_{ij} \leq 3 \cdot Avg \\ 0.42 & 0.5 \cdot Avg < n_{ij} \leq Avg \\ 0.17 & n_{ij} \leq 0.5 \cdot Avg \end{cases} \quad (5)$$

where Avg is the average

239 number of endpoints in each cell.

240 3.2.3. The CWT model

241 The CWT method, developed by Hsu et al. (2003), was used to make a distinction between
 242 strong sources from weak ones, considering the limitation of the PSCF method in which grid
 243 cells can have the same PSCF value when the sample concentrations are either only slightly
 244 higher or much higher than the criterion. In the CWT method, each grid cell is assigned a
 245 weighted concentration by averaging the sample concentrations that have associated trajectories
 246 that crossed the grid cell as follows:

$$C_{ij} = \frac{\sum_{h=1}^M C_h \times \tau_{ijh}}{\sum_{h=1}^M \tau_{ijh}} \times W(n_{ij}) \quad (6)$$

248 where C_{ij} is the mean weight concentration of the back trajectory 'h' in the ij cell; C_h represents
 249 $PM_{2.5}$ concentration in the trajectory 'h' through ij cell; τ_{ijh} represents the actual time that
 250 trajectory 'h' resides in the ij cell. Besides, $W(n_{ij})$ used in CWT is the same as that in PSCF to
 251 reduce the uncertainty in cells.

252 Both WPSCF and WCWT analyses were calculated using the MeteoInfo software-TrajStat
 253 Plugin (Wang et al., 2009), which has been proven useful to identify potential source areas of
 254 $PM_{2.5}$ and its chemical species.

255

256 **3.3. Discrimination of aerosol types**

257 An investigation of major aerosol types were identified over urban Nanjing site via., the
258 AOD₄₄₀ versus AE₄₄₀₋₈₇₀ cluster technique previously used in several studies (Kaskaoutis et al.,
259 2009; Bibi et al., 2016; Patel et al., 2017; Boiyo et al., 2018; Kumar et al., 2018). The method is
260 based on the sensitivity of the two wavelength-dependent parameters (AOD and AE) to
261 discriminate different aerosol types following the defined threshold value which varies
262 depending on the geographical locations (Pace et al., 2006). In the present work, AOD₄₄₀ and
263 AE₄₄₀₋₈₇₀ values were in the range of 0.1–1.9 and 0.4–1.8, respectively. Therefore, the values of
264 AOD < 0.15 and AE < 0.9 are represented as the clean maritime (MA) aerosol type. On the other
265 hand, AOD < 0.15 and AE > 1.0 is used for the polluted continental (PC) type of aerosols.
266 However, AOD > 0.3 and AE > 1.0 are characterized as the urban-industrial/biomass burning
267 (UI/BB), while, AOD > 0.6 with AE < 0.7 indicated the desert dust (DD) aerosol types. The
268 remaining gaps were considered as a mixed type of aerosols (MX). The MX aerosols are difficult
269 to be discriminated against, bearing in mind the distinct effects of various aerosol-mixing
270 processes in the atmosphere such as coagulation, condensation, humidification, and gas-to-
271 particle conversion (Kaskaoutis et al., 2009).

272 **3.4. Aerosol modification scheme**

273 The AE is dependent on the spectral band used for its estimation and reflects different
274 aerosol-types (Eck et al., 1999). At shorter wavelengths, it provides information regarding size
275 variation of fine-mode particles; while at longer wavelengths, it yields information about the
276 variation in fine- to the coarse-mode ratio (Kumar et al., 2013). The aerosol modification
277 processes at Nanjing were examined using the graphical scheme proposed by Gobbi et al. (2007)
278 and utilized by several authors (Kaskaoutis et al., 2011; Yu et al., 2016; Kang et al., 2016b; Patel

279 et al., 2017; Boiyo et al., 2018). The method relies on the combined analysis of $AE_{440-870}$ and its
280 spectral curvature represented by $dAE = AE_{440-675} - AE_{675-870}$ with fine-mode effective radii (R_f)
281 and FMF (η) as grid parameters in a grouped AOD. The graphical framework has been drawn
282 using Mie theory for a typical refractive index of air $m=1.4-0.001i$. More details concerning the
283 scheme and sensitivity analysis are discussed by Gobbi et al. (2007).

284 The change in AE (dAE) pairs with increasing $AOD_{675\text{ nm}}$ provides crucial information
285 regarding aerosol-modification processes in the atmosphere (i.e., cloud contamination, hydration,
286 and coagulation) (Kumar et al., 2014). Negative values of dAE are associated with negative
287 curvature (i.e., larger values of AE indicate the dominance of fine-mode particles). Positive
288 values are associated with positive curvature (i.e., smaller values of AE suggest the dominance
289 of coarse-mode particles associated with bimodal size distribution). A zero value of dAE
290 represents the absence of AE spectral variability (Kaskaoutis et al., 2011). Therefore, to
291 investigate the spectral variation of AE, its difference (dAE) at shorter (440–675 nm) and longer
292 (675–870 nm) wavelength are calculated and plotted against $AE_{440-870}$ as a function of AOD_{675} .
293 The aerosol modification processes in Nanjing were examined using daily averaged observations
294 of AE and AODs, with the latter (at 675 nm) represented by different colors and symbol size of
295 increasing turbidity.

296 **3. Results and discussion**

297 **3.1. Frequency distributions**

298 The frequency distributions of $PM_{2.5}$, PM_r , AOD_{440} , and $AE_{440-870}$ during the two seasons
299 for the entire study period are shown in Fig. 1. The bin interval in the present study was set to 0.1
300 for PM_r , AOD_{440} , and $AE_{440-870}$, and 10 for $PM_{2.5}$. However, we considered daily averaged
301 values of $PM_{2.5}$ and AOD_{440} in the range $0-260\ \mu\text{g m}^{-3}$, and $0-3.3$ respectively, while $0-1.8$ for

302 PM_r, and AE₄₄₀₋₈₇₀, respectively. The PM_{2.5} showed a wide unimodal distribution, with
303 significant seasonal heterogeneity (Fig. 1a) attributed to complex seasonal patterns (as for
304 thermal inversion and foggy conditions) of the atmosphere and variability in meteorological
305 phenomena (Che et al., 2014). The PM_{2.5} autumn season peaked at a bin interval of 20–40 $\mu\text{g m}^{-3}$
306 ³, accounting for 33.50% of the total distribution. On the other hand, it showed the strongest
307 mode in the bin interval of 40–60 $\mu\text{g m}^{-3}$ contributing 21.5%, signifying the dominance of
308 absorbing fine aerosol type during winter. Furthermore, PM_r distribution significantly skewed
309 towards larger values (Fig. 1b) in both seasons throughout the study period is more likely related
310 to large seasonal heterogeneity and massive biomass burning in the YRD. In autumn, almost
311 60% of the values were in the interval of 0.4–0.6; whereas, in winter, nearly 48% of the values
312 were found between 0.6 and 0.8, indicating that PM₁₀ concentrations were found considerably
313 higher than PM_{2.5}. In both the seasons, AOD₄₄₀ showed the strongest modes in the bin interval of
314 0.4–0.6 (autumn) and 0.6–0.8 (winter), accounting for 25.1% and 25.5%, respectively (Fig. 1c).
315 Meanwhile, the AE₄₄₀₋₈₇₀ also illustrated a sharp peak at larger values (>1.0), with its wide
316 distribution (Fig. 1d) similar to the results reported by Kang et al. (2016a) and Kumar et al.
317 (2018) over Nanjing using the MODIS satellite data. The occurrence of strong modes at
318 relatively higher size bins, with seasonal means > 1.0 (Table 1) during both the seasons interprets
319 more contribution of fine- relative to coarse-mode particles.

320 **3.2. Temporal changes in PM and AOD**

321 The seasonal mean AOD₄₄₀ was found high with 0.97 ± 0.31 and 0.86 ± 0.45 during
322 autumn and winter of 2015, respectively (Table 1). However, the annual mean values of AOD₄₄₀
323 for the entire study period was 0.74 ± 0.25 (Fig. 2), which is higher compared to that recorded at
324 other sites (Kaskaoutis et al., 2009; Kumar et al., 2013; Bibi et al., 2016; Patel et al., 2017; Boiyoy

325 et al., 2018). Further, the mean AOD₅₀₀ value of 0.65 ± 0.14 observed during winter reported by
326 Srivastava et al. (2014) at an urban Delhi is comparable with our findings obtained in this study.
327 Moreover, the MODIS AOD₅₅₀ was found close to 0.9, while the Cimel Sunphotometer (CSP)
328 measured AOD₄₄₀ showed a high value of 1.48, with a mean of 0.74 ± 0.25 for the study area.

329 The monthly mean precipitable water vapor content (PWC) (Fig. 2a) varied from 0.48 to
330 3.40 cm, with its seasonal mean found highest (1.58 ± 0.80 cm) in autumn during the study period
331 observed at Nanjing. The PWC was recorded very high (3.4 cm) in September 2014 with the
332 AOD and PWC almost followed a similar pattern, except in 2017 with each other, which
333 signifies the hydrophilic nature of the aerosol particles on the humid regions of East China. This
334 aspect implies that the higher amount of water content helped the growth of existing particles,
335 and subsequently results in the formation of new particles in the same vicinity, which in turn
336 results in higher AOD (Kumar et al., 2013).

337 The annual mean AE₄₄₀₋₈₇₀ was found 1.14 ± 0.13 for the study area (Fig. 2b). However,
338 AE was found low (< 1.0) during winter indicates the presence of mixed aerosols with the
339 dominance of fine- particles. Even, it increased to 1.1 in the other years of study attributed to
340 enhanced anthropogenic activities; and sometimes reached maximum values around 1.4 due to
341 intense biomass burning in the vicinity of the study region. However, the contribution from the
342 windblown mineral dust aerosols from the surrounding continents due to long-range transport
343 cannot be ruled out. Whereas, the mean value of β was estimated to be about 0.31 ± 0.12 during
344 the period 2014-2017 (Fig. 2b). The aerosol loading was extremely high (0.65) at the initial
345 period of study in September 2014 but suddenly dropped down to 0.15 during September 2017
346 resulting in decreased aerosol concentrations attributed to strict implementation of new
347 environmental policies in China (Hu et al., 2018), apart from the meteorological phenomena.

