



# Article Recent Abnormal Hydrologic Behavior of Tibetan Lakes Observed by Multi-Mission Altimeters

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Received: 28 July 2020; Accepted: 11 September 2020; Published: 14 September 2020



Abstract: Inland lakes in the Tibetan Plateau (TP) with closed catchments and minimal human disturbance are an important indicator of climate change. However, the examination of changes in the spatiotemporal patterns of Tibetan lakes, especially water level variations, is limited due to inadequate access to measurements. This obstacle has been improved by the development of satellite altimetry observations. The more recent studies revealed that the trend of central TP to grow decreased or reversed between 2010 and 2016. However, thus far, this trend has not been investigated to determine whether this pattern would last for the following years. This study aims to combine the traditional (launched before 2010, e.g., TOPEX/POSEIDON, ERS-1, ERS-2, Jason-1/-2, and Envisat) and recently advanced (launched after 2010, e.g., SARAL and Sentinel-3) altimetry observations to understand the Tibetan lake changes further in recent years. Therefore, we acquired information on the continuous lake level changes in Tibetan lakes using the lake level sequence integration method based on multisource altimetry satellites. The results revealed that water level changes in 22 examined lakes showed abrupt rises in 2016–2018, but the onsets and magnitudes of the rises varied among the lakes. During the study period, the water levels of the lakes (except Nam Co) revealed a drastic rising tendency with a mean rate of 0.74 m/a, which was remarkably higher than the average rate of water level rise over the period 2010–2015 (approximately 0.28 m/a). Specifically, the water level of the nine lakes in the Northern TP (NTP) displayed a significant rising trend, with an average rate of 0.82 m/a. In the Central TP (CTP), the lake level changes were generally divided into two categories. The water levels for the lakes in the Western CTP rose rapidly, while, in the Eastern CTP, the lake water levels rose slowly, with an average rising rate less than 0.40 m/a. The water levels for the lakes in the Northeastern TP (NETP) and Northwestern TP (NWTP) kept a stable rising tendency. According to the results of the climate analysis, the spatial differences of the lake level rise rates were primarily caused by the spatial and temporal changes of precipitation over the TP.

Keywords: Tibetan Plateau; lake; water level; satellite altimetry; climate change; precipitation

#### 1. Introduction

Lakes are an important part of the global hydrological and biogeochemical cycles, as well as an essential resource for human societies [1–3]. These inland waterbodies are quite sensitive to climate changes and human activities. The last several decades have witnessed widespread and tremendous changes in lakes in the context of global changes at the regional and global scales [4–8]. Lake changes have profound effects on the regional water balance, ecosystems, biogeochemical cycles, exchange of energy, and tracing gases with the atmosphere and human water consumption. Therefore, the investigation of lake variability is critical for a wide range of socioeconomic, political, and scientific interests. Several recent efforts, such as Pekel et al. [9], Donchyts et al. [10], Yao et al. [11], and Busker et al. [12], elucidated the lake dynamics at the global scale using multitemporal satellite optical and altimetry observations. The prominent characteristics revealed in these studies suggest that many large lakes, such as the Caspian Sea, Aral Sea, and Urmia Lake in the endorheic basins of the Eurasian hinterlands, shrank remarkably in recent decades [13]. In contrast, widespread lake growths have been observed in Earth's highest plateau, the Tibetan Plateau (TP).

The TP is characterized by vast distributions of glaciers and lakes. There are approximately 1200 lakes (lakes with sizes  $\geq 1 \text{ km}^2$ ), with a total area of about 46,500 km<sup>2</sup>, according to recent investigations [14,15]. The lack of historical in-situ data in the remote and sparsely populated TP has greatly limited the investigations of the hydrological and climate responses of the lakes to the changing climate. However, this obstacle has been reduced by the rapidly emerging satellite observations from various sensors. Many prior studies employed optical satellite images to examine lake changes across space and time in the TP [14–24]. However, the acquisition of high-quality optical images is usually impeded by unfavorable weather conditions, which limit the monitoring of interannual and seasonal changes in the lake area. Moreover, derived lake area variations only indirectly indicate the lake water volume changes. Therefore, increasing satellite altimetry measurements have been synergized for the investigation of the lake dynamics across the TP.

