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Potential impacts of future reduced aerosols on internal dynamics characteristics of precipitation based on model simulations over southern China

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**Highlights:**

- Precipitation records over southern China exhibit relatively weak long-term correlation characteristics.
- The CESM1 model has the better performance in simulating the internal dynamics characteristics of precipitation series in southern China.
- The future reduced aerosols emissions will contribute to strengthening the long-term correlation of precipitation records in Huai river basin in summer during the late-21st century (2071-2100).

1 **Potential impacts of future reduced aerosols on internal dynamics characteristics**  
2 **of precipitation based on model simulations over southern China**

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22

23 **Abstract:** In this study, the scaling behaviors of precipitation records over southern  
24 China are investigated by using the detrended fluctuation analysis (DFA) method. It is  
25 found that the precipitation records over southern China exhibit relatively weak  
26 long-term correlation characteristics. The scaling exponents in coastal areas are close  
27 to 0.6 showing long-term correlation while in inland areas, uncorrelation can be found  
28 with the scaling exponents close to 0.5. Based on the long-term correlation  
29 characteristics of the observed precipitation records, the performance of Community  
30 Earth System Model (CESM1) in simulating precipitation over southern China is  
31 evaluated and the results show that the CESM1 has the ability to simulate the internal  
32 dynamics characteristics of precipitation series in southern China. As indicated by the  
33 DFA results of simulated precipitation data from CESM1, the long-term correlation of  
34 precipitation records during the late-21st century (2071-2100) will increase in Huai  
35 river basin under the RCP8.5 simulation scenario in summer and decrease in most  
36 regions of southern China under both the RCP8.5 and RCP8.5\_FixA scenarios in  
37 comparison with the present condition (1987-2016). And the differences of scaling  
38 exponents of precipitation series between the RCP8.5 and RCP8.5\_FixA simulation  
39 scenarios further reveal that the future reduced aerosols emissions will contribute to  
40 strengthening the long-term correlation of precipitation records in Huai river basin in  
41 summer during the late-21st century (2071-2100) and the scaling exponents will  
42 increase by more than 0.1 in comparison with the present condition (1987-2016).

43 **Keywords:** precipitation; aerosols; detrended fluctuation analysis (DFA)

## 45 **1 Introduction**

46 As we all know, the weather changes with short-term correlation. This feature can be  
47 maintained for a week (the duration of a weather system) (Havlin et al., 1999).  
48 However, it is noteworthy that, in recent years, results of the nonlinear research have  
49 shown that changes in the climate system have self-memory. That is to say, the past  
50 climate evolution has a long-term effect on the future evolution of the climate system  
51 (Monetti et al., 2003). This effect is the so-called “long-term correlation”, “long-term  
52 memory” or “long-term persistence” in nonlinear science. Long-term correlation  
53 essentially describes the self-similarity of the evolution of climate system at different  
54 timescales. Therefore, climate variability characteristics on a large timescale can be  
55 inferred from the features of the climate system itself on a small timescale (Shukla,  
56 1998). In fact, in early 1951, Hurst, a British hydrologist (Hurst, 1951), discovered  
57 that long-term correlation characteristics existed in nature when he studied the  
58 variation of water flow in the Nile. He found that if there is a large amount of water  
59 flowing in the first year, the river’s water flow of the following year would be larger,  
60 which showed that the system had a clear long-term correlation characteristic. With  
61 the unremitting research of scientists, long-term correlation characteristics have been  
62 found in many climate variables, such as precipitation series (Bunde et al., 2013; Zhao  
63 and He, 2015; Jiang et al., 2017), temperature records (Monetti et al., 2003; Király et  
64 al., 2005; Yuan et al., 2015), wind speed records (Koçak, 2009), ozone (Varotsos et al.,  
65 2006), and outgoing longwave radiation (Shen et al., 2017). People cannot accurately  
66 quantify the intensity of long-term correlation by the simple calculation of

67 autocorrelation function and power spectral density due to the non-stationarity of time  
68 series (Talkner et al., 2000). In order to solve this problem, the detrended fluctuation  
69 analysis (DFA) technique was introduced and generalized by Peng et al. (1994) and  
70 Bunde et al. (2000). DFA can effectively filter out the strong trend components caused  
71 by external forcing in time series and excavate the fluctuation components induced by  
72 the long-term correlation. Due to its advantages and applicability (Bashan et al., 2008),  
73 in recent years, this method has been widely used to study the evolution of the climate  
74 system (Kantelhardt et al., 2006; Fraedrich et al., 2009; Zhu et al., 2010; Bunde et al.,  
75 2013; Jiang et al., 2015, 2017; Yuan et al., 2015; He et al., 2015) and to evaluate the  
76 performance of climate models (Govindan et al., 2002; Blender and Fraedrich, 2003;  
77 Zhao and He, 2015). Zhao and He (2015) evaluated the precipitation simulation skills  
78 of the BCC\_CSM1.1 (m) by comparing the differences of the scaling exponents  
79 between the climate model-simulated and observational data. Jiang et al. (2017)  
80 investigated the scaling behaviors of precipitation over China and found that the  
81 precipitation time series in some stations exhibit long-term correlation and the scaling  
82 exponents are less than 0.62, showing weak correlation characteristics.

83 Recently, urban air pollution in China has become increasingly severe and the haze  
84 weather frequently occurs in major cities such as Beijing due to high pollutant  
85 emissions (Yin et al., 2015; Cai et al., 2017). Aerosols, as one of the significant  
86 components of pollutants, would be reduced in the future for their harmful effects on  
87 human health as well as the environment (Wang et al., 2017). Precipitation, as one of  
88 the most important climate variables, exerts a huge impact on agriculture, water

89 resource management and human activities (Walther et al., 2002). At present, much  
90 effort has been made to detect and analyze the trend in the occurrence of extreme  
91 precipitation events under future reduced aerosols emissions. Wang et al. (2016)  
92 analyzed the effect of reduced aerosols emissions on projected extreme precipitation  
93 events and found that some extreme precipitation events would become more and  
94 more frequent in the future. However, the potential impacts of future reduced aerosols  
95 on internal dynamics characteristics of precipitation series have not been widely  
96 discussed before. Southern China, located in regions south of the Huai River and  
97 Qinling Mountains and east of the Qinghai-Tibet Plateau, is significantly affected by  
98 weather systems (e.g., subtropical high, southwest vortex) and climate factors (e.g., El  
99 Niño Southern Oscillation, Pacific Decadal Oscillation, Indian Ocean Dipole Mode  
100 and Atlantic Multi-decadal Oscillation) (Yang et al., 2004; Si and Ding, 2016; Zhu et  
101 al., 2016). Due to its complex climate characteristics, the prediction skill for  
102 short-term and long-term climate change is limited in southern China. In this paper,  
103 the DFA technique is used to analyze potential impacts of future reduced aerosols on  
104 internal dynamics characteristics of precipitation in southern China. Such a study may  
105 contribute to understanding the inherent precipitation evolution and further improve  
106 forecast skills in southern China.

107 This study is organized as follows. In Section 2, the source and details of data used in  
108 this paper are introduced and the calculation steps of the DFA technique are described.  
109 In Section 3, the scaling behaviors of precipitation in southern China are detected and  
110 the performance of Community Earth System Model (CESM1) is evaluated. In

111 Section 4, the potential impacts of reduced aerosols on internal dynamics  
112 characteristics of precipitation in southern China are investigated. The conclusion and  
113 possible mechanisms are given in Section 5.