348 The monthly mean $PM_{2.5}$ was found moderate ($22 \mu\text{g m}^{-3}$) to high ($98 \mu\text{g m}^{-3}$), with an
349 annual mean of $58.51 \pm 21.03 \mu\text{g m}^{-3}$ over the study region during 2014-2017. We found
350 maximum values of annual $PM_{2.5}$ in winter ($81.2 \pm 41.5 \mu\text{g m}^{-3}$) compared to the autumn, with
351 monthly mean larger values noticed in January 2015 and smaller in October 2016. The increased
352 concentrations of $PM_{2.5}$ in winter are attributed to biomass and fossil fuel burning (Wang et al.,
353 2014; Chen et al., 2019). Also, less amount of precipitation and lower wind speed during winter
354 results in the maximum concentration of $PM_{2.5}$ over the study site. This results in the
355 deterioration of air quality over the region (Tao et al., 2014). Also, a high density of fire spots
356 was noticed closer to Nanjing resulting in high $PM_{2.5}$ and more amounts of absorption aerosols.
357 Apart from these, the urban, industrial and vehicle emissions play an important role for higher
358 PM concentrations in some cities (Cheng et al., 2012). The AQI follows a similar pattern as that
359 of $PM_{2.5}$ throughout the study; whereas, the visibility noticed opposite variation with $PM_{2.5}$
360 indicating an increase in $PM_{2.5}$ resulted in a decrease of visibility, and vice-versa. Furthermore,
361 the seasonal ratio in $PM_{2.5}/PM_{10}$ was found to be lower (0.53) in autumn and higher (0.60) in
362 winter seasons, with an annual mean of 0.57 (Table 1) attributed to the dominance of fine
363 particles. Sharma and Maloo (2005) reported similar values (~ 0.58) at three different sampling
364 locations in Kanpur, a highly polluted city in the northern part of India during the winter of
365 2002-2003. Tiwari et al. (2009) also reported the PM ratio (0.48 ± 0.21) varying between 0.18
366 (June) and 0.86 (February) across Delhi, suggested the dominance of fine-mode particles in
367 winter.

368 **3.3. Source analysis from trajectory models**

369 The PSCF and CWT models were used to reveal the potential origins of $PM_{2.5}$
370 geographically. To investigate the aerosol sources from its pathways during different seasons

371 arriving in Nanjing, the weighted analyses of PSCF and CWT models for $PM_{2.5}$ are performed on
372 a seasonal and annual basis (Fig. 3). Since, weighted-PSCF (WPSCF) is a method for estimating
373 the potential source locations for long-range transported pollutants, such as $PM_{2.5}$ (Tao et al.,
374 2014). However, the PSCF cluster analysis suggested that potential sources of $PM_{2.5}$ over the
375 study region are local as well as transported from distant places, mostly situated in north and
376 northwest of the observation site. This depends on the meteorological conditions and regional
377 anthropogenic activities that vary 0.1-0.5 with the season. The reason is most of the cities located
378 along the YRD are industrial as well as regional transport from these areas would have a
379 potentially high impact on the formation of $PM_{2.5}$ (Zong et al., 2018). Furthermore, the weighted-
380 CWT (WCWT) was also performed, to determine the possible emission sources causing $PM_{2.5}$
381 variability at the receptor site. The seasonal mean WCWT values of $> 40 \mu\text{g m}^{-3}$ and $> 80 \mu\text{g m}^{-3}$
382 revealed large influence of potential regional sources of aerosols originating from the north and
383 northwest of measurement site during autumn and winter, respectively; while the small source of
384 aerosols transporting from the East influenced with the WCWT of $< 20 \mu\text{g m}^{-3}$ was found on an
385 annual scale. This emphasizes the contribution of coarse particles, typically originated from the
386 upper parts of Jiangsu region located in East China. However, the air masses in autumn and
387 winter seasons were mostly originated from the northwestern directions, carrying similar types of
388 aerosols.

389 ***3.4. Relationship between surface and column aerosols***

390 The scatter plots of PM pollutant concentrations versus PM_{10} , PM_r , and AOD_{440} observed
391 at Nanjing during 2014-2017 are considered for correlation studies (Fig. 4). With the
392 corresponding linear regressions (using daily averages), significant strong correlations were
393 observed between $PM_{2.5}$ (0.94), $PM_{10-2.5}$ (0.82) and PM_{10} (Fig. 4a) indicates the dominant

394 sources of fine and inhalable particles in Nanjing have similar origins (such as direct motor
395 vehicle emissions and suspended road dust). It is revealed from the slope observed between
396 $PM_{2.5}$ and PM_{10} (0.67) indicates the PM ratio is dominant with fine-mode particles in Nanjing.
397 The regression analysis between mass concentrations of PM versus daily values of PMr is plotted
398 in Fig. 4b. Among the three PM pollutants, $PM_{2.5}$ exhibited a fair correlation of 0.54 with PMr
399 indicating a significant contribution of $PM_{2.5}$ in the PM ratio. It is observed that the $PM_{2.5}$ and
400 PM_{10} are positively correlated with PMr, except $PM_{10-2.5}$ which showed a negative relationship.
401 The correlation was found relatively low for PM_{10} (0.26) with PMr due to the high variability,
402 although it was slightly better (0.32) with AOD_{440} (Fig. 4c). However, both $PM_{2.5}$ (0.41) and
403 PM_{10} (0.32) showed positive correlations with AOD observed during the study period, with later
404 relatively weak than the former pollutant.

405 **3.6. Association between AOD, $PM_{2.5}$, and meteorology**

406 Fig. 5 illustrates the scatter plots to present the relationships between column AOD and
407 meteorological parameters. The intercept in the regression equation represents the background
408 air temperature and slope represents a coefficient used for converting AOD in temperature ($^{\circ}C$)
409 and relative humidity (%). However, Fig. 5a provides the correlation between RH and
410 temperature with intercept values of 63.80 and 10.61, respectively and correlation coefficients of
411 0.24 and 0.06, respectively during the study period over Nanjing. This is likely a result of the
412 common trends of human activities during this time. Another reason is that the site receives
413 usually heavy rainfall during the summer and autumn seasons that subsequently causes the
414 temperature to fall and increases RH.

415 Regression analysis between the precipitation and wind speed versus AOD_{440} is shown in
416 Fig. 5b. The resultant linear fitting has a slope of -3.62 and 0.09, and an intercept of 10.01 and

417 2.12 for precipitation and wind speed, respectively against AOD₄₄₀. Moreover, the correlation
418 between AOD₄₄₀ and precipitation was found weak (0.04). Also, a low intercept of 8.94 km was
419 observed between visibility versus AOD₄₄₀ (Fig. 5c), with a correlation coefficient of 0.43.
420 However, all the meteorological parameters were found negatively correlated with PM_{2.5} (Table
421 2), presenting the strong and weak correlations with visibility (-0.58) and rainfall (-0.09). In an
422 early study by Fu et al. (2014) reported that wind speed is an important factor influencing PM_{2.5}.
423 However, the wind speed can affect the accumulation and diffusion of aerosol particles; and the
424 quantity PM_{2.5} impinge on the visibility. The direction of the wind can affect the transport of
425 pollutants and can determine the spatial distribution of visibility (Tao et al., 2014). However, the
426 temporal trends of low visibility are mainly affected by the emissions of pollutants, aerosol
427 compositions and meteorological conditions (Che et al., 2014).

428 **3.7. Influence of winds on PM_{2.5} concentration**

429 The bivariate polar plots are shown in Fig. 6 presented the season-wise mean concentration
430 of PM_{2.5}. The wind speed and direction were used to identify the sources responsible for PM_{2.5}
431 concentrations. In both the seasons, the high PM_{2.5} concentrations were associated with lower
432 wind speeds indicating the presence of local pollutants. But for the combined seasonal analysis
433 (Fig. 6c) during the entire period of study, the highest concentrations were found mostly due to
434 strong winds representing the presence of local emissions as well as long-range transported
435 aerosols. In winter, air masses with a maximum concentration of particles are transported from
436 almost every quadrant i.e., with high wind speed from the northwest (6 ms⁻¹), and comparatively
437 low from northeast (2 ms⁻¹), southwest (4 ms⁻¹) and southeast (3 ms⁻¹) directions. Furthermore,
438 the air masses showed bi-directional pathways from the southwest and northeast in autumn.
439 However, PM_{2.5} had higher concentrations in all directions during the autumn and winter seasons

440 with low wind speed (< 4 m/s), which does not favor the dilution and dispersion processes of
441 pollutants. Most importantly, higher concentrations of $PM_{2.5}$ in winter are associated with
442 vehicular emissions. Eventually, the study area is also influenced by the highest $PM_{2.5}$
443 concentrations with the moderate breeze in autumn and winter seasons. This indicates the
444 influence of emissions from an industrial area as well (a few km distances away from the study
445 region) affecting the $PM_{2.5}$ concentrations. Similar results were also found for $PM_{2.5}$ over
446 Thiruvananthapuram (India) (Sumesh et al., 2018) during winter and autumn seasons, suggesting
447 that such a trend tends to be a general feature over India during the winter period with a large
448 contribution of industrial aerosols.