Phan et al. [25], Song et al. [18], Wang et al. [26], and Zhang et al. [27] examined the changes in the lake water levels in the TP using the Ice, Cloud and land Elevation Satellite (ICESat) laser altimetry data and found that most Tibetan lakes experienced rapid water level rises during 2003–2009. However, the long revisit cycle and sparse sampling of the ICESat mission is insufficient for tracing continuous and real-time changes in the water levels, thereby making it difficult to explore the trends and spatial differences of lake changes in the TP. Favorably, the height measurements of satellite radar altimeters that have been launched continuously since the early 1990s can extend the lake-level observations to a longer temporal coverage. The advantages are applications that are not affected by the weather and a reasonable high frequency of revisits from monthly to ten-day time scales so that we can better solve the spatial and temporal patterns of the water level changes and driving mechanisms in the TP. The combination of multi-radar altimetry missions, including TOPEX/POSEIDON, ERS-1/-2, Envisat, Cryosat-2, and Jason-1/-2/-3, at the height accuracy ranging from decimeters to centimeters have been used to investigate the decadal water level changes of the lakes in the TP [28–30]. However, these radar altimeters are only suitable for large water bodies, and many lakes are only accessible by one or two satellites because of their large-size footprints and along-track/cross-track spacing.

The recent advances in satellite altimeters, such as CryoSat-2 (2010–), SARAL/AltiKa (2013–), Jason-3 (2016–), and Sentinel-3 (2016–), provide alternative options of measuring relatively smaller water bodies (relative to the lakes observed by traditional radar altimetry satellites). For example, compared with the previous altimetry satellites, the improvement of SARAL/AltiKa was that it was the first altimeter to measure in the Ka band, with higher spatial resolutions. In addition, the enhanced bandwidth of AltiKa (500 MHz) resulted in a higher vertical resolution and higher pulse repetition frequency [31]. Therefore, the more recent studies benefiting from the longer and finer observations revealed that the growth tendency of the lakes in the Central TP were decelerated or reversed during the period 2010–2016 [32–35]. However, whether the deceleration or hiatus has lasted in the following years remains unclear. Although there have been studies that update the lake level changes in the

TP [34,36,37], the focus has mostly been on extending the lake level observation times. However, there has been no systematic analysis of abnormal lake level changes in recent years. Thus, this study aims to combine the traditional and recently advanced radar altimetry measurements to update our understanding of the changing characteristics of Tibetan lakes in the recent past several years. Furthermore, we will explore the potential climate-driving mechanisms of the recent varying tendency

#### 2. Materials and Methods

of lake changes in the TP.

## 2.1. Study Region

The TP, located in Central Eurasia, has a geographic area about 3 million km<sup>2</sup> [18] and mean altitude of approximately 4000 m. The climate in the TP is marked by low temperatures and strong solar radiation [38]. Precipitation in the TP is characterized by strong seasonality. Most of the precipitation (60–90% of the annual total precipitation) occurs between June and September, while the rainfall from November to February is less than 10%. [39]. The TP and its surroundings are sometimes known as "Earth's Third Pole" and the "Water Tower of Asia" because of the abundant water resources. It is the birthplace of many large rivers (Yangtze River, Yellow River, Mekong River, etc.) in China and throughout Asia. In addition, the region is densely distributed with many lakes, which are currently undergoing significant changes due to the impacts of climate changes [40]. In this study, a total of 22 lakes (with sizes from 111.93 km<sup>2</sup> to 4493.52 km<sup>2</sup>) in the TP were selected for analysis based on the altimetry data coverage. The characteristics of the studied 22 TP lakes covered by satellite altimetry are presented in Figure 1.



**Figure 1.** Locations and satellite altimetry overpass of the selected lakes over the Tibetan Plateau (TP) and nearby meteorological stations.

#### 2.2. Study Data

# 2.2.1. LEGOS Hydroweb and DAHITI Altimetry Data Archives

Observation data provided by altimetry satellites, including TOPEX/POSEIDON(T/P), ERS-1, ERS-2, Jason-1/-2/-3, Envisat, and Cryosat-2, have been gradually increased since the early 1990s (Table 1). The combination of different altimetry satellites can increase the spatial coverage of altimetry measurements and extend the temporal span to nearly thirty years. At present, the long-term lake-level changes reconstructed by many institutions are based on the altimetry satellite data of different service years. As a data source of our analyzed water levels, we obtained lake levels from the Laboratoire d Etudes en Géophysique et Océanographie Spatiales (LEGOS) Hydroweb altimetry data archive and the Hydrological Time Series of Inland Waters (DAHITI) altimetry data archive. The LEGOS Hydroweb altimetry data archive was derived from merged Topex/Poseidon, Jason, Envisat, and Sentinel data [41,42]. All background rectifications have been used to produce this data archive. Since different altimetry satellites are used for the same lake, the water level calculation is mainly divided into three steps. First, each satellite data was processed separately to obtain the altimetry measurement data of different satellites, and then, Topex/Poseidon data was used as the reference standard to minimize potential radar instrument deviations and geoid differences between different satellites. Finally, lake levels from the different satellites were merged. The DAHITI data archive was launched by the Deutsches Geodätisches Forschungsinstitut (DGFI, now DGFI-TUM) in 2013. The processing procedure of the DAHITI database in detail can be referred to in Schwatke et al. [43], which was based on an extended outlier detection and a Kalman filtering of multisource-merged altimetry measurements. In this study, because the algorithms adopted by the two data archives were different, we only used one of them for a lake. When a lake was covered by the two data archives at the same time, the Hydroweb data archive was selected as the priority.