114

## 115 **2 Materials and Method**

### 116 **2.1 Data**

117 The grid precipitation data with 0.5x0.5 horizontal resolution used to evaluate the  
118 performance of CESM1 climate model is selected from the National Meteorological  
119 Information Center of China (<http://data.cma.cn/site/index.html>). The data covers  
120 11315 days from 1961 to 2007 (Xie et al., 2007). In this paper, we select the data from  
121 1976 to 2005 as observations in order to keep the consistency of the output data  
122 period from a climate model. The global climate model (GCM) used in this paper is  
123 Community Earth System Model (CESM1), which is a fully coupled  
124 ocean-atmosphere-land-sea ice model and can simulate the Earth's climate states for  
125 the past, present and future (Hurrell et al., 2013). The simulation scenarios for the 21st  
126 century is documented in Meehl et al., (2013). The horizontal resolution is 0.9°  
127 latitude x 1.25° longitude for the atmosphere and land and 1° latitude x 1°  
128 longitude for the ocean. The CESM1 includes three-mode modal aerosols scheme  
129 (MAM3 scheme) that can simulate aerosols concentrations and internal mixtures for  
130 dust, organic carbons, sulfate, sea salt and black carbon (Liu et al., 2012). The daily  
131 precipitation data from CESM1 history simulation during 1976-2005 are used to  
132 evaluate the performance of CESM1. The daily precipitation data from the RCP8.5



133 Large Ensemble simulation (Kay et al., 2014) and RCP8.5 with fixed aerosols  
134 (RCP8.5\_FixA) simulation (Xu et al., 2015; Wang et al., 2016) are used to study the  
135 potential impacts of reduced aerosols on internal dynamics characteristics of  
136 precipitation. The RCP8.5 simulation scenario assumes that the energy demand and  
137 greenhouse gas (GHG) emissions are comparatively high in the future due to the  
138 imbalance between the rapidly increasing population and relatively slow  
139 technological innovation and improvements of energy structure. It includes 30  
140 ensemble members forced by the same trajectory of aerosols emissions and the  
141 equivalent GHG but starts from a different atmospheric initial condition (1920-2100).  
142 The RCP8.5\_FixA (15 members) simulation scenario uses the same forcing as the  
143 RCP8.5 simulation scenario except that all aerosols emissions and tropospheric  
144 oxidants are fixed at 2005 levels. The variation of internal dynamics characteristics of  
145 precipitation is examined during the late-21st century (2071-2100) in comparison with  
146 the present condition (1987-2016) under the climate scenario RCP8.5 and  
147 RCP8.5\_FixA. Therefore, the results are compared to assess the effect of reduced  
148 aerosols emissions on internal dynamics characteristics of precipitation under the  
149 different RCP8.5 simulation scenarios.

## 150 **2.2 The DFA method**

151 Recently, more and more scientists are beginning to focus on the non-stationary  
152 characteristics of climate series. Many different analysis techniques, such as the  
153 rescaled range (R/S) analysis technique (Hurst et al., 1951), the fluctuation analysis  
154 technique (Peng et al., 1992), the detrending moving average analysis technique

155 (Alessio et al., 2002) and the detrended fluctuation analysis technique (Peng et al.,  
 156 1994), have been proposed. DFA was proposed by Peng et al. (1994) to detect the  
 157 long-term correlation in climate time series. The more detailed introduction is  
 158 documented in Bunde et al. (2000) and Kantelhardt et al. (2001). Here, a brief  
 159 introduction about the generalized DFA technique is summarized as follows.

160 **The precipitation time series  $\{x(i), i = 1, 2, \dots, N\}$  with annual cycle removed is**  
 161 **considered and its cumulative sum**, which is better known as the profile, is defined as

$$162 \quad y(j) = \sum_{i=1}^{i=j} \Delta x(i) \quad (j = 1, 2, \dots, N).$$

163 Then, the profile  $y(j)$  is divided by  $m$  non-overlapping segments of size  $s$ . In  
 164 order to make full use of the remaining data, we can obtain  $2m$  segments by  
 165 redividing the profile from its back to front. In each segment, the local trend  
 166  $p(v, j)$  is calculated by least squares fit (n-order DFA corresponding n-order  
 167 polynomial) and the residuals are obtained as

$$168 \quad Q(v, j) = y(v, j) - p(v, j) \quad (v = 1, 2, \dots, 2m; j = 1, 2, \dots, s).$$

169 The variance for each segment is defined as

$$170 \quad H(v, s) = \frac{1}{s} \sum_{j=1}^s Q^2(v, j) \quad (v = 1, 2, \dots, 2m). \text{ Finally, the fluctuation function can be}$$

171 obtained by the arithmetic mean of the fluctuation function in all the segments

$$172 \quad F(s) = \left[ \frac{1}{2m} \sum_{v=1}^{2m} H(v, s) \right]^{1/2}.$$

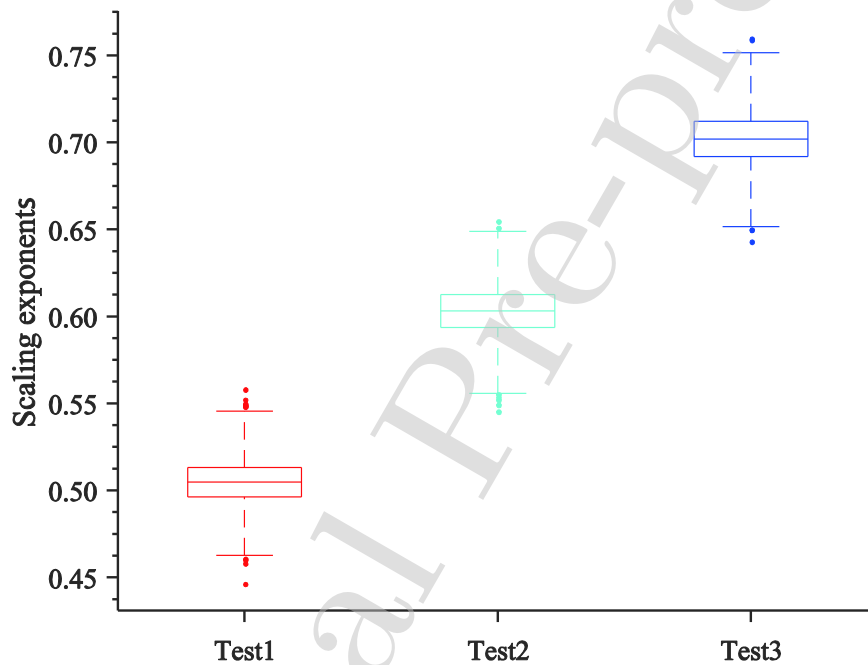
173 For long-term corrected time series, the fluctuation function and timescale  $s$  exhibit  
 174 that the power law relationship  $F(s) \sim s^\alpha$  and  $F(s)$  will increase with the  
 175 timescale  $s$ .  $\alpha$  represents the scaling exponent of time series. If  $0.5 < \alpha < 1$

176 ( $0 < \alpha < 0.5$ ), the time series exhibit a positive (negative) long-term correlation  
177 characteristics. If  $\alpha=0.5$ , the time series are uncorrelated. In this paper, the  
178 second-order DFA (DFA2) technique is adopted, which has been commonly employed  
179 to study the long-term memory characteristics of climate variability (Zhao and He,  
180 2015; He et al., 2015; Shen et al., 2017).

