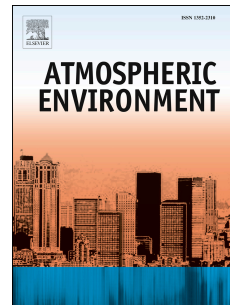


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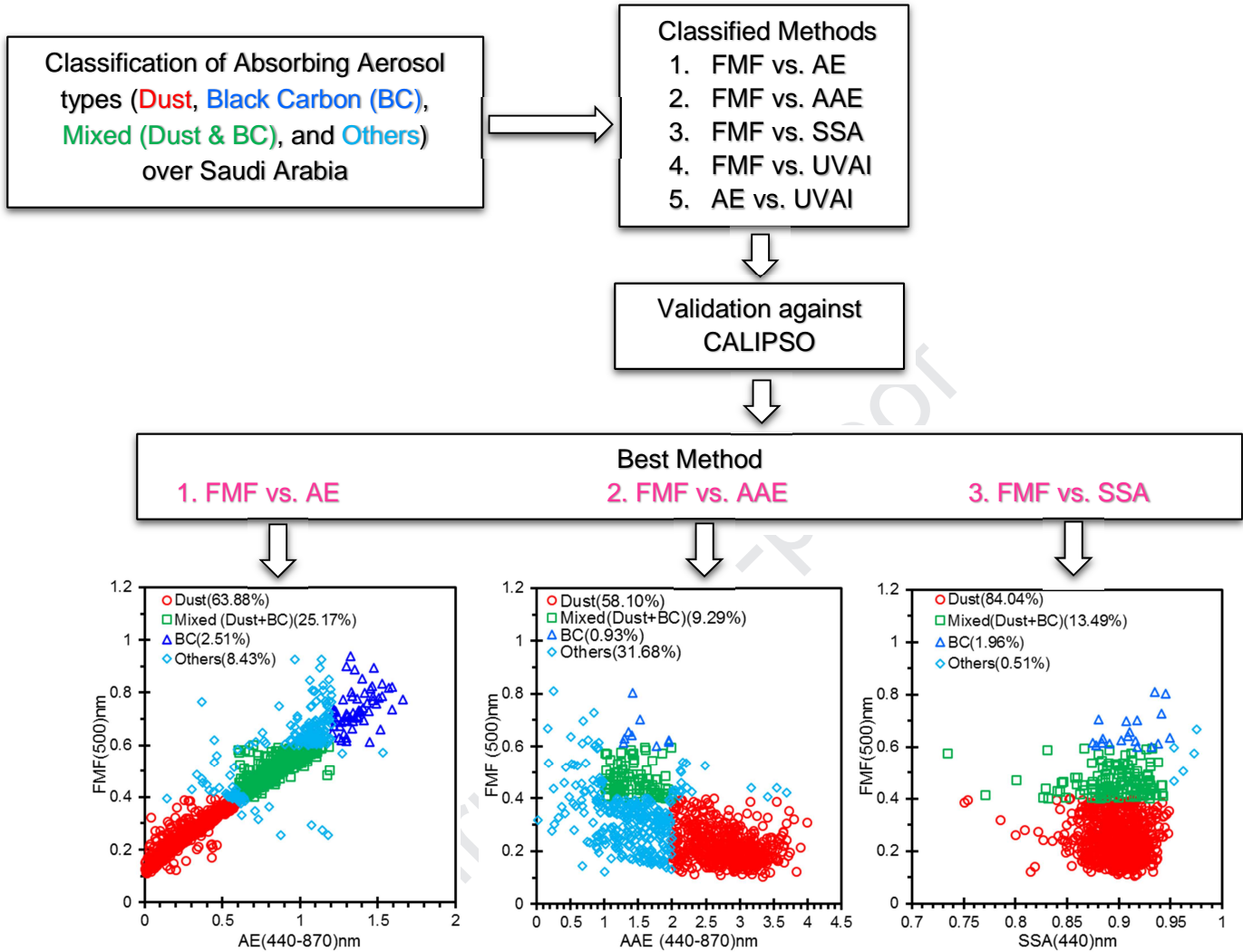
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CRedit authorship contribution statement

Md. Arfan Ali: Conceptualization, Data curation, Methodology, Formal analysis, Investigation, Validation, Visualization, Writing - original draft. **Janet E. Nichol:** Supervision, Investigation, Writing - review & editing. **Muhammad Bilal:** Conceptualization, Investigation, Visualization, Writing - review & editing. **Zhongfeng Qiu:** Supervision, Writing - review & editing. **Usman Mazhar:** Data curation. **Md Wahiduzzaman:** Data curation. **Mansour Almazroui:** Data curation. **M. Nazrul Islam:** Data curation.