449 **3.9. Aerosol classification and modification processes**

450 *3.9.1. Identification of aerosol types*

451 Fig. 7 shows the scatter plot of AOD versus AE to represent the major aerosol types with
452 their seasonal percentage contributions at Nanjing. There is a wide range of AE values for low-
453 to-moderate values of AOD reflecting large variability in the aerosol properties and suggesting
454 several types of aerosols mixed in the atmosphere over Nanjing. The UI/BB type of aerosols was
455 found the most contributors ($>75\%$) reaching a maximum in winter (76.8%), indicating high
456 turbid conditions due to enhanced fine-mode particles produced from the burning of fossil fuel
457 and biomass (agricultural residue) or urban-industrial aerosols at the study area. A recent study
458 by Kumar et al (2018) has found about 60.9% of anthropogenic contribution with the dominance
459 of UI/BB aerosol type based on 12 years of MODIS AOD obtained over Nanjing. However, due
460 to the neighboring arid regions in South China and Asia (that emit large amounts of dust in
461 spring), the DD aerosols possess an important factor of 2.8% during autumn+winter (combined)
462 and comparatively less in winter ($\sim 0.9\%$). Furthermore, MA and PC aerosol types hold no

463 significant fraction over the study region; and no visible contribution has been found throughout
464 the study period at Nanjing. However, cases that do not belong to any of the above categories are
465 characterized as mixed (MX) types of aerosols were found to be 38.1%, 22.3%, 24.7% for the
466 autumn, winter, and autumn+winter, respectively.

467 3.9.2. Aerosol classification scheme

468 Over the entire study region, nearly all the points are within the classification scheme
469 represented as a function of AOD_{675} with different colors is shown in Fig. 7. The fine-mode
470 particles are dominated in autumn, while some coarse-mode particles are observed in winter with
471 a limited contribution during autumn+winter. This could be attributed to some dust particles
472 associated with prolonged dry and cold winters in Nanjing. During autumn, decreasing AOD
473 shows a shift to larger values of $AE_{440-870}$ (0.9–1.1) with R_f between 0.10 and 0.15 μm .
474 Furthermore, a similar trend has been found in winter which is inconsistent with the recent
475 findings of Patel et al. (2017) and Boiyo et al. (2018). However, with high AE (> 1.6), $\eta \sim 90\%$
476 and very low (< 0) values of dAE during winter interpret predominance of the same fine mode
477 particles indicates hygroscopic growth of fine mode aerosols or increase in fine mode particles
478 significantly due to cold breezes in winter (Kang et al., 2016b). The negative dAE indicates that
479 these aerosols are of the bimodal distribution having a large fine-mode fraction (between 0.2 and
480 0.15). While negative dAE values with $\eta > 70\%$, suggesting a predominance of fine-mode
481 particles, likely due to biomass burning activities and urban-industrial aerosols (Boiyo et al.,
482 2018). In contrast to the separate analysis of autumn and winter seasons, the combined
483 (autumn+winter) data set gives slightly different results with high AE (> 1.5), $\eta < 90\%$ and 0.2-
484 0.15 values of R_f , which interprets again the predominance of fine mode particles. However, a
485 slightly lower inhomogeneity with significant contributions of various aerosol types ($AOD \sim$

486 1.2–1.8) was found in winter and autumn plots. Similar results were also found over the Bay of
487 Bengal (Kaskaoutis et al., 2011) and urban Hyderabad (Sinha et al., 2012) for the same seasons.
488 It can, therefore, be interpreted that high AODs at Nanjing may be associated with increased
489 concentration of particularly, fine mode particles, and were significantly contaminated by clouds.
490 Moreover, the fine-mode particles are dominated in autumn and winter seasons particularly, the
491 coarse-mode particles are observed with a limited contribution during the autumn and winter
492 seasons.

493 **3.8. Spatial variations and trends in AOD and AE**

494 The optical parameter AOD is very useful for regional evaluation and investigating the
495 characteristics of aerosols, atmospheric pollution and environment (Eck et al., 1999). The spatial
496 distribution of AOD₅₅₀ showed >0.70 (Fig. 8) in autumn for the study area, was higher than that
497 in the surrounding area during the study period. However, higher values of AOD (≥ 0.7) were
498 found in the East of the study region (between 32°N to 34°N and 121°E to 122°E). In the center
499 towards the south part of Nanjing (between 26°N to 31°N and 116°E to 118°E), moderate aerosol
500 concentrations ($0.3 < \text{AOD} < 0.6$) were noticed, and clean environment conditions with AOD₅₅₀ <
501 0.3 were found often in the southeast of the study area. Further, it is revealed that the seasonal
502 averaged AOD in winter observed during 2014-2017 noticed higher than in autumn, and is more
503 widely distributed in Jiangsu province. The spatial distributions of AE showed for winter and
504 autumn seasons indicated a remarkably large amount of fine-mode aerosols in the north and west
505 of the study region (Nanjing) attributed largely from biomass burning, and mixed aerosol in
506 central and East of the study area, respectively.

507 The annual spatial trend distributions of AOD and AE (Fig. 9) are evaluated for two
508 different seasons and combined during 2014-2017. We found positive trends in the east of the

509 study region, with a negative trend in the north and other areas surrounding the study site. The
510 significant positive trends in winter for the study period around the site are attributed to large
511 anthropogenic activities and biomass burning. Even there are significant differences in the AE
512 spatial variability between seasonal and annual means. The AE was found almost higher during
513 the annual time than in the seasonal mean, especially during winter at the study site.

514 Fig. S3 of SM shows the spatial correlation analysis of AOD and AE values in eastern
515 China from 2014 to 2017. The spatial correlation values in autumn were found to be similar to
516 the spatial correlation between AOD and AE values in autumn+winter (combined) for the study
517 period. Moreover, most of the regions were found with a positive correlation of ~ 0.6 , ranging
518 from 0 to 0.7. However, the spatial correlation was found negative in winter near the study
519 region ranging from -0.5 to 0. Furthermore, our results are agreement with the findings of
520 Mamun et al. (2014) investigated over the northern and southern parts (Chittagong division) of
521 Bangladesh using the MODIS data from 2004 to 2013. They found the correlation coefficient of
522 -0.94 between AOD and AE and RMSE as 0.09.

523 4. Conclusions

524 Monitoring and estimating the air pollutant concentrations are essential for evaluating air
525 quality and its effects on the climate, environment, and human health. This study aims to
526 examine the potential characteristics of the surface measured PM ($PM_{2.5}$ and PM_{10}) concentration
527 and its association with column AOD and meteorological variables during autumn and winter
528 seasons of 2014-2017 over Nanjing in the YRD region, East China. The annual mean values of
529 $PM_{2.5}$ and PM_{10} in autumn (winter) were observed as $45.8 \pm 29.3 \mu\text{g m}^{-3}$ ($74.1 \pm 42.4 \mu\text{g m}^{-3}$) and
530 $83.5 \pm 45.1 \mu\text{g m}^{-3}$ ($120.9 \pm 56.6 \mu\text{g m}^{-3}$), respectively during the study period. The observations
531 showed with high values of $PM_{2.5}$ and PM_{10} exceeding the national air quality standards greatly

532 and significantly affecting the visibility, air quality, and climate system. Further, the mean ratio
533 of $PM_{2.5}/PM_{10}$ was about 0.57, with the lowest (0.53) observed in autumn and highest (0.60)
534 during winter for the entire period. This revealed the dominance of fine-mode particles at
535 Nanjing, attributed to anthropogenic activities (biomass burning, urban-industrial, vehicular
536 emissions, and other pollutants generated by construction activities) and conducive
537 meteorological phenomena. These resulting in serious pollution episodes at different locations in
538 the YRD are likely related to several anthropogenic activities.

539 Different statistical techniques include the effect and temporal variation of winds through
540 the CBPF analysis was used to analyze $PM_{2.5}$ in the study area. The bivariate plot results showed
541 an increase in $PM_{2.5}$ concentration attributed to the presence of local anthropogenic emissions
542 and favorable meteorological conditions over the study region. Apart from this, the source
543 identification was carried out from the PSCF and CWT models utilizing the air mass trajectories
544 obtained from the HYSPLIT model. The results showed that the regional transport of smoke
545 particles from biomass-burning resulting in high surface PM concentration and column AOD.
546 The present study also aims to investigate spatiotemporal evolution and trend in the aerosol
547 optical properties, qualitatively identify different aerosol types and sources in Nanjing. For this
548 purpose, the Collection 6 Level-2 data obtained from the MODIS sensor onboard Terra and
549 Aqua satellites for the period between 2014 and 2017 have been analyzed. A notable
550 spatiotemporal heterogeneity was observed in the optical properties of aerosols on the seasonal
551 scale over Nanjing. The seasonal mean AOD_{440} ($AE_{470-870}$) was found to be maximum with 0.97
552 ± 0.31 during autumn (1.14 ± 0.23) and a minimum of 0.62 ± 0.34 (1.29 ± 0.19) in winter.
553 However, the spatial correlation between AOD and AE was found to be negative during winter,
554 with the correlation coefficient ranging from -0.5 to 0. The results derived in this study

555 contribute to an in-depth understanding of PM concentrations during autumn and winter seasons
556 in urban Nanjing and forms a basis for the extension of future research over this aerosol hotspot
557 region (YRD) in China.