### 181 **2.3 Uncertainties in the DFA results**

182 The long-term memory characteristics of a given climate variable can be  
183 quantitatively described by using the DFA method. However, there are some  
184 uncertainties due to the length of the observed data. Therefore, the uncertainties  
185 should be estimated with the quantitative analysis of the long-term memory  
186 characteristics of precipitation series over southern China. Because the scaling  
187 exponent of precipitation series is between 0.5 and 0.7, the uncertainty of  $\alpha=0.5$ ,  
188  $\alpha=0.6$  and  $\alpha=0.7$  are examined as follows. Each test produces 20000 artificial  
189 series (10000 artificial series for daily precipitation and another 10000 artificial series  
190 for seasonal precipitation) with the given scaling exponents ( $\alpha=0.5$ ;  $\alpha=0.6$ ;  $\alpha=0.7$ )  
191 by Fourier filtering (Peng et al., 1991; Lennartz and Bunde et al., 2011), and the  
192 length of the artificial series is the same as that of the precipitation series over  
193 southern China used in this study. By using the DFA, scaling exponents of the  
194 artificial series are shown in Figure 1. For  $\alpha=0.5$ , 99% of the scaling exponents are  
195 between 0.46 and 0.54 as 0.5% to 99.5% range; the result of scaling exponents has a  
196 deviation ranging from -0.04 to 0.04 in the confidence interval of 99%. For  $\alpha=0.6$ ,  
197 99% of the scaling exponents are between 0.56 and 0.65 as 0.5% to 99.5% range; that

198 is to say, the result of scaling exponents has a deviation ranging from -0.04 to 0.05 in  
 199 the confidence interval of 99%. For  $\alpha=0.7$ , 99% of the scaling exponents are  
 200 between 0.65 and 0.75 as 0.5% to 99.5% range, indicating that the result of scaling  
 201 exponents has a deviation ranging from -0.05 to 0.05 in the confidence interval of  
 202 99%. According to the above discussions, we define the variation ranging from -0.05  
 203 to 0.05 as the 0.01 significant test level of precipitation series over southern China.



204  
 205 **Figure.1** The DFA analysis of the artificial series generated with given scaling  
 206 exponents by Fourier filtering.

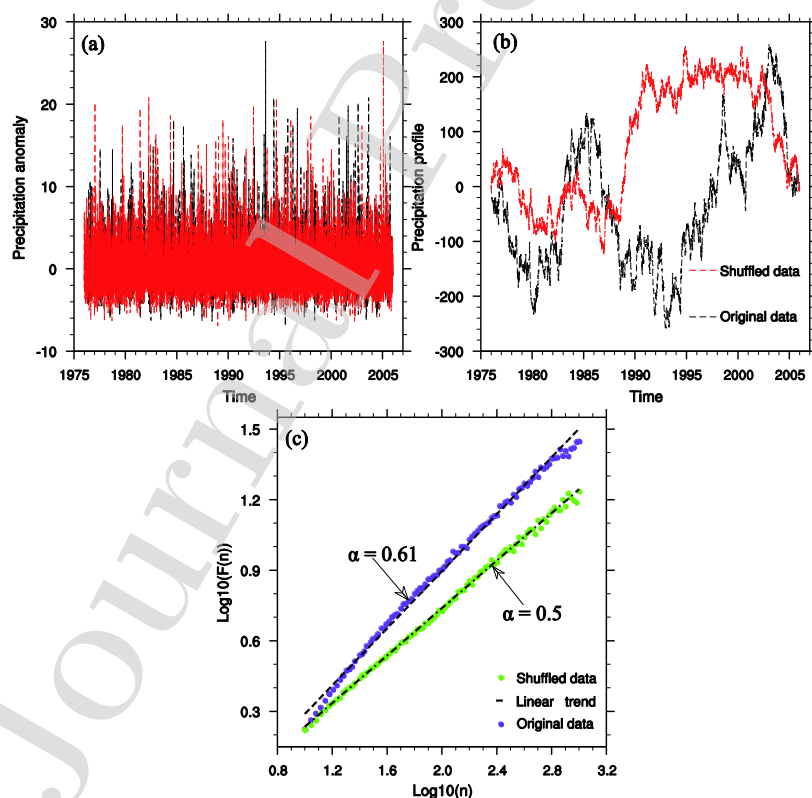
207

### 208 **3 Results**

#### 209 **3.1 Scaling behaviors of precipitation series**

210 The temporal evolution and profile of regional average precipitation time series over  
 211 southern China during 1976-2005 are shown in Figure 2a and 2b, respectively. We

212 notice that the precipitation anomaly series exhibit irregular high-frequency  
 213 fluctuations and their profiles show the inter-annual oscillation. The double  
 214 logarithmic curve of the fluctuation function  $F(s)$  with timescales is shown in Figure  
 215 2c for the time series of regional average precipitations over southern China. By linear  
 216 fitting, it is seen that the precipitation time series (regional average) over southern  
 217 China exhibit relatively weak long-term correlation characteristics with scaling  
 218 exponents of 0.61. Moreover, the scaling exponent of randomly shuffled precipitation  
 219 records is 0.5 with white noise characteristics, which further verifies the fact that the  
 220 long-term correlation characteristics of precipitation series over southern China are  
 221 caused by the fractal characteristics.



222

223 **Figure.2** (a) The regional mean precipitation anomaly time series over southern  
 224 China (the black dashed line represents raw precipitation time series and the red