Mission	Altimeter	Revisit Cycle (Days)	Along-Track Spacing	Footprint	Operational
Topex/Poseidon	Poseidon-1	10	~620 m	2.2 km	1992-2005
Jason-1	Poseidon-2	10	~294 m	2.2 km	2001-2013
Jason-2	Poseidon-3	10	~294 m	2.2 km	2008-active
Jason-3	Poseidon-3B	10	~294 m	2.2 km	2016-active
ERS-2	RA	35	~374 m	1.7 km	1995-2011
Envisat	RA2	35	~374 m	1.7 km	2002-2010
Cryosat-2	SIRAL	369 (30)	~173 m	1.6 km	2010-active
SARAL	Altika	35	~173 m	8 km	2013-2016
Sentinel-3	SRAL	27	~300 m	0.3 km	2016–active

Tał	ole	1.	Lis	t o	f al	l al	timeter	missio	ns used	l in	this	study	, toget	her	with	their	main	chara	cteristics.

Both data archives cover a certain number of lakes in the TP (Table 2). The long-term water level change information of 16 lakes was obtained based on the Hydroweb and DAHITI data archives. However, the number of lakes cannot fully meet the needs of our lake-level monitoring (refer back to Figure 1). Therefore, in this study, SARAL and Sentinel-3 altimetry data were used to supplement the missing observations from the LEGOS Hydroweb and DAHITI data archives. We interpreted and combined SARAL and Sentinel-3 altimetry data to obtain a total of 12 lake-level time series during the period 2013–2018. Six of them, covering the same lakes as the Hydroweb and DAHITI data archives, were used to verify the reliability of SARAL and Sentinel-3 altimetry data. Others were used as the data supplement.

Lake Name	Latitude	Longitude	Lake Area (km <sup>2</sup> )	Data Source(s)	Satellite(s)	
Jieze Caka	33.95	80.89	116.08	SARAL/Sentinel-3	SARAL; Sentinel-3	
Lumajangdong Co	34.04	81.65	389.95	Hydroweb	Envisat; SARAL; Sentinel-3	
Heishibei Lake	35.58	82.77	111.93	SARAL/Sentinel-3	SARAL; Sentinel-3	
Ngangla Ringco	31.57	83.05	493.09	Hydroweb	Envisat; SARAL; ERS2; Sentinel-3	
Zhari Namco	30.94	85.57	991.73	Hydroweb	SARAL; Topex/Poseidon; Jason2; Jason 3; Sentinel-3	
Tangra Yumco	31.02	86.55	841.95	Hydroweb	Envisat; SARAL; ERS2; Sentinel-3	
Ngangze Co	31.04	87.13	491.67	Hydroweb	Envisat; SARAL; Topex/Poseidon; Jason2; Jason3; ERS2	
Dagze Co	31.90	87.52	305.02	Hydroweb	Envisat; SARAL; Topex/Poseidon; Jason2; ERS2; Sentinel-3	
Aqqikkol	37.09	88.39	537.63	DAHITI	Cryosat-2; Sentinel-3	
Gyariog Co	31.09	88.42	474.56	SARAL/Sentinel-3	SARAL; Sentinel-3	
Dogai Coring	34.56	88.99	485.26	Hydroweb	Envisat; SARAL; Jason2; Jason3; ERS2	
Seling Co	31.81	89.03	2393.53	LEGOS Hydroweb	Envisat; SARAL; Topex/Poseidon; Jason; ERS2; Sentinel-3	
Dogaicoring QiangCo	35.32	89.25	394.53	Hydroweb	Envisat; SARAL; ERS2; Sentinel-3	
Ayakkum	37.54	89.43	979.85	Hydroweb	Envisat; SARAL; ERS2; Sentinel-3	
Lexiewudan	35.75	90.21	270.95	Hydroweb	Envisat; SARAL; ERS2; Sentinel-3	
Migriggyangzham Co	33.39	90.27	1066.43	Hydroweb	SARAL; ERS2; Topex/Poseidon; Jason2; Jason3	
Xijin Ulan Lake	35.21	90.32	539.76	SARAL/Sentinel-3	SARAL; Sentinel-3	
Úlan Ul Lake	34.84	90.42	646.31	SARAL/Sentinel-3	SARAL; Sentinel-3	
Nam Co	30.71	90.57	2010.60	Hydroweb	Envisat; SARAL; ERS2; Sentinel-3	
Mingjing Lake	35.06	90.58	122.61	SARAL/Sentinel-3	SARAL; Sentinel-3	
Kusai Lake	35.73	92.89	327.79	DAHITI	SARAL; Jason2; Jason3; Sentinel-3	
Qinghai Lake	36.92	100.19	4493.52	Hydroweb	Envisat; SARAL; Topex/Poseidon; Jason; ERS2; Sentinel-3	