558 **Conflicts of interest**

559 The authors declare and confirm that we don't have any competing financial interest concerning
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578 **References**

- 579 Adesina, A.J., Piketh, S., Kumar, K.R., Venkataraman, S., 2017. Characteristics of columnar
580 aerosol optical and microphysical properties retrieved from the sunphotometer and its
581 impact on radiative forcing over Skukuza (South Africa) during 1999–2010. *Environ Sci*
582 *Pollut Res.* 24, 16160-16171
- 583 Ångström, A., 1961. Techniques of determining the Turbidity of the Atmosphere. *Tellus XIII*,
584 214-223.
- 585
586 Bibi, S., Alam, K., Chishtie, F., Bibi, H., 2017. Characterizing of absorbing aerosol types using
587 ground and satellite-based observations over an urban environment. *Atmospheric*
588 *Environment* 150, 126-135.
- 589
590 Boiyo R, Kumar KR, Zhao T. 2018. Spatial variations and trends in AOD climatology over East
591 Africa during 2001–2015: A comparative study using three satellite datasets. *International*
592 *Journal of Climatology*. <https://doi.org/10.1002/joc.5446>.
- 593
594 Boiyo, R., Kumar, K.R., Zhao, T., Guo, J., 2019. A 10-year record of aerosol optical properties
595 and radiative forcing over three environmentally distinct AERONET sites in Kenya, East
596 Africa. *Journal of Geophysical Research-Atmospheres* 124, 1596-1617.
- 597
598 Carabali, G., Estevez, H.R., Valdes-Barron, M., et al., 2017. Aerosol climatology over the
599 Mexico City basin: Characterization of optical properties. *Atmospheric Research* 194, 190-
600 201.
- 601
602 Charlson RJ, Schwartz SE, Hales JM, Cess RD, Coakley JA, Hansen JE, Hofmann DJ. 1992.
603 *Climate Forcing by Anthropogenic Aerosols. Science.* 255: 423–430. [https://](https://doi.org/10.1126/science.255.5043.423)
604 doi.org/10.1126/science.255.5043.423.
- 605
606 Che, H., Xia, X., Zhu, J., Li, Z., Dubovik, O., Holben, B.N., et al., 2013. Column aerosol optical
607 properties and aerosol radiative forcing during a serious haze-fog month over the North
608 China Plain in 2013 based on ground-based Sunphotometer measurements. *Atmos. Chem.*
609 *Phys.* 14, 2125–2138.
- 610
611 Che, H., Zhang, X. Y., Xia, X., Goloub, P., Holben, B., Zhao, H., Wang, Y., et al., 2015.
612 Ground-based aerosol climatology of China: aerosol optical depths from the China Aerosol
613 Remote Sensing Network (CARSNET) 2002-2013. *Atmospheric Chemistry and Physics*
614 15, 7619-7652.
- 615
616 Chen, G., Morawska, L., Zhang, W., Li, S., Cao, W., Ren, H., Wang, B., Wang, H., Knibbs,
617 L.D., Williams, G., Guo, J., Guo, Y., 2018. Spatiotemporal variation of PM₁ pollution in
618 China. *Atmospheric Environment* 178, 198-205.
- 619
620

- 621 Chen, P., Kang, S., Yang, J., Pu, T., Li, C., Guo, J., Tripathi, L., 2019. Spatial and temporal
622 variations of gaseous and particulate pollutants in six sites in Tibet, China during 2016-
623 2017. *Aerosol and Air Quality Research* 19, 516-527.
624
- 625 Eck, T.F., Holben, B.N., Reid, J.S., Dubovik, O., Smirnov, A., O'Neill, N. T., et al., 1999.
626 Wavelength dependence of the optical depth of biomass burning, urban, and desert dust
627 aerosols. *Journal of Geophysical Research*. <https://doi.org/10.1029/1999JD900923>.
628
- 629 Gobbi, G.P., Kaufman, Y.J., Koren, I., Eck, T.F., 2007. Classification of aerosol properties
630 derived from AERONET direct sun data. *Atmospheric Chemistry and Physics Discussions*,
631 6, 8713–8726. <https://doi.org/10.5194/acpd-6-8713-2006>.
632
- 633 Han, L. J., W. Q. Zhou, W. F. Li, L. Li, 2014. Impact of urbanization level on urban air quality:
634 A case of fine particles (PM_{2.5}) in Chinese cities. *Environ.Pollut.*, 194, 163-170.
635
- 636 Hou, X., Zhu, B., Kumar, K.R., Lu, W., 2019. Inter-annual variability in fine particulate matter
637 pollution over China during 2013-2018: Role of meteorology. *Atmospheric Environment*
638 214, 116842.
639
- 640 IPCC, 2013: Climate Change. 2013: The Physical Science Basis. Contribution of Working
641 Group I to the Fifth 4 Assessment Report of the Intergovernmental Panel on Climate
642 Change [Stocker, T. F., D. Qin, G.-K.5 Plattner, M. Tignor, S. K. Allen, J. Boschung, A.
643 Nauels, Y. Xia, V. Bex and P. M. Midgley (Eds.)].6 Cambridge University Press,
644 Cambridge, United Kingdom And New York, NY, USA, 1535. [https://doi.org/ 10.1017/CBO9781107415324](https://doi.org/10.1017/CBO9781107415324).
645
646
- 647 Kalluri, R.O.K., Gugamsetty, B., Kotalo, R.G., Nagireddy, S.K.R., et al., 2016. Direct radiative
648 forcing properties of atmospheric aerosols over semi-arid region, Anantapur in India.
649 *Science of the Total Environment* 566-567, 1002-1013.
650
- 651 Kang, N., Kumar, K.R., Hu, K., Yu, X., and Yin, Y., 2016a. Long-term (2002–2014) evolution
652 and trend in Collection 5.1 Level-2 aerosol products derived from the MODIS and MISR
653 sensors over the Chinese Yangtze River Delta. *Atmospheric Research*. 181, 29–43.
654
- 655 Kang, N., Kumar, K.R., Yu, X., Yin, Y., 2016b. Column-integrated aerosol optical properties
656 and direct radiative forcing over the urban-industrial megacity Nanjing in the Yangtze
657 River Delta, China. *Environmental Science and Pollution Research* 23, 17532-17552.
658
- 659 Kaskaoutis, D.G., Badarinath, K.V.S., Kharol, S.K., Sharma, A.R., Kambezidis, H.D., 2009.
660 Variations in the aerosol optical properties and types over the tropical urban site of
661 Hyderabad, India. *Journal of Geophysical Research-Atmospheres*, 114(22), 1–20.
662
- 663 Kaskaoutis, D.G., Kambezidis, H.D., Hatzianastassiou, N., Kosmopoulos, P.G., Badarinath, K.
664 V. S., 2007. Aerosol climatology: dependence of the Angstrom exponent on wavelength
665 over four AERONET sites. *Atmos. Chem. Phys. Discuss.* 7, 7347–7397.
666 <https://doi.org/10.5194/acpd-7-7347-2007>.

- 667
668 Kaskaoutis, D.G., Kharol, S.K., Sinha, P.R., Singh, R.P., Kambezidis, H.D., Rani Sharma, A.,
669 Badarinath, K.V.S., 2011. Extremely large anthropogenic-aerosol contribution to total
670 aerosol load over the Bay of Bengal during winter season. *Atmospheric Chemistry and*
671 *Physics*.11, 7097–7117. <https://doi.org/10.5194/acp-11-7097-2011>.
672
- 673 Kumar KR, Kang N, Yin Y. 2018. Classification of key aerosol types and their frequency
674 distributions based on satellite remote sensing data at an industrially polluted city in the
675 Yangtze River Delta, China. *International Journal of Climatology*. 38(1): 320-336. [https://](https://doi.org/10.1002/joc.5178)
676 [doi.org/ 10.1002/joc.5178](https://doi.org/10.1002/joc.5178).
677
- 678 Kumar, K.R., Kang, N., Sivakumar, V., Griffith D., 2017. Temporal characteristics of columnar
679 aerosol optical properties and radiative forcing (2011-2015) measured at AERONET's
680 Pretoria_CSIR_DPSS site in South Africa. *Atmospheric Environment*. 165, 274-289.
681
- 682 Kumar, K.R., Sivakumar, V., Reddy, R.R., Gopal, K.R., Adesina, A.J., 2013. Inferring
683 wavelength dependence of AOD and Ångström exponent over a sub-tropical station in
684 South Africa using AERONET data: Influence of meteorology, long-range transport and
685 curvature effect. *Science of the Total Environment*. 461–462, 397–408.
686
- 687 Kumar, K.R., Sivakumar, V., Yin, Y., Reddy, R.R., Kang, N., Diao, Y., Adesina, A.J., Yu, X.,
688 2014. Long-term (2003-2013) climatological trends and variations in aerosol optical
689 parameters retrieved from MODIS over three stations in South Africa. *Atmos. Environ*. 95,
690 400–408.
691
- 692 Kumar, K.R., Yin, Y., Sivakumar, V. N., Yu., X., Diao Y., Adesina A. J., Reddy, R.R., 2015.
693 Aerosol climatology and discrimination of aerosol types retrieved from MODIS, MISR and
694 OMI over Durban (29.88°S, 31.02°E), South Africa. *Atmospheric Environment*. 117, 9–18.
695
- 696 Pace, G., Sarra, A., Meloni, D., Piacentino, S., Chamard, P., 2006. Aerosol Optical Properties at
697 Lampedusa (Central Mediterranean). Influence of Transport and Identification of Different
698 Aerosol Types. *Atmospheric Chemistry and Physics Discussions*. 5: 4929–4969. [https://](https://doi.org/10.5194/acpd-5-4929-2005)
699 doi.org/10.5194/acpd-5-4929-2005.
700
- 701 Patel, P.N., Dumka, U.C., Kaskaoutis, D.G., Babu, K.N., Mathur, A.K., 2017. Optical and
702 radiative properties of aerosols over Desalpar, a remote site in western India: Source
703 identification, modification processes and aerosol type discrimination. *Science of the Total*
704 *Environment*, 575, 612–627.
705
- 706 Ramanathan V, Crutzen PJ, Kiehl JT, Rosenfeld D. 2001. Aerosols, Climate, and the
707 Hydrological Cycle. *Science* (New York, N.Y.). 294: 2119–2124. [https://](https://doi.org/10.1126/science.1064034)
708 doi.org/10.1126/science.1064034.
709
- 710 Rosenfeld D. 2000. Suppression of Rain and Snow by Urban and Industrial Air Pollution.
711 *Science* 287: 1793–1796. [https:// doi.org/10.1126/science.287.5459.1793](https://doi.org/10.1126/science.287.5459.1793).
712