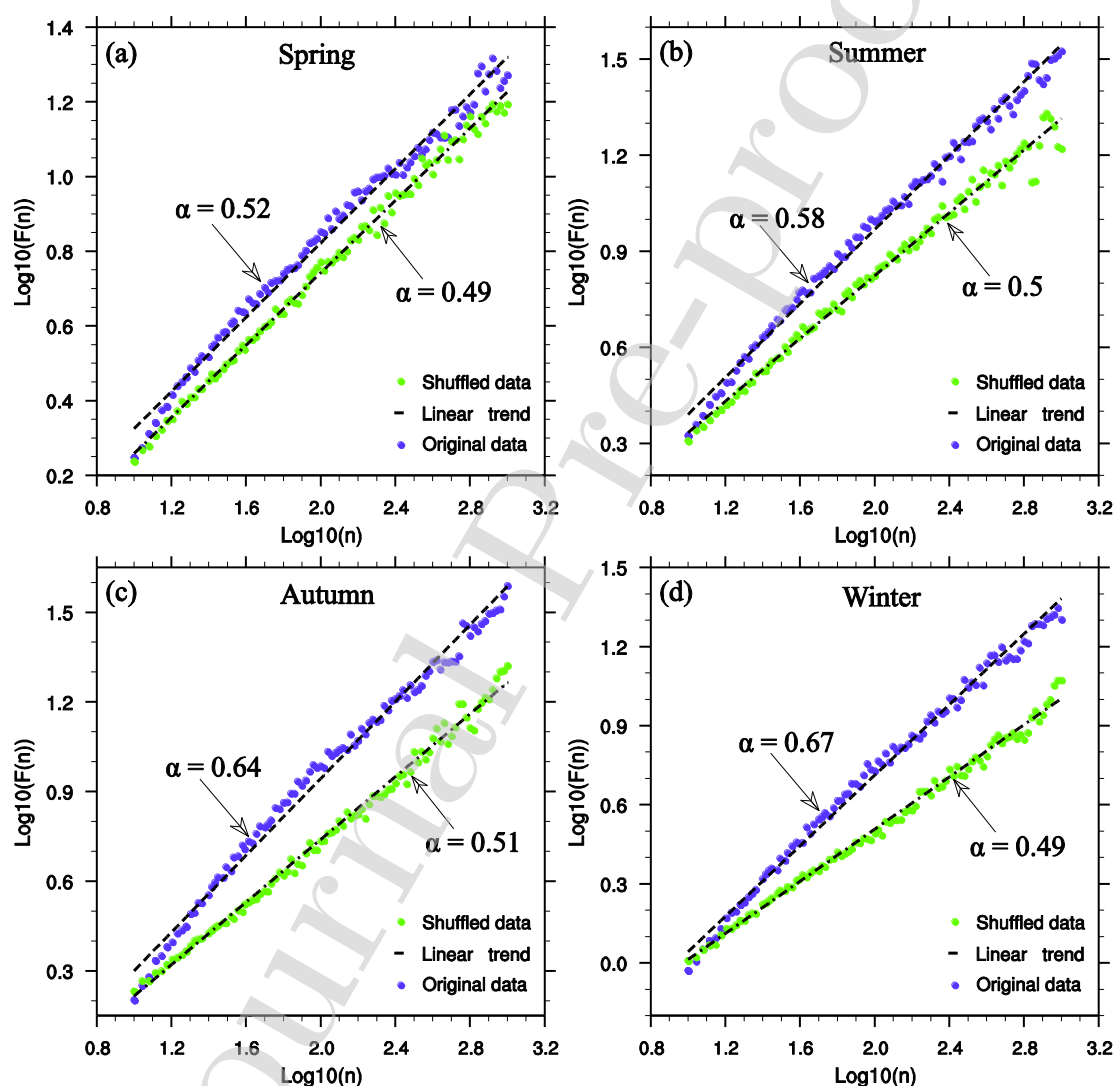
225 dashed line represents randomly shuffled precipitation time series). **(b)** The  
226 cumulative deviation for raw (black dashed line) and randomly shuffled (red dashed  
227 line) precipitation time series over southern China. **(c)** The DFA analysis results for  
228 annual precipitation time series (regional mean) over southern China from 1976 to  
229 2005 (the purple (raw time series) and green (randomly shuffled time series) solid  
230 point represent the obtained fluctuation function; the black dashed line represents the  
231 linear fitting line of the red solid point, and the slope can be a good index for  
232 long-term memory characteristics on this timescale).

233

234 The scaling analysis of precipitation series (regional average) over southern China in  
235 four seasons is shown in Figure 3 to further investigate whether daily precipitation  
236 data has long-term correlation characteristics in all four seasons. The scaling  
237 exponents for precipitation series in the four seasons are 0.52, 0.58, 0.64 and 0.67,  
238 respectively. The more obvious long-term correlation characteristics are shown in  
239 autumn and winter, while the scaling exponents are close to 0.5 in spring and summer  
240 suggesting the uncorrelation. Figure 4 shows the spatial distribution of scaling  
241 exponents of the precipitation series over southern China. It is seen that the scaling  
242 exponents of annual precipitation records show a decreasing trend from the coast to  
243 the inland over southern China. The scaling exponents in the coastal areas,  
244 Guangdong, southern Anhui and southeastern Jiangxi provinces in particular, are  
245 close to 0.6 showing strong long-term correlation, while they are close to 0.5 showing  
246 the uncorrelation in inland areas, Chongqing and western Hubei province in particular.

247 The distribution feature for scaling exponents of precipitation series is closely related  
248 to the climatic conditions in the region. The precipitation in coastal areas is mainly  
249 affected by weather systems (e.g., subtropical high, southwest vortex) and climate  
250 factors (e.g., El Niño Southern Oscillation, Pacific Decadal Oscillation, Indian Ocean  
251 Dipole Mode and Atlantic Multi-decadal Oscillation), while precipitation in inland  
252 areas is mainly attributed to the re-evaporation of surface water, which is produced by  
253 the inland circulation system. The precipitation affected by weather systems is more  
254 organized. Hence it presents stronger long-term correlation. However, the  
255 precipitation produced by the inland circulation system is more random, so that it  
256 shows the less correlation. In spring, the scaling exponents of precipitation sequence  
257 are ranged from 0.48 to 0.52 in most areas of southern China, showing the  
258 uncorrelation. The regions with the scaling exponents between 0.58 and 0.62 are  
259 mainly located in Guangdong, Zhejiang, northwestern Anhui and southeastern Henan  
260 provinces indicating that the precipitation records exhibit positive long-term  
261 correlation in these regions. However, the scaling exponents are smaller than 0.46 in  
262 the northeastern Chongqing and southwestern Hubei province showing weak negative  
263 long-term correlation. In summer, the precipitation records exhibit obvious positive  
264 long-term correlation with the scaling exponents greater than 0.62 in the downstream  
265 areas of the Yangtze River. At the same time, the scaling exponents are close to 0.5  
266 showing the uncorrelation in most of the other southern China regions. In autumn, the  
267 precipitation records exhibit relatively weak long-term correlation with the scaling  
268 exponents greater than 0.55 in most of the southern China regions. And the regions

269 showing strong long-term correlation located in the southeastern Guangxi, southern  
 270 Hunan and central Jiangxi provinces. In winter, the scaling exponents of precipitation  
 271 records are significantly greater than 0.66 exhibiting strong long-term correlation in  
 272 the provinces of Henan, southern Zhejiang, Fujian, Guangdong and Jiangxi, and  
 273 especially high in Guangdong province.