Table 2. Our studied lakes in the Tibetan Plateau (TP) measured by satellite altimetry.

## 2.2.2. Supplementary Altimetry Data

## SARAL/AltiKa Data

SARAL is a common mission of the Indian Space Research Organization (ISRO) and the France space agency [31]. The mission, launched on 25 February 2013, is a continuation of the ERS series and Envisat missions. The orbit and revisit cycle of SARAL satellite were the same as Envisat, but after July 2016, the SARAL satellite was then transformed to drifting ground track mode. We downloaded the SARAL data (spanning from March 2013 to June 2016) from ftp://avisoftp.cnes.fr.

## Sentinel-3/SRAL Data

The Surface Topography Mission uses the Synthetic Aperture Radar Altimeter (SRAL) instrument, which transmits pulses occasionally at Ku band for altimeter range measurements and is supplemented by a C-band frequency to correct range delay errors. The SRAL instrument acquires data in two measurement modes: the low-resolution mode (LRM), which is useful over homogeneous open-ocean surfaces, and the synthetic aperture radar (SAR) mode, which achieves high along-track resolution over relatively flat surfaces. Sentinel-3 satellite has multiple new functions and modes with the dual-frequency SAR radar altimeter (SRAL) [44]. Among them, the tracking mode is the most critical characteristic. We downloaded the needed data from March 2016 to December 2018 through https://sentinel.esa.int/web/sentinel/ sentinel-data-access.

## 2.2.3. Climate Data

The annual and monthly precipitation data from 10 stations in the TP were collected from the China Meteorological Network (http://data.cma.cn) to investigate the climate impacts of the lake water level changes in 1990–2018. The locations of the ten stations are shown in Figure 1. In addition, we analyzed the changes of water vapor transport modes over the TP and its surroundings based on ERA-interim reanalysis data with a spatial resolution of  $0.75^{\circ} \times 0.75^{\circ}$  [45].

## 2.3. Methods

## 2.3.1. Altimetry Water Level Processing

We retrieved important variables, such as time, latitude, longitude, water level elevation, and geoid, from the collected raw data files. Since the parameter information contained in each altimeter satellite was not completely consistent, the water level calculations needed to refer to the data manual for each altimeter satellite.

## 2.3.2. Removal of Measurement Outliers

For all lakes, the original tracks of the satellite are observed as they pass through the lake, and then, the track data falling within the lake will be selected. As the observations contain outliers, it is necessary to eliminate them for subsequent work. Outlier elimination for the selected orbit data was divided into two steps [35,46]:

Step 1: We removed the outliers that were more than 3 standard deviations from the median for all the measurements.

Step 2: We removed the data other than twice the standard deviation for single-day measurements.

2.3.3. Unifying the Reference System of Different Altimetry Sources

In order to reduce the differences caused by different altimeter instruments, all lake levels from multisource altimetry satellites were first converted to the same geographic reference system. The system biases calculated by the data from different altimetry satellites at the same period were subtracted from the elevation measurements for each lake [47,48].

The specific procedures are shown in Figure 2.



Figure 2. Flowchart of the combining multisource altimetry data.