- 713 Schuster GL, Dubovik O, Holben BN, 2006. Angstrom exponent and bimodal aerosol size
714 distributions. *J Geophys Res Atmos* 111(D7), doi:10.1029/2005JD006328.
715
- 716 Sinha, P.R., Kaskaoutis, D.G., Manchanda, R.K., and Sreenivas, S. (2012). Characteristics of
717 aerosols over Hyderabad in southern Peninsular India: synergy in the classification
718 techniques. *Anna. Geophys.* 30: 1393–1410.
719
- 720 Srivastava, A.K., Yadav, V., Pathak, V., Singh, S., Tiwari, S., Bisht, D., Goloub, T., 2014.
721 Variability in radiative properties of major aerosol types: A year long study over Delhi-An
722 urban station in Indo-Gangetic Basin. *Science of Total Environment* 473-474, 659-666.
723
- 724 Sumesh, R.K., Rajeevan, K., Resmi, E.A., Unnikrishnan, C. K., 2017. Particulate matter
725 concentrations in the southern tip of India: Temporal variations, meteorological influences,
726 and source identification. *Earth System Environment* 13, 1-18.
727
- 728 Tao, J., Gao, J., Zhang, L., Zhang, R., Che, H., Zhang, Z., Lin, Z., Jing, J., Cao, J., Hsu, S.C.,
729 2014. PM_{2.5} pollution in a megacity of southwest China: Source apportionment and
730 implication. *Atmospheric Chemistry and Physics* 14, 8679-8699.
731
- 732 Tiwari, S., Hopke, P.K., Pipal, A.S., Srivastava, A.K., Bisht, D.S., Tiwari, S., Singh, A.K., Soni,
733 V.K., Attri, S.D., 2015. Intra-urban variability of particulate matter (PM_{2.5} and PM₁₀) and
734 its relationship with optical properties of aerosols over Delhi, India. *Atmospheric Research*
735 166, 223-232.
736
- 737 Wang, Y.G., Ying, Q., Hu, J.L., Zhang, H.L., 2014. Spatial and temporal variations of six criteria
738 air pollutants in 31 provincial capital cities in China during 2013-2014. *Environment*
739 *International* 73, 413-422.
740
- 741 Wang, Y.Q., Zhang, X.Y., Draxler, R.R., 2009. TrajStat: GIS-based software that uses various
742 trajectory statistical analysis methods to identify potential sources from long-term air
743 pollution measurement data. *Environmental Modeling Software* 24, 938-939.
744
- 745 Yu, X., Kumar, K.R., Lü, R., Ma, J., 2016a., Changes in column aerosol optical properties during
746 extreme haze-fog episodes in January 2013 over urban Beijing. *Environmental Pollution*,
747 210, 217–226. <https://doi.org/10.1016/j.envpol.2015.12.021>.
748
- 749 Yu, X., Lu, R., Kumar, K.R., Ma, J., Zhang, Q., Jiang, Y., Kang, N., Yang, S., Wang, J., Li, M.,
750 2016b. Dust aerosol properties and radiative forcing observed in spring during 2001–2014
751 over urban Beijing, China. *Environ. Sci. Pollut. Res.* 23, 15432-15442.
752
- 753 Zhang J, Reid JS. 2010. A decadal regional and global trend analysis of the aerosol optical depth
754 using a data-assimilation grade over-water MODIS and Level 2 MISR aerosol products.
755 *Atmospheric Chemistry and Physics*. 10, 10949–10963.
756
- 757 Zhang, Y. L., F. Cao, 2015. Fine particulate matter (PM_{2.5}) in China at a city level. *Sci. Rep.*, 5,
758 14884.

Table 1. Inter-annual variations of PM concentrations and aerosol optical properties observed at Nanjing during autumn and winter seasons of 2014-2017.

Year	PM _{2.5}		PM ₁₀		PM _{10-2.5}		PM _{2.5} /PM ₁₀		AOD ₄₄₀		AE ₄₄₀₋₈₇₀	
	Aut	Win	Aut	Win	Aut	Win	Aut	Win	Aut	Win	Aut	Win
2014	58.9±27.5	78.6±41.6	111.5±47.6	135.6±57.1	48.1±19.2	58.1±27.7	0.59±0.10	0.57±0.14	0.94±0.44	0.72±0.30	1.05±0.20	1.02±0.21
2015	47.9±28.2	81.2±41.5	87.2±43.2	128.0±52.2	39.4±20.0	46.8±20.9	0.54±0.12	0.62±0.13	0.97±0.31	0.86±0.45	1.14±0.23	0.96±0.29
2016	33.7±19.1	61.7±30.9	65.8±34.5	101.5±46.5	32.1±17.4	39.8±19.3	0.51±0.10	0.60±0.10	-	0.67±0.31	-	1.30±0.22
2017	35.2±21.9	75.7±50.4	69.6±38.9	118.3±64.1	34.3±20.2	42.6±24.4	0.50±0.12	0.63±0.15	0.60±0.32	0.62±0.34	1.25±0.19	1.29±0.19
Mean	45.8±29.3	74.1±42.4	83.5±45.0	120.9±56.6	37.7±19.8	46.8±24.2	0.53±0.12	0.60±0.13	0.84±0.26	0.72±0.22	1.15±0.14	1.14±0.20

Table 2. Statistics obtained from the linear regression fitting between $PM_{2.5}$ and meteorological parameters during the study period. The symbols m, c, r, RMSE, n, and p represents slope, intercept, Pearson's coefficient, root mean square error, total data points, and probability value to test significance of data, respectively. The respective p-values <0.05 and >0.05 indicates significance of data at 95% confidence level and least significant.

Parameters	$PM_{2.5}$ vs Temp	$PM_{2.5}$ vs RH	$PM_{2.5}$ vs WS	$PM_{2.5}$ vs RF	$PM_{2.5}$ vs Vis
m	-0.056	-0.032	-0.011	-0.045	-0.048
c	14.791	76.603	2.985	8.119	8.681
r	-0.28	-0.09	-0.37	-0.22	-0.58
RMSE	7.598	13.251	0.986	7.866	2.704
n	725	725	725	211	725
p	<0.05	<0.05	<0.05	>0.05	<0.05

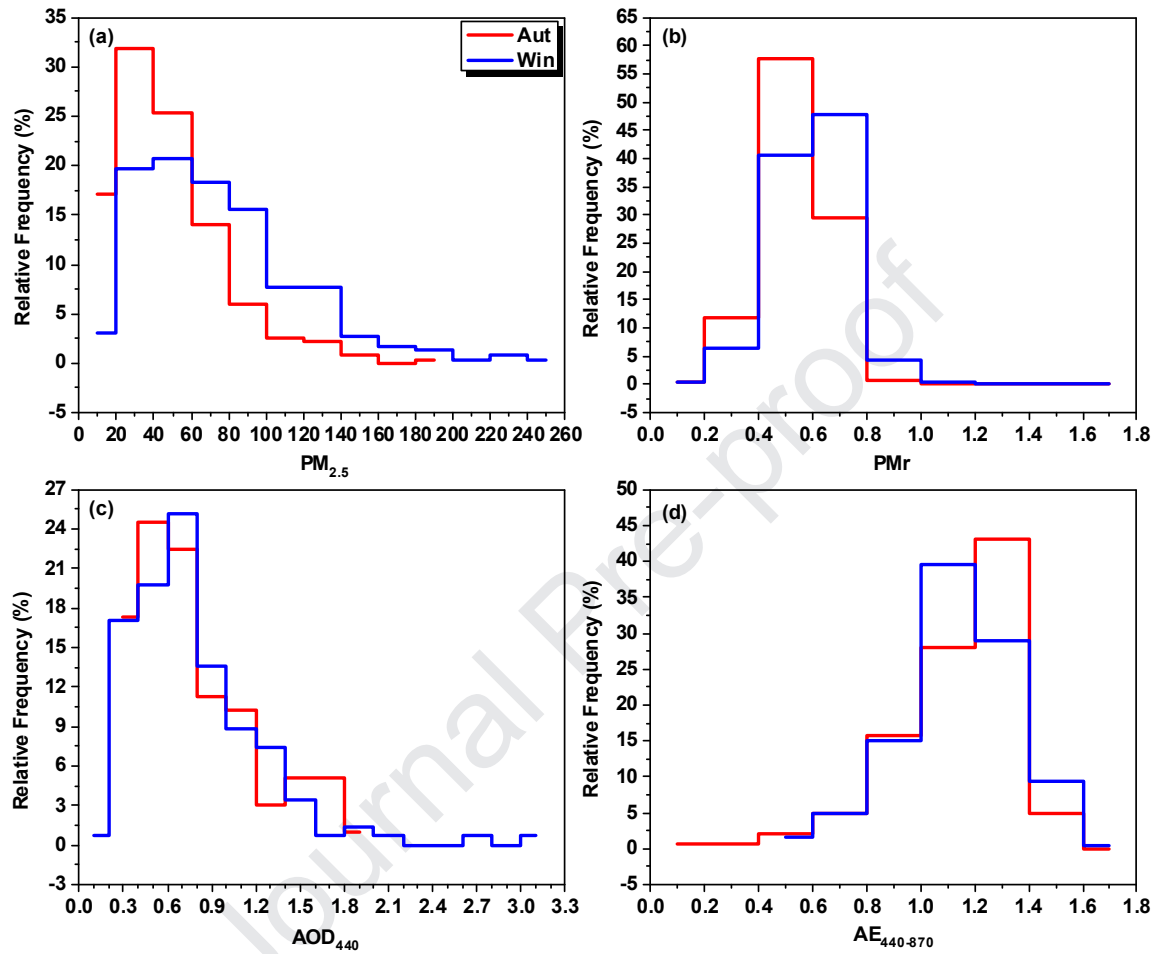


Fig. 1. Relative frequency distributions in measured (a, b) surface and (c, d) column aerosol properties at Nanjing during autumn and winter seasons.