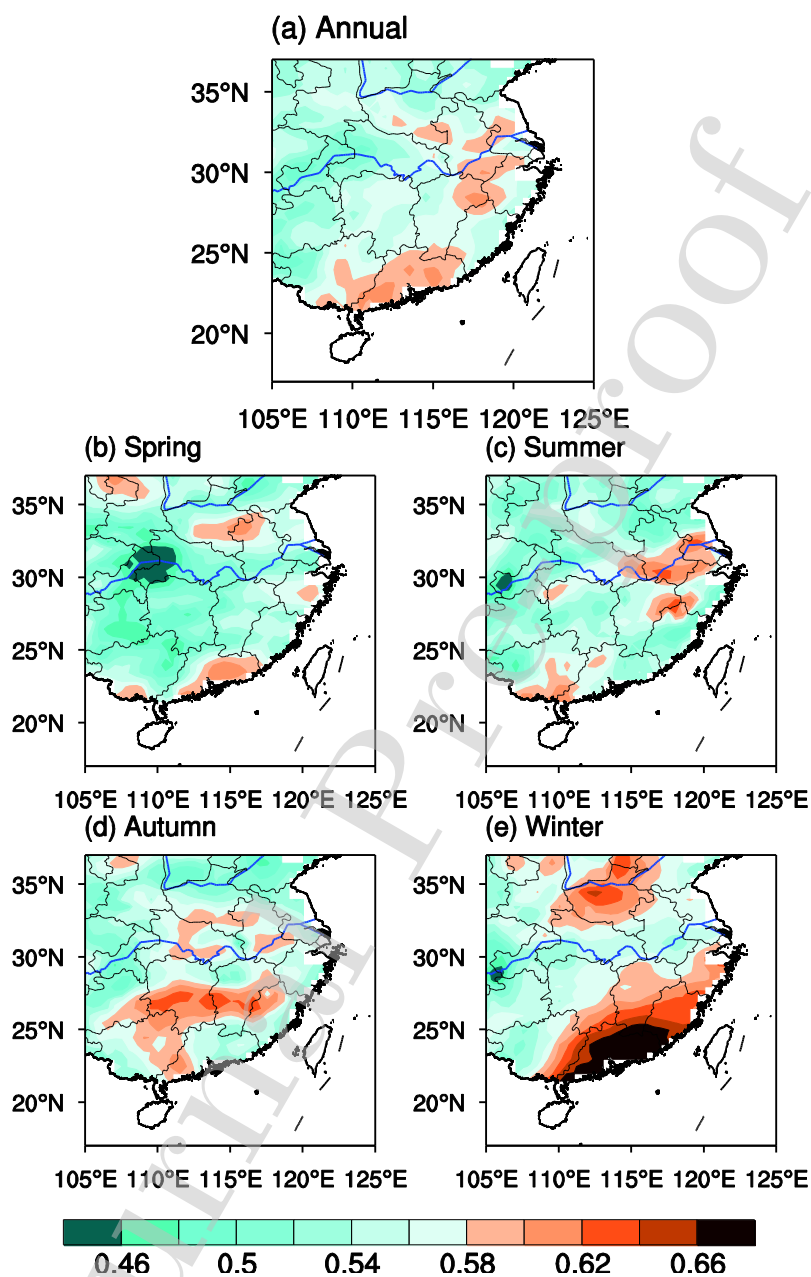


274

275 **Figure.3** The DFA analysis result for four seasons time series (regional mean) over  
 276 southern China from 1976 to 2005 (the purple (raw time series) and green (randomly  
 277 shuffled time series) solid point represent the obtained fluctuation function; the black  
 278 dashed line represents the linear fitting line of the red solid point, and the slope can be



279 a good index for long-term memory characteristics on this timescale); (a) Spring; (b)  
 280 Summer; (c) Autumn; (d) Winter.



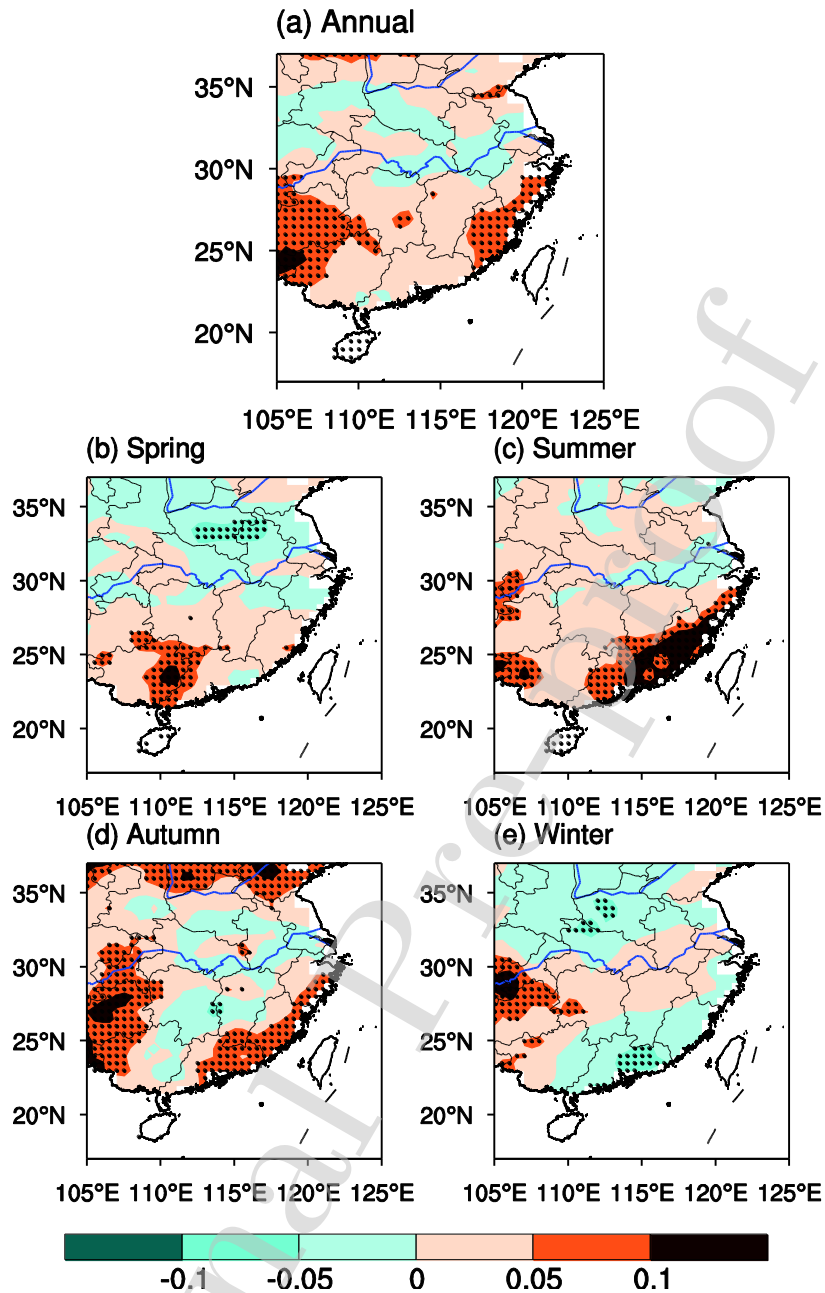
281

282 **Figure.4** The spatial pattern of scaling exponents for precipitation records from 1976  
 283 to 2005 over southern China in annual and four seasons. (a) Annual; (b) Spring; (c)  
 284 Summer; (d) Autumn; (e) Winter.

285

286 **3.2 Evaluation of the performance of CESM1**

287 Although the performance of CESM1 to simulate precipitations in China have been  
288 evaluated (Wang et al., 2017), the differences of scaling exponents of precipitation  
289 records between observations and simulations are shown in Figure 5 to evaluate the  
290 performance of CESM1 in simulating the internal dynamics characteristics of  
291 precipitation over southern China. Compared with the observation, the simulated  
292 scaling exponents of annual precipitation records are close to observation in most of  
293 southern China regions, and only 9% of grid points pass the 99% errors significance  
294 test. However, in places like Guizhou, Fujian and western Guangxi provinces the  
295 simulated scaling exponents are apparently overvalued. As for the seasonal  
296 evaluations, the simulated scaling exponents of precipitation records in spring are  
297 larger in eastern Guangxi province and smaller in northwestern Anhui and central  
298 Henan provinces than the observed ones. In summer, the simulated results are  
299 significantly greater than observations in Guangdong and Fujian provinces with errors  
300 more than 0.1. When it comes to autumn, Guizhou, Chongqing, northeastern  
301 Guangdong, southern Shanxi and southern Shandong are regions with distinguishably  
302 higher simulated values. While in winter, the simulated scaling exponents are almost  
303 consistent with the observations in most of southern China regions and only 9% of  
304 grid points pass the 99% errors confidence test. Overall, there are 90%, 79%, 72% and  
305 91% of grid points display the almost identical scaling exponents with the observed  
306 scaling exponents in spring, summer, autumn and winter, respectively, indicating that  
307 the CESM1 has the ability to simulate the internal dynamics characteristics of  
308 precipitation series in southern China.