## 3. Results

The best way of altimetry data validation is to compare the altimeter-derived lake level with the in-situ data. At present, only a few lakes, such as Qinghai Lake and Nam Co, have long-term continuous water level observations in the TP, which are extremely limited compared with more than 1200 lakes (>1 km<sup>2</sup>) across the TP [49]. The altimetry data archives, such as Hydroweb and DAHITI, play a critical role in verifying the lake level changes obtained by other satellite altimetry data. In order to verify the reliability of the Hydroweb and DAHITI altimetry data archives, the annual means of the water level for Qinghai Lake were validated against the in-situ water levels measured by the hydrological station. As shown in Figure 3, the altimetry levels demonstrate a robust linear correlation with in-situ measurements, with an R<sup>2</sup> value of 0.90. This agreement indicates that the lake level time series obtained from Hydroweb and DAHITI have a high reliability and, therefore, were used as a reference for validating our level time series processed from the SARAL and Sentinel-3 altimeters.



**Figure 3.** Correlation between in-situ water level observations and satellite altimetry observations (from Hydroweb and DAHITI) for Qinghai Lake.

Figure 4 presents the time series of water levels for eight large TP lakes from the Hydroweb and DAHITI data archives from the 1990s to 2018. All of the eight lakes showed rapid water level increases since around 2000. By comparison, the abrupt changing points varied with different lakes. For example, the lake levels for Seling Co, Nam Co, Ngangze Co, and Tangra Yumco began to rise rapidly before 2000, while Migriggyangzham Co and Qinghai Lake began to grow rapidly after 2003 and 2005, respectively. The most significant expansion occurred in Seling Co in Central Tibet. From May 1998 to December 2010, the lake level for Seling Co rose from 4533.66 m to 4543.77 m, with an increasing rate exceeding 0.80 m/a. However, for these lakes, the increasing trends became weaker or even reversed after 2010. The findings were consistent with other previous studies [32–35,50]. However, the lake levels for some lakes, such as Qinghai Lake, Zhari Namco, Tangra Yumco, Ngangze Co, and Migriggyangzham Co, began to increase significantly again after 2016. Among them, the abrupt rise of Zhari Namco was the most remarkable, with a water level increase during 2016–2018 of more than the total during the previous 16 years (2000–2016). By contrast, the lake level for Nam Co showed interannual fluctuations without an evident trend during the recent several years. The level time series of other lakes are provided in the Supplementary Materials (Figure S1).



**Figure 4.** Long-term lake level variations of eight large lakes in the TP during 1995–2018 based on the Hydroweb and DAHITI altimetry data archives.

## 3.2. Lake Level Variations Based on SARAL and Sentinel-3 Altimetry Data

In order to examine whether abrupt water level increases during 2016–2018 occurred similarly in other Tibetan lakes, we investigated the water level changes of a total of 12 lakes from 2013 to 2018 by combining SARAL and Sentinel-3 altimetry data. Six of them were aimed to intercompare with the Hydroweb and DAHITI altimetry data. Figure 5 presents the comparisons of the water-level time series from the Hydroweb/DAHITI data archives and SARAL/Sentinel-3 data between 2013 and 2018 for Qinghai Lake, Nam Co, Ayakkum, Migriggyangzham Co, Zhari Namco, and Tangra Yumco. The water level time series for all the lakes indicated that general trends derived from two data sources are comparable, and abrupt changes and intra-annual variations of the lake levels are generally consistent. For example, both data sources clearly depicted a sudden increase of about 1.0 m during the wet season (July–September) of 2017 for Zhari Namco. The characteristics of the water levels rises in the wet seasons and declines in the dry seasons are also in good agreement for all the lakes, suggesting that SARAL and Sentinel-3 data have reliable results.



**Figure 5.** Comparison of lake level changes based on the Hydroweb/DAHITI and SARAL/Sentinel-3 altimetry data (relative levels changes refer to lake levels relative to the beginning of the time series).

Based on the above validation, the SARAL and Sentinel-3 data were used to monitor the lakes larger than 100 km<sup>2</sup> in the TP, and the water level time series are shown in Figure 6. The six lakes display an evident trend of water level increases at varying rates during 2016–2018. For example, the alpine lakes between the Tanggula and Kunlun Mountains, including the Xijin Ulan Lake, Mingjing Lake, and Ulan Ul Lake, showed a significant rise (exceeding 1 m/a). By contrast, the lake level for Gyariog Co showed an interannual fluctuation.