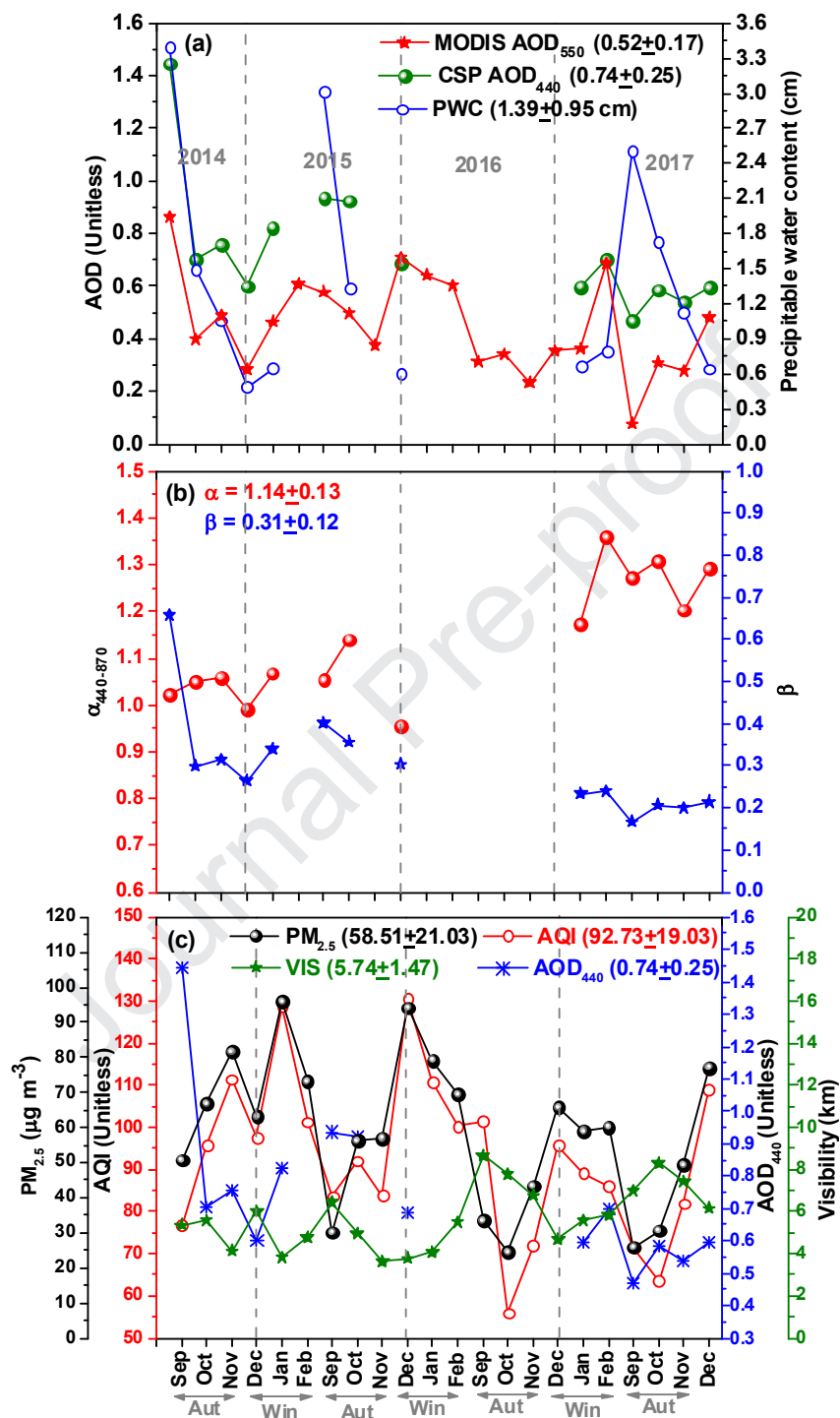


Fig. 2. Temporal variability in (a, b) aerosol optical properties measured from the Sun photometer and surface measured PM_{2.5} concentration at Nanjing during the study period. AQI and visibility are also presented in panel (c) to show their relationship with PM_{2.5} and AOD₄₄₀.

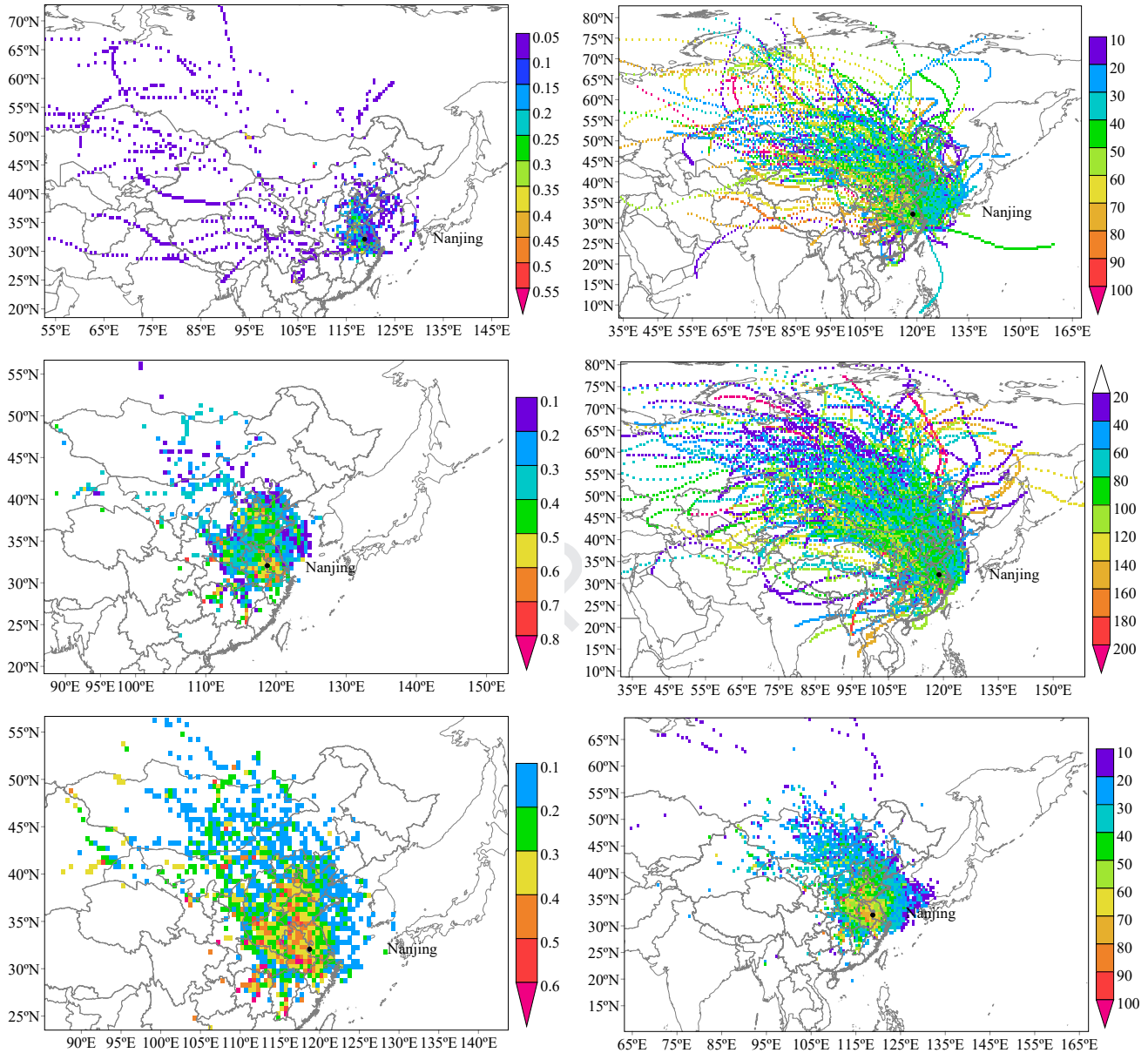


Fig. 3. WPSCF (left) and WCWT (right) maps to identify the sources of PM_{2.5} derived from the PSCF and CWT models arriving Nanjing at a height of 500 m above ground level. The maps are shown for (top) autumn, (middle) winter, and (bottom) all seasons during 2014-2017. The color corresponds to the PM_{2.5} pollutant concentration.

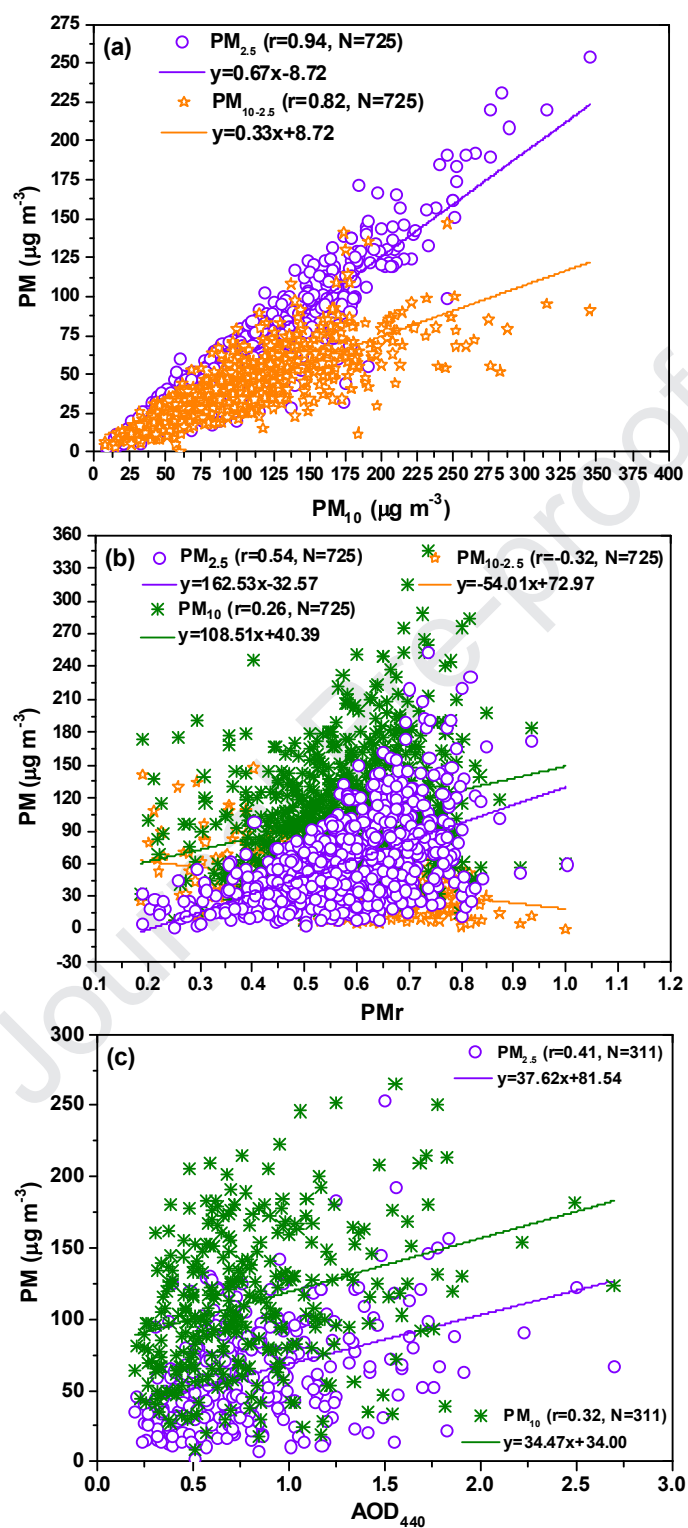


Fig. 4. (a, b) Scatter plots shown for different PM pollutant concentrations and its ratio. (c) Relationship between surface and column aerosol properties observed at Nanjing for the entire study period. The solid lines correspond to the straight lines obtained from the regression fitting. Also, the regression equation, Pearson's correlation coefficient (r) and number of paired data points (N) are given in the panels.