309

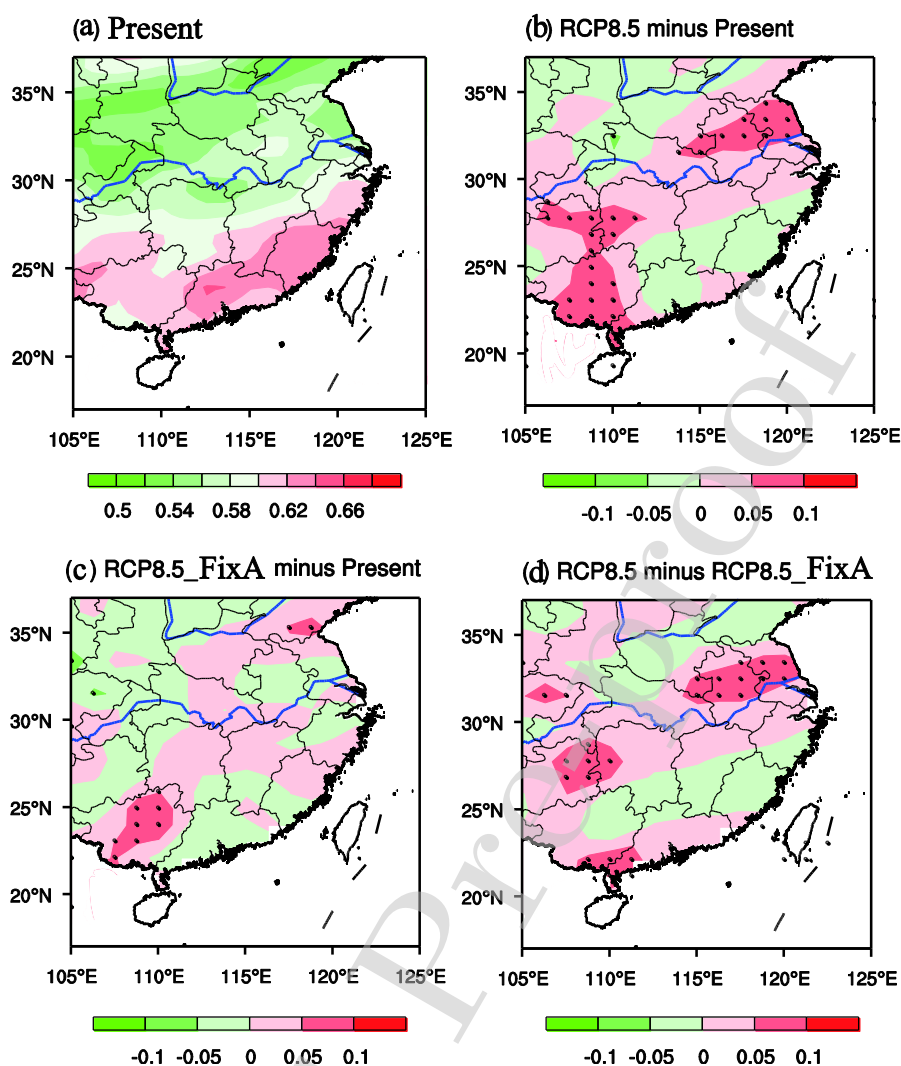
310 **Figure.5** The spatial pattern of the differences for scaling exponents between  
 311 observations and CESM1 simulated precipitation records from 1976 to 2005 over  
 312 southern China in annual and four seasons. (a) Annual; (b) Spring; (c) Summer; (d)  
 313 Autumn; (e) Winter (The black dots indicate that the differences of scaling exponents  
 314 between observation and simulation results pass the 99% significance test).

315

316 **3.3 The impacts of future reduced aerosols on internal dynamics characteristics**

317 **of precipitation series**

318 Figure 6 shows the spatial pattern of changes in scaling exponents of annual  
319 precipitation series over southern China during the late-21st century (2071-2100)  
320 under the RCP8.5 and RCP8.5\_FixA simulation scenarios in comparison with the  
321 present condition (1987-2016). It can be seen that the long-term correlation of  
322 precipitation records in Huai river basin will increase in values during the late-21st  
323 century (2071-2100) under the RCP8.5 simulation scenario (Figure 6b). However, the  
324 long-term correlation of precipitation records over all southern China regions are  
325 consistent with present condition (1987-2016) during the late-21st century  
326 (2071-2100) under the RCP8.5\_FixA simulation scenario (Figure 6c). Based on the  
327 differences between the RCP8.5 and RCP8.5\_FixA simulation scenarios (Figure 6d),  
328 we found that the future reduced aerosols will significantly strengthen the long-term  
329 correlation characteristics of precipitation records in Huai river basin.