## 3.3. Spatial Distribution of Lake Level Changing Rates During 2016–2018

Our results of the lake level changing rates are shown in Figure 7 and Table 3. Obviously, most lakes experienced a significant increase during the period 2016–2018. To be specific, except the lake level of Nam Co (remained relatively stable, with little fluctuations), the water level time series of the other lakes revealed a rising tendency with a mean rate of 0.74 m/a. The three lakes in the Northwestern TP: Jieze Caka, Lumajangdong Co, and Heishibei Lake showed significant differences in water level rising rates, with the level for Heishibei Lake increasing the fastest at 1.39 m/a. The northern part of the TP is located in the Hoh Xil Region, which is known for a complex and harsh geographical environment. A large number of endorheic lakes are distributed in this region. As listed in Table 3, most of the lakes have expanded, with an average level increase of larger than 0.82 m/a. The level for Xijin Ulan Lake increased at the greatest rate of 1.52 m/a. Ayakkum, the largest lake in this region, also exhibited a significant increase of 0.83 m/a. The average increasing rate was 0.66 m/a in the central part of the TP. The lake levels for Dagze Co and Zhari Namco experienced a rapid rise, exceeding 1 m/a. In Qinghai Province, the level for Qinghai Lake maintained a steady increasing trend of



0.47 m/a. Overall, most lakes in the Northern Tibetan Plateau experienced relatively faster rates of water level rises compared to those in the Northeastern TP (NETP) and Eastern CTP.

Figure 6. Lake level changes during the period 2013–2018 based on SARAL and Sentinel-3 altimetry data.



**Figure 7.** Spatial distribution of lake level changing rates in the Tibetan Plateau during 2016–2018 (red rectangles show the boundaries of different TP regions). NWTP: Northwestern TP, NTP: Northern TP, NETP: Northeastern TP, and CTP: Central TP.

Lake Name	Latitude	Longitude	Elevation (m)	Area (km²)	Lake Type	Change Rate (m/a)
Jieze Caka	33.95	80.89	4531	113.5	Y	0.20
Lumajangdongco	34.04	81.65	4817	389.9	Y	0.86
Heishibei Lake	35.58	82.77	5054	111.9	Y	1.39
Ngangla Ringco	31.57	83.05	4715	493.1	Y	0.15
Zhari Namco	30.94	85.57	4613	991.7	Y	1.34
Tangra Yumco	31.02	86.55	4537	841.9	Y	0.70
Ngangze Co	31.04	87.13	4688	462.2	Y	0.87
Dagze Co	31.90	87.52	4471	305.0	Y	1.43
Aqqikkol	37.09	88.39	4251	537.6	Y	0.58
Gyariog Co	31.09	88.42	4647	474.6	Ν	0.40
Dogai Coring	34.56	88.99	5038	485.3	Y	0.47
Seling Co	31.81	89.03	4544	2393.5	Y	0.32
Dogaicoring Qiangco	35.32	89.25	4787	394.5	Y	0.82
Ayakkum	37.54	89.43	3876	979.8	Y	0.83
Lexiewudan	35.75	90.21	4870	271.0	Y	0.90
Migriggyangzham Co	33.39	90.27	4933	1066.4	Y	0.70
Xijin Ulan Lake	35.21	90.32	4772	539.8	Y	1.52
Úlan Ul Lake	34.84	90.42	4855	646.3	Y	0.61
Namco	30.71	90.57	4724	200.6	Y	0.01
Mingjing Lake	35.06	90.58	4797	122.6	Y	1.49
Kusai Lake	35.73	92.89	4475	327.8	Y	0.17
Oinghai Lake	36.92	100 19	3194	4228.3	Y	0.47

**Table 3.** Statistics of lake level changing rates in the Tibetan Plateau from 2016 to 2018 ("Y" represents terminal lakes and "N" open lakes).

#### 4. Discussion

## 4.1. The Driving Factors of Abnormal Lake Level Changes

Since the late 1990s, the expansion of Tibetan lakes has been mainly affected by regional climate changes, such as an increased water vapor content, potential evaporation reduction, and significantly enhanced air convection [51,52]. Most of the previous studies only analyzed the correlation between the lakes and climate changes from a qualitative perspective. However, which driving factors dominate the lake dynamics is still under debate. Favorably, benefiting from research based on quantitative analyses in recent years, it has become a consensus that precipitation variations play a leading role in lake water storage [53,54]. Therefore, with the acceleration of climate warming and humidification, the precipitation could be also gradually increase [52], which supports the observation of the rapid lake expansion in the TP.