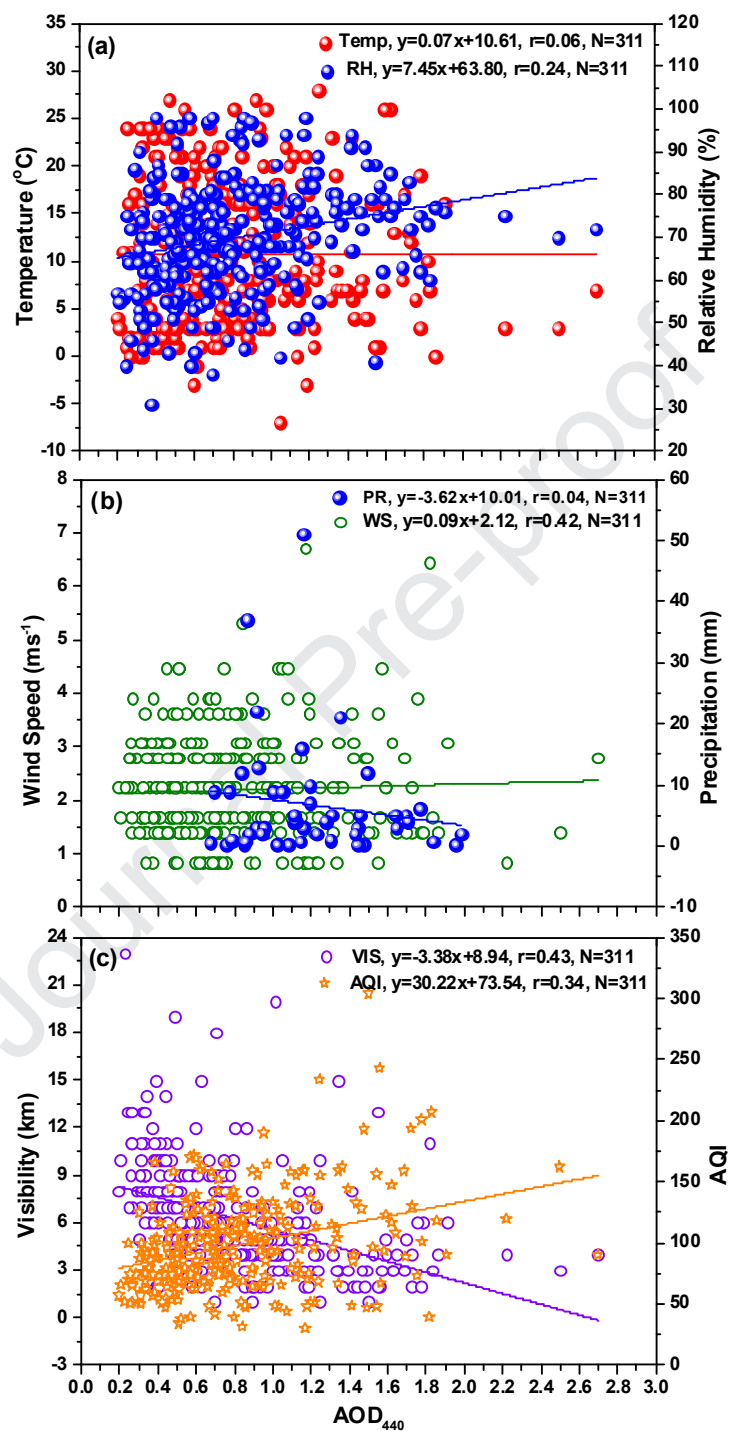


Fig. 5. Association of AOD_{440} with the ground-based measured major meteorological parameters at Nanjing. The regression analysis and its statistics are also presented.

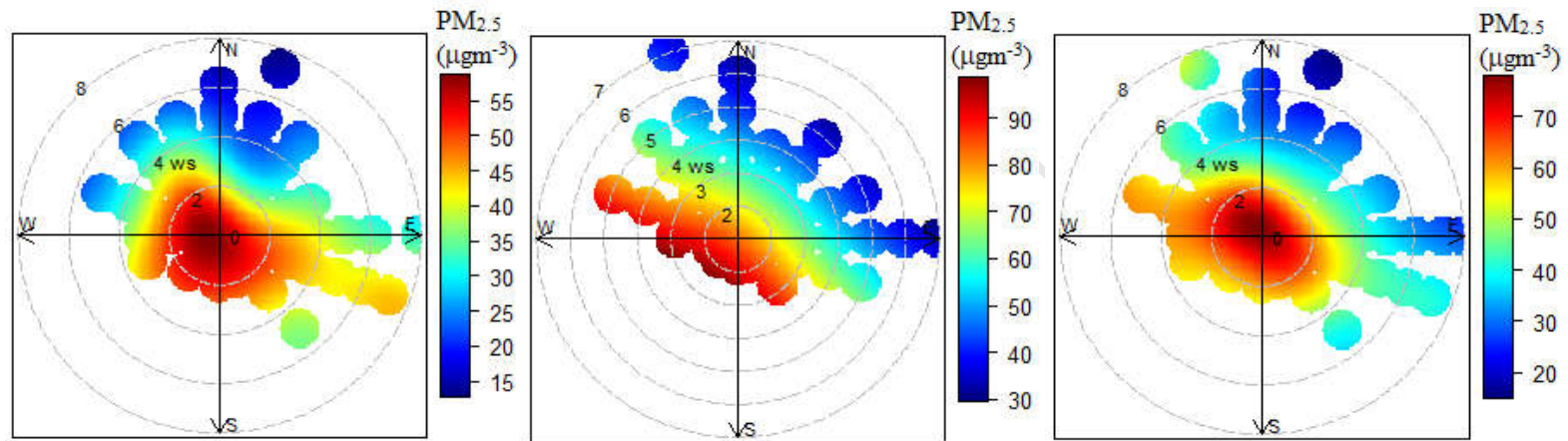


Fig. 6. Concentration bivariate probability function (CBPF) plots of PM_{2.5} concentration during (left) autumn, (middle) winter, and (right) all seasons of 2014-2017. The contour represents the concentration of PM_{2.5} corresponding to the wind speed and direction.

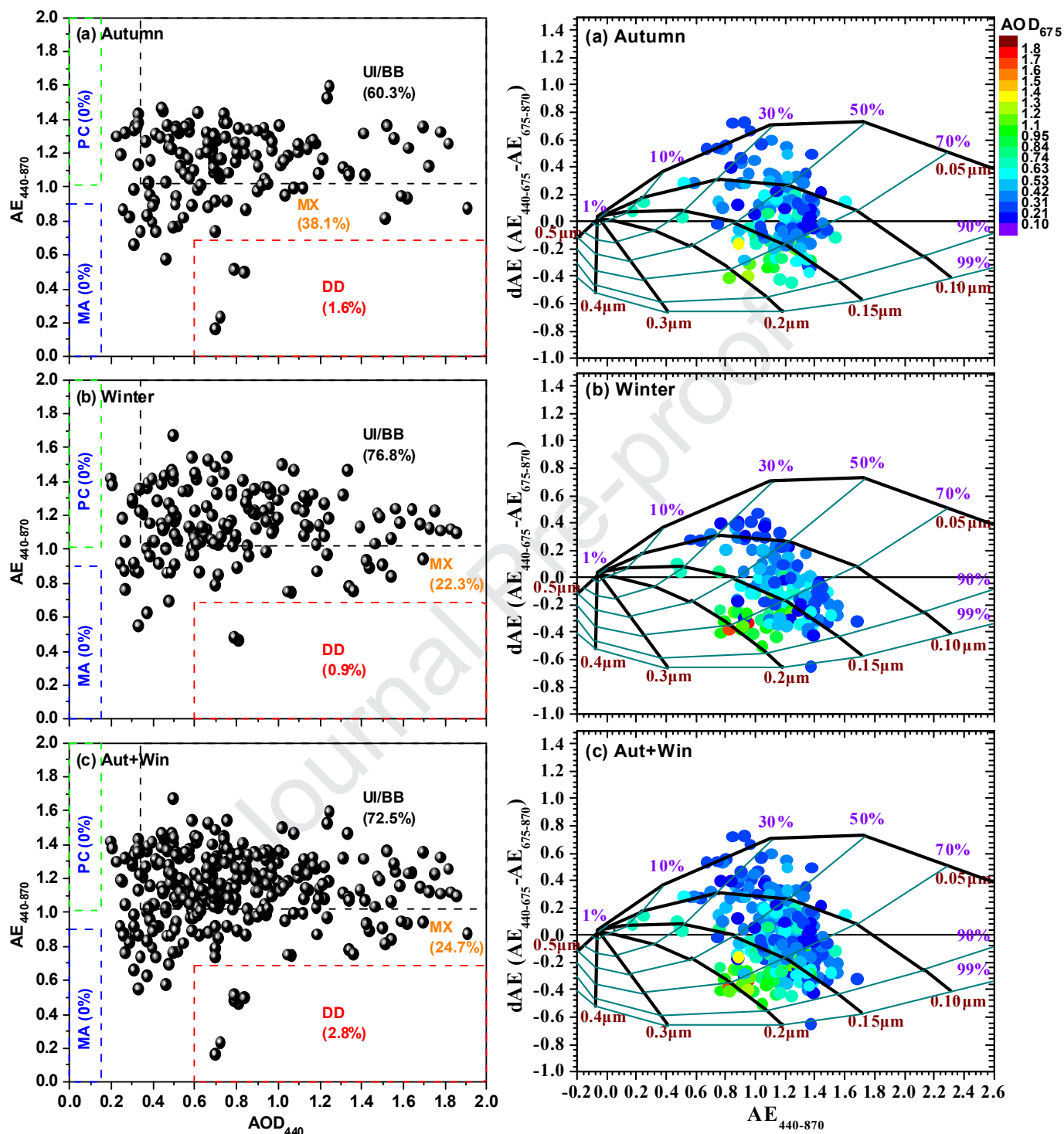


Fig. 7. Aerosol type discrimination (left) and its modification process (right) techniques observed at Nanjing from Sun photometer data during 2014-2017. The major aerosol types are identified and their percentage contributions are also presented.