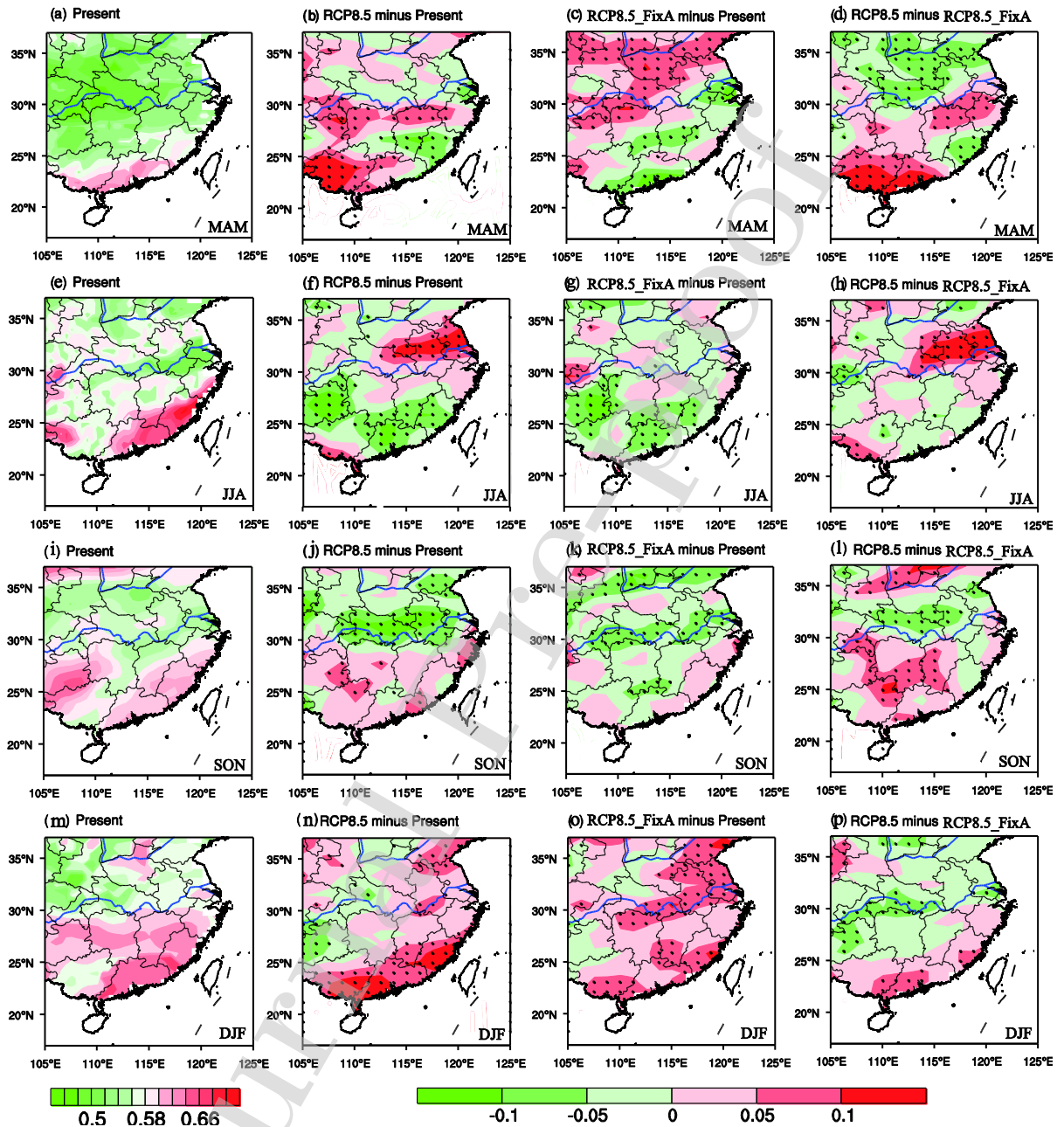
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331 **Figure.6** The spatial pattern of changes in scaling exponents of annual precipitation  
 332 series over southern China during the late-21st century (2071-2100) under the RCP8.5  
 333 and RCP8.5\_FixA scenarios in comparison with the present condition (1987-2016). (a)  
 334 The scaling exponents of the present condition (1987-2016); (b) the differences of  
 335 scaling exponents between the RCP8.5 and present scenarios; (c) the differences of  
 336 scaling exponents between the RCP8.5\_FixA and present scenarios; (d) the  
 337 differences of scaling exponents between the RCP8.5 and RCP8.5\_FixA scenarios.  
 338 The black dots indicate that the differences of scaling exponents between the two  
 339 scenarios pass the 99% significance test.

340 Figure 7 presents the variation characteristics of scaling exponents of precipitation  
341 series in the four seasons by using the DFA technique. In spring, the distribution of  
342 variations on scaling exponents is complicated in southern China during the late-21st  
343 century (2071-2100) under the RCP8.5 simulation scenario (Figure 7b). The  
344 significant change mainly occurs in the south of the Yangtze River. Augmentation of  
345 the long-term correlation can be found in Guangxi, northeastern Guizhou,  
346 northwestern Hunan, northern Jiangxi and Ningxia provinces while in Fujian and  
347 central Jiangxi provinces the long-term correlation declines. Under the RCP8.5\_FixA  
348 simulation scenario (Figure 7c), An obvious seesaw between the south and north of  
349 the Yangtze River can be observed in the long-term correlation of precipitation  
350 records, with distinct intensification in the north and reduction in the south. Figure 7d  
351 shows the differences of scaling exponents calculated by precipitation series between  
352 the RCP8.5 and RCP8.5\_FixA simulation scenarios. It indicates that the future  
353 reduced aerosols emissions will weaken the long-term correlation of precipitations in  
354 the north of the Yangtze River but strengthen the long-term correlation of  
355 precipitations in the most of southern Yangtze River regions, Guangdong and Guangxi  
356 provinces in particular. In summer, regions with the obvious change of scaling  
357 exponents of precipitation series located in Huai river basin and the scaling exponents  
358 will increase during the late-21st century (2071-2100) under the RCP8.5 simulation  
359 scenario in comparison with the present condition (1987-2016) (Figure 7f).  
360 Meanwhile, there is a noticeable decrease in the long-term correlation of precipitation  
361 records in a few areas (Guizhou and Guangdong provinces) both under the RCP8.5

362 (Figure 7f) and RCP8.5\_FixA simulation scenario (Figure 7g). By analyzing the  
363 differences of scaling exponents of precipitation series between the RCP8.5 and  
364 RCP8.5\_FixA simulation scenarios (Figure 7h), we can find that the future reduced  
365 aerosols emissions will contribute to reinforcing the long-term correlation of  
366 precipitation records in Huai river basin in summer since the scaling exponents  
367 increase by more than 0.1. In autumn, the long-term correlation of precipitation  
368 records will decrease in the north of the Yangtze River and slightly increase in the  
369 south under the RCP8.5 simulation scenario (Figure 7i). When it comes to the  
370 RCP8.5\_FixA simulation scenario (Figure 7k), more regions present the reduced  
371 long-term correlation during the late-21st century (2071-2100) in comparison with the  
372 present condition (1987-2016). Therefore, the future reduced aerosols emissions may  
373 result in reinforcing of the long-term correlation of precipitation records in southern  
374 Hunan and southwestern Jiangxi provinces and diminishing the long-term correlation  
375 of precipitation records in the north of Hubei and the south of Henan provinces  
376 (Figure 7l). In winter, the long-term correlation of precipitation records will increase  
377 in most regions of southern China under the RCP8.5 and RCP8.5\_FixA simulation  
378 scenarios during the late-21st century (2071-2100) in comparison with the present  
379 condition (1987-2016), especially in Guangxi, Guangdong and Fujian provinces under  
380 the RCP8.5 simulation scenario and in Huai river basin under the RCP8.5\_FixA  
381 simulation scenario (Figure 7n and 7o). The differences of scaling exponents of  
382 precipitation series between the RCP8.5 and RCP8.5\_FixA simulation scenarios in  
383 winter are less distinct as other seasons, suggesting the future reduced aerosols

384 emissions will generate less effect on the internal dynamics characteristics of  
 385 precipitation series in winter over southern China (Figure 7p).



386

387 **Figure.7** Same as the **Figure.6** but for four seasons. (a-d) Spring; (e-h)