As shown in Figure 4 and the Supplementary Materials (Figure S1), the trends of lake level changes since 1990 are generally consistent with previous studies. Our findings revealed that the drastic water level changes in Tibetan lakes since 2016 are a result of continued precipitation increases after a relatively dry period from 2012 to 2016. This suggested that precipitation remains to be a first-order driver of Tibetan lake dynamics. It was found that the drastic increase of lake levels in 2017–2018 was consistent with the increase of precipitation in the recent past several years, so the monthly precipitation data in recent years was selected and compared with the lake level. According to the study, the monthly precipitation and the lake water levels during the period 2013–2018 were selected to study the influence of specific monthly precipitations on the lake water levels. The most prominent feature was the good correspondence between lake level increases and the summer precipitations. In addition, we can find that the decrease of water levels in individual years (e.g., 2015) corresponded to less summer precipitation amount. For example, the drastic lake level rises between 2016 and 2018 in the TP corresponded to high precipitation, as shown in Figure 8. This suggested that

the excessive precipitation during the wet season played a crucial role in the recent lake level increases. Although the lake level changes and precipitation do not correspond well for some lakes in individual years, meteorological stations only represent the meteorological characteristics of a local region and cannot indicate water volume variations in the whole watershed. The station data can only relatively show the trend of regional precipitation changes due to the obvious spatial heterogeneity of rainfall. Moreover, we calculated the total annual precipitation at the Gaize and Tuotuohe Meteorological Stations. We can find that the summer precipitation in 2018 was not as significant as that in 2017 at the Gaize Meteorological Station, but the total annual precipitation was still significantly higher than in ordinary years (2013–2015).



**Figure 8.** Observed lake level changes and precipitation variations for typical lakes between 2013 and 2018. The right Y-axis "lake increase" refers to level changes relative to the previous year (shown as red dots in the plot). The highest water level during the year is used as the annual lake level. Blue bars represent monthly precipitation at the Gaize and Tuotuohe meteorological stations, and the total annual precipitation was calculated. The light blue stripes indicate the corresponding changes in precipitation and lake levels.

The rise of lake levels is not only affected by precipitation but, also, by glacier mass loss and permafrost thawing. However, limited by the in-situ data availability in the TP, the current quantitative studies on Tibetan lakes only focus on individual lakes. For the causes of lake changes such as climate (precipitation and temperature) or cryosphere (glacier and permafrost), qualitative analyses have been the norm [54]. However, this was implied when warming-induced melting did not reverse the slow-down or hiatus of lake expansions before 2016 (when precipitation was less), and the reoccurrence of precipitation increases eventually triggered another round of lake level increases.

## 4.2. Spatial Linkage of Increased Precipitation to Enhanced Water Vapor Circulations

Previous studies have identified the main sources of moisture for precipitation in the TP, including regions from the Indian subcontinent to the Southern Hemisphere, the Bay of Bengal, and the northwestern part of the TP [55,56]. It is proven that these circulation changes in the increase or decrease of different moisture sources provide favorable water vapor conditions for rainfall in the TP [57]. Thus, the changes in moisture sources contributed to different precipitation changes across the entire TP during different periods. As indicated in the preceding results, the lake water levels increased rapidly during the period 2016–2018 relative to the previous stage. The phenomenon was analyzed with the distribution of water vapor transport in the TP. Although the water vapor transport began to change in 2016 (Figure 9C), the changes in 2017/2018 were more significant. Therefore, the water levels of lakes in only some regions rose rapidly in 2016, unlike the general drastic rises in lake levels in 2017/2018.



**Figure 9.** Anomaly distributions of the water vapor and its flux over the TP and its surrounding regions in 2014 (**A**), 2015 (**B**), 2016 (**C**), 2017 (**D**), and 2018 (**E**) and composite analysis of the water vapor relative to the years with rapid increases and decreases of precipitation (**F**). The black arrows represent the transport of water vapor by the horizontal movement of the atmosphere. The green rectangular boxes represent the location of the TP.

The differences can be observed during the four years when the water vapor is transported by abnormal northwesterly and southeast wind and south wind in 2015, 2016, 2017, and 2018, respectively (Figure 9B–E). By comparison, the total water vapor increased, with the most significant anomaly in 2017–2018. In this study, precipitation in the Northern TP (NTP) was more obviously influenced by the East Asian monsoon and the Indian summer monsoon. From Figure 9D, we can also find that there is a phenomenon of internal water vapor circulation and retention, which might cause water

vapor subsidence and significant rainfall over the Western CTP. In general, the change in water vapor transport during the study period supports the changing characteristics of the precipitation and lake level for the TP, as analyzed above.