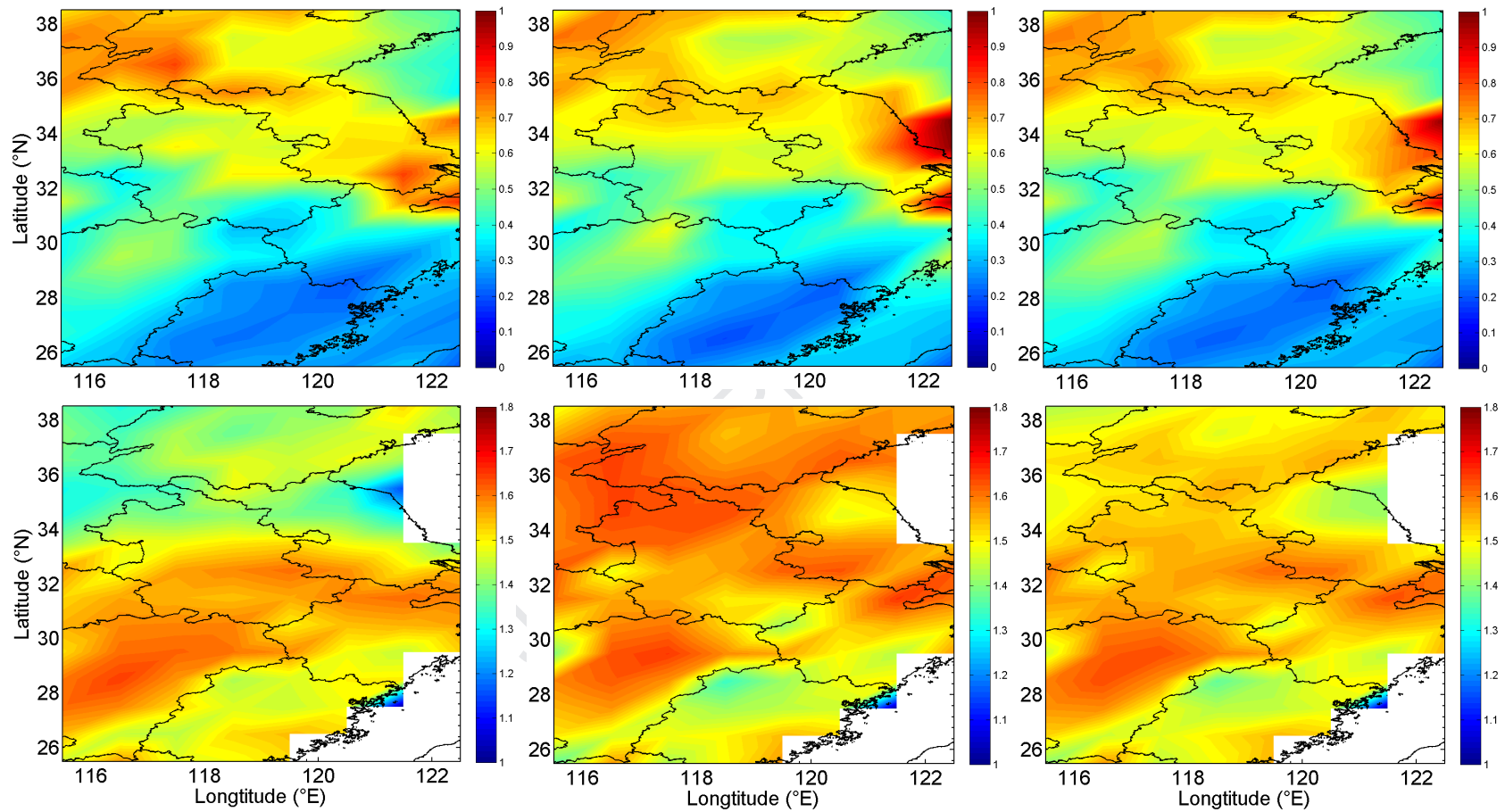


Fig. 8. Spatial distributions of MODIS Aqua satellite-derived AOD₅₅₀ (top panels) and AE₄₄₀₋₆₆₀ (bottom panels) during 2014-2017. The first, middle, and last panels in two rows corresponds to the autumn, winter and autumn+winter, respectively observed over East China. There are no units for AOD and AE.

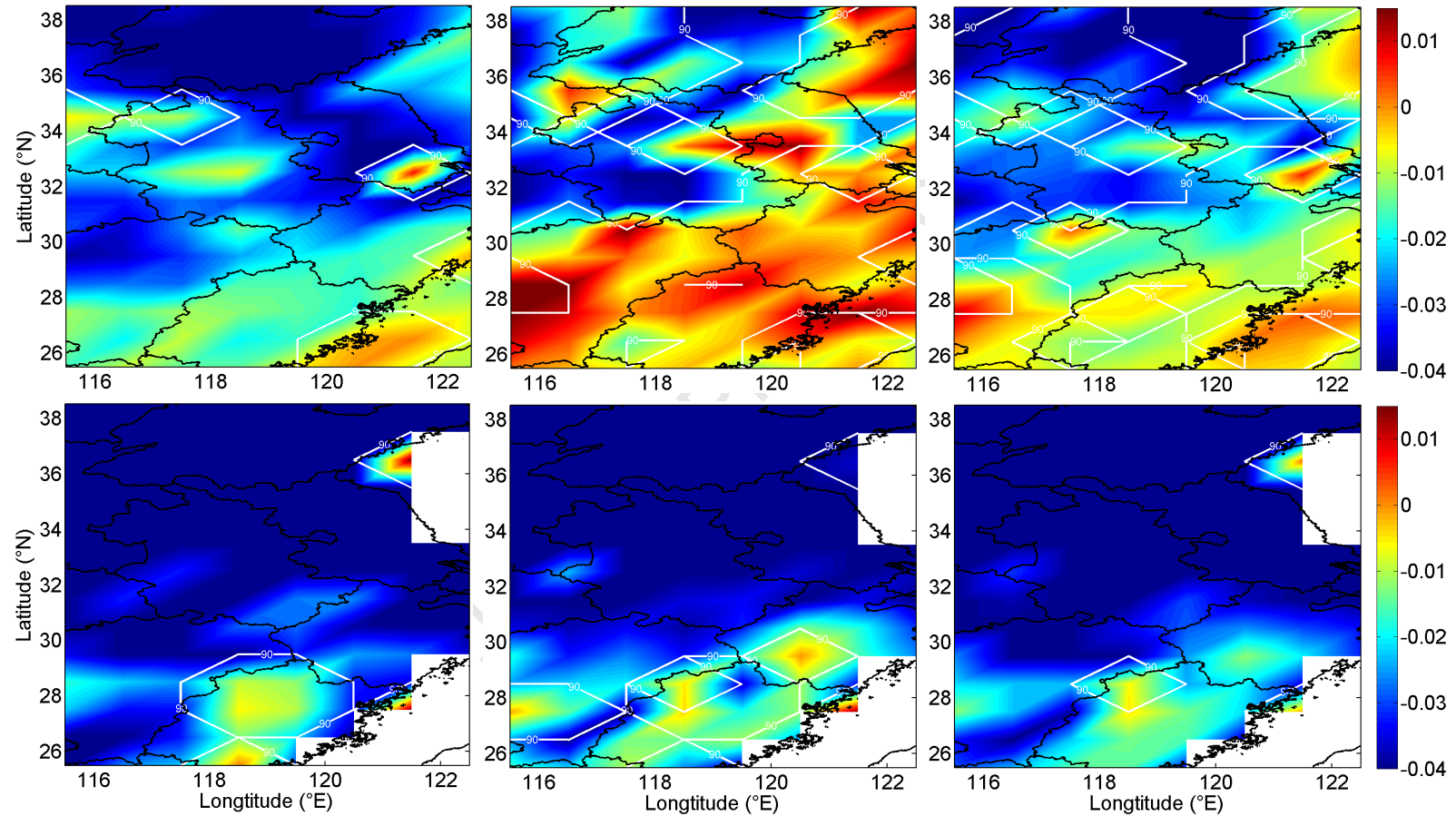


Fig. 9. Spatial distributions of trends observed in AOD₅₅₀ (top panels) and AE₄₄₀₋₆₆₀ (bottom panels) derived from the MODIS Aqua satellite during 2014-2017. The first, middle, and last panels in two rows corresponds to the autumn, winter and autumn+winter, respectively observed over East China. There are no units for the observed trends in AOD and AE. The white lines in all panels represent trends significant at 90% confidence levels and rest is least significant. The positive and negative values in color scale correspond to the respective trends.

RESEARCH HIGHLIGHTS

1. Quantification of PM concentrations and air quality during autumn and winter seasons.
2. The concentrations of PM_{2.5} found maximum in winter relative during autumn.
3. AOD showed high in autumn clearly demonstrates negative correlation with PM.
4. The CWT analysis revealed a significant contribution to PM during winter.
5. The meteorological variables showed negative relationship with PM in all seasons.