388 Autumn; (m-p) Winter.

389

390 **4 Conclusion and possible mechanisms**

391 In this study, the scaling behaviors of precipitation records over southern China are



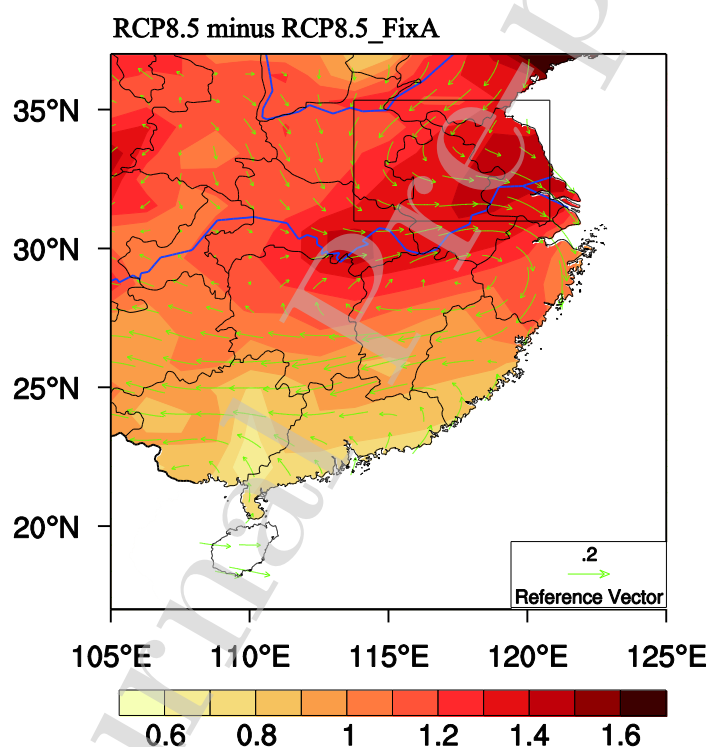
392 investigated by using the DFA method. It is found that the precipitation records over  
393 southern China exhibit relatively weak long-term correlation characteristics. The  
394 scaling exponents in coastal areas are close to 0.6 showing long-term correlation but  
395 the values in inland areas are close to 0.5 showing the uncorrelation. The shuffled  
396 precipitation time series is analyzed by the DFA technique to further verify the fact  
397 that the long-term correlation characteristics for precipitation series over southern  
398 China are caused by fractal characteristics. Based on the long-term correlation  
399 characteristics of the observed precipitation records, the performance of CESM1 in  
400 simulating precipitations over southern China is evaluated. The simulation results  
401 show that 90%, 79%, 72%, and 91% of grid points for precipitation records over  
402 southern China are close to observation results in the four seasons, respectively, which  
403 indicate that the CESM1 have the ability to simulate the internal dynamics  
404 characteristics of precipitation series in southern China. Based on simulated  
405 precipitation data by CESM1, the changes of scaling exponents of precipitation series  
406 over southern China are investigated during the late-21st century (2071-2100) in  
407 comparison with the present condition (1987-2016) under the climate scenarios  
408 RCP8.5 and RCP8.5\_FixA. Then the differences between the climate scenarios  
409 RCP8.5 and RCP8.5\_FixA are analyzed to assess the contribution of future reduced  
410 aerosols emissions on the internal dynamics characteristics of precipitation. It is found  
411 that there are obvious changes of internal dynamics characteristics of precipitation  
412 records over southern China during the late-21st century (2071-2100) under both the  
413 RCP8.5 and RCP8.5\_FixA simulation scenarios in comparison with the present

414 condition (1987-2016) in summer. During the late-21st century (2071-2100), the  
415 long-term correlation for precipitation will be intensified in Huai river basin under the  
416 RCP8.5 simulation scenario. Meanwhile, there is noticeable decrease in the long-term  
417 correlation of precipitation records in a few areas (Guizhou and Guangdong provinces)  
418 both under the RCP8.5 (Figure 7f) and RCP8.5\_FixA simulation scenario (Figure 7g).  
419 And the differences of scaling exponents of precipitation series between the RCP8.5  
420 and RCP8.5\_FixA simulation scenarios further indicate that the future reduced  
421 aerosols emissions will contribute to strengthening the long-term correlation of  
422 precipitation records in Huai river basin in summer with an increment of more than  
423 0.1 in the scaling exponents during the late-21st century (2071-2100) in comparison  
424 with the present condition (1987-2016).

425 The present paper demonstrated the fact that the future reduced aerosols emissions  
426 will contribute to strengthening the long-term correlation of precipitation records in  
427 Huai river basin in summer during the late-21st century (2071-2100) in comparison  
428 with the present condition (1987-2016). But the possible mechanisms of how the  
429 future reduced aerosols emissions manage to induce the long-term correlation of  
430 precipitation records to strengthen in Huai river basin in summer require further  
431 investigations and discussions. Previous studies have proposed that aerosols have a  
432 significant impact on precipitation through multiple processes. (i) Aerosols can  
433 directly change the energy of solar radiation reaching the land surface through its own  
434 optical properties. (ii) Aerosols have important microphysical effects on clouds and  
435 precipitation through their influence on cloud drop nucleation, which can influence

436 cloud lifetime, cloud albedo, and precipitation (Ramaswamy et al., 2001; Qian et al.,  
437 2009). (iii) Aerosols indirectly influence the precipitation through affecting the  
438 strength of the East Asian summer monsoon (Xie et al., 2016), weather systems (e.g.,  
439 subtropical high, southwest vortex) and climate factors (e.g., El Niño Southern  
440 Oscillation, Indian Ocean Dipole Mode and North Atlantic Oscillation). Figure 8  
441 shows the anomalies of temperature and 850hPa horizontal wind under the RCP8.5  
442 and RCP8.5\_FixA scenarios from 2071 to 2100 in southern China. The reduction of  
443 aerosols will increase the amount of solar radiation that reaches the land surface,  
444 which results in the strong positive temperature anomalies ( $\sim 1.4$  °C) in Huai river  
445 basin in summer. Therefore, more heat is available for evaporating water and  
446 energizing convective rain clouds. Meanwhile, more warming of the air above the  
447 surface can destabilize the low atmosphere and promote the generation of convective  
448 clouds. On the other hand, fewer aerosols mean relatively less cloud condensation  
449 nuclei (CCN), which is expected to speed up the conversion of cloud water into  
450 smaller raindrops and increase the precipitation from shallow and short-lived clouds.  
451 In addition, there is an anomalous cyclonic circulation over Huai river basin, which is  
452 favorable for wind convergence at the low troposphere (850hPa). The anomalous  
453 vertical upward movement combined with sufficient moisture can lead to more  
454 precipitation in Huai river basin. In summer, there is a large amount of precipitation in  
455 Huai river basin due to more favorable background conditions. Moreover, the  
456 anomalous precipitation brought by the future reduced aerosols emissions can further  
457 increase the intensity and duration of precipitation, which is beneficial to intensify the

458 correlation among precipitations in Huai river basin. To some extent, it can explain  
459 the fact that the future reduced aerosols emissions will contribute to strengthening the  
460 long-term correlation of precipitation records in Huai river basin in summer during  
461 the late-21st century (2071-2100) in comparison with the present condition  
462 (1987-2016). Although the current knowledge about the effects of aerosols on  
463 precipitations has made great progress, it is still very limited. Thus, more physical  
464 mechanisms of the impact of aerosols on internal dynamics characteristics of  
465 precipitation need to be further investigated in the future.



466  
467 **Figure.8** The climatic averaged differences of temperature (shading in °C) and  
468 850hPa horizontal wind (vector in m/s) between the RCP8.5 and RCP8.5\_FixA  
469 scenarios from 2071 to 2100 over southern China. The changes are significant based  
470 on a two sample student's test at the 95% confidence intervals.

471

472 Finally, it should be pointed out that long-term correlation is quite common in the  
473 climate system. The increase of scaling exponents of precipitation records in Huai  
474 river basin in summer indicates that cumulative historical effects on current  
475 precipitation state will become stronger with reduced aerosols emissions in the future,  
476 which is significant to further improve forecasting skills of precipitations in Huai river  
477 basin.

478

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485 (<https://www.earthsystemgrid.org/dataset/ucar.cgd.cesm4.output.html>).

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**\*Declaration of Interest Statement**

**Declaration of interests**

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

Journal Pre-proof