Figure 7 and Table 3 present the differences in the rising rates of the lake levels at different parts of the TP in this study. The water level rising rates for some typical lakes, such as Xijin Ulan Lake, Mingjing Lake, Lexiewudan, and Ayakkum, exceed 0.80 m/a in the Northern TP. There are some glaciers distributed in the Northern TP, and there is rich rainfall in the rainy season. Therefore, the rapid increase of the water level is affected by both rainfall and glacial meltwater. The rising rate of the lake levels in Central TP is mainly divided into two categories. For example, the rising rates of Zhari Namco and Dagze Co in the Western CTP are 1.34 and 1.43 m/a, respectively. The rising rates of the lakes in the Central and Eastern CTP are slower, generally less than 0.40 m/a. The reason for this difference may be because the phenomenon of internal water vapor circulation and retention mainly occur in the Western CTP. The Northwestern TP (NWTP) was mainly affected by the westerlies, and the distinction in the water vapor transport distribution in different years led to a significant difference in the water level rise rates are caused by various water vapor transport patterns and main sources in different regions.

#### 5. Conclusions

The more recent studies benefiting from the longer and finer observations revealed that the growth tendencies of lakes in the Central TP were decelerated or reversed during the period 2010–2016. However, whether the deceleration or hiatus would last in the following years remains unclear. It has not been systematically investigated thus far. Thus, in this study, 22 large inland lakes were investigated to understand the changing characteristics of Tibetan lakes further during the period 2016–2018 by combining the traditional and recently advanced radar altimetry measurements. Furthermore, we explored the potential climate-driving mechanisms of the recent varying tendencies of lake changes in the TP. For most lakes in the TP, the lake levels showed an abrupt rise during 2016–2018, compared with the earlier stage during 2010–2016, but the onsets and magnitudes of the water level rises varied with the subzones and lakes.

Our results revealed that the lake water levels in the NTP displayed a sharp rise, with an average rate of 0.90 m/a, except Kusai Lake (outburst of its upstream lake). In the Central TP, the lake level changes were divided into two categories during the period 2016–2018, which was different from the dramatic increasing pattern for lakes in the Northern TP. The water level for Qinghai Lake in the Northeastern TP kept a stable growth tendency during the recent years at 0.47 m/a. The water level rising rates for the three lakes in the Northwestern TP (Lumajangdong Co, Jieze Caka, and Heishibei Lake) are extremely different from each other. The spatial differences of the lake level rise rates are primarily caused by the changes in precipitation over the TP, which may be related to large-scale atmospheric circulation. Although this study could be beneficial for our understanding of the driving mechanisms of the rapid water level rises of Tibetan lakes and the hydrological cycle under future climate change conditions, more comprehensive investigations of the precise physical processes are required to clarify the climate-driven mechanisms of Tibetan lake level changes in further studies.

**Supplementary Materials:** The following are available online at http://www.mdpi.com/2072-4292/12/18/2986/s1, Figure S1: Long-term lake level variations of the other eight large lakes in the TP during 1995–2018 based on the Hydroweb and DAHITI altimetry data archives. Figure S2: Comparison of the data distribution between the in-situ water level observations and satellite altimetry observations (from Hydroweb and DAHITI) for Qinghai Lake.

Author Contributions: Conceptualization, C.S. and P.Z.; data curation, P.Z., L.K., W.L., and C.S.; formal analysis, P.Z., C.S., W.L., and J.W.; funding acquisition, C.S. and J.W.; investigation, P.Z., J.W., and L.K.; methodology, P.Z., C.S., and L.K.; project administration, C.S.; resources, P.Z., C.S., and L.K.; software, P.Z., L.K., and J.W.; supervision, C.S. and J.W.; validation, P.Z., C.S., J.W., K.L., and T.C.; visualization, P.Z. and C.S.; writing—original draft, P.Z. and C.S.; and writing—review and editing C.S., J.W., L.K., P.Z., K.L., and T.C. All authors have read and agreed to the published version of the manuscript.

**Funding:** This work was partly funded by the Strategic Priority Research Program of the Chinese Academy of Sciences (Grant No. XDA23100102); the Second Tibetan Plateau Scientific Expedition and Research (2019QZKK0202); the National Key Research and Development Program of China (Grant No. 2019YFA0607101, 2018YFD0900804, and 2018YFD1100101); the Thousand Young Talents Program in China (Grant No. Y7QR011001); and the National Natural Science Foundation of China (No. 41971403 and 41801321).

**Acknowledgments:** We are also grateful to the Laboratoire d'Etudes en Géophysique et Océanographie Spatiales (LEGOS), the Hydrological Time Series of Inland Waters (DAHITI), the NASA EOSDIS Physical Oceanography Distributed Active Archive Center (PO.DAAC), Centre National d'Etudes Spatiales (CNES), the China Meteorological Data Sharing Service System, and the European Centre for Medium-Range Weather Forecasts for providing satellite altimetry data and meteorological data for this study.

Conflicts of Interest: The authors declare no conflict of interest.

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