Classification of aerosols over Saudi Arabia from 2004–2016

Md. Arfan Ali^{1,†}, Janet E. Nichol^{2,†}, Muhammad Bilal^{1*}, Zhongfeng Qiu¹, Usman Mazhar³,
Md Wahiduzzaman⁴, Mansour Almazroui⁵, M. Nazrul Islam⁵

¹School of Marine Sciences, Nanjing University of Information Science and Technology,
Nanjing 210044, China

²Department of Geography, School of Global Studies, University of Sussex, UK

³School of Remote Sensing and Geomatics Engineering, Nanjing University of Information
Science and Technology, Nanjing 210044, China

⁴Institute for Climate and Application Research, Nanjing University of Information Science and
Technology, Nanjing 210044, China.

⁵Center of Excellence for Climate Change Research/Department of Meteorology, King
Abdulaziz University, Jeddah 21589, Saudi Arabia

*Email: muhammad.bilal@connect.polyu.hk

†Authors with equal contributions.

Abstract

Knowledge of aerosol size and composition is very important for investigating the radiative forcing impacts of aerosols, distinguishing aerosol sources, and identifying harmful particulate types in air quality monitoring. The ability to identify aerosol type synoptically would greatly contribute to the knowledge of aerosol type distribution at both regional and global scales, especially where there are no data on chemical composition. In this study, aerosol classification techniques were based on aerosol optical properties from remotely-observed data from the Ozone Monitoring Instrument (OMI) and Aerosol Robotic Network (AERONET) over Saudi Arabia for the period 2004–2016 and validated using data from the Cloud-Aerosol Lidar and Infrared Pathfinder Satellite Observation (CALIPSO). For this purpose, the OMI-based

26 Aerosol Absorption Optical Depth (AAOD) and Ultra-Violet Aerosol Index (UVAI), and
27 AERONET-based AAOD, Ångström Exponent (AE), Absorption Ångström Exponent (AAE),
28 Fine Mode Fraction (FMF), and Single Scattering Albedo (SSA) were obtained. Spatial analysis
29 of the satellite-based OMI-AAOD showed the dominance of absorbing aerosols over the study
30 area, but with high seasonal variability. The study found significant underestimation by OMI
31 AAOD suggesting that the OMAERUV product may need improvement over bright desert
32 surfaces such as the study area. Aerosols were classified into (i) Dust, (ii) Black Carbon (BC),
33 and (iii) Mixed (BC and Dust) based on the relationships technique, between the aerosol
34 absorption properties (AAE, SSA, and UVAI) and size parameters (AE and FMF). Additionally,
35 the AE vs. UVAI and FMF vs. UVAI relationships misclassified the aerosol types over the study
36 area, and the FMF vs. AE, FMF vs. AAE and FMF vs. SSA relationships were found to be
37 robust. As expected, the dust aerosol type was dominant both annually and seasonally due to
38 frequent dust storm events. Also, fine particulates such as BC and Mixed (BC and Dust) were
39 observed, likely due to industrial activities (cement, petrochemical, fertilizer), water desalination
40 plants, and electric energy generation. This is the first study to classify aerosol types over Saudi
41 Arabia using several different aerosol property relationships, as well as over more than one site,
42 and using data over a much longer time-period than previous studies. This enables classification
43 and recognition of specific aerosol types over the Arabian Peninsula and similar desert regions.

44 **Keywords:** Aerosols; AERONET; Single Scattering Albedo; Absorption Ångström Exponent;
45 Ozone Monitoring Instrument; Aerosol Absorption Optical Depth.

46

47 **1 Introduction**

48 Atmospheric aerosol particles comprise solid and liquid materials differing in size from a
49 few nanometers to larger than 100 micro-meters, with intricate composition and volatility in their
50 physiochemical properties (Ali et al., 2019; Ali and Assiri, 2019; Almazroui, 2019). Over Asia,
51 an immense diversity of aerosol types exist, due to rapid industrialization and urbanization. This
52 creates uncertainty in assessing global climate change (Eck et al., 2010). Atmospheric aerosols
53 are considered a major element of the earth's climate system, as they remodel the climate and
54 radiative balance directly by scattering and absorbing incoming solar radiation (Ali et al., 2017),
55 whilst indirectly changing cloud optical properties and providing condensation nuclei (Kaufman

56 et al., 2005). Classification of aerosols into different types can improve the precision of radiative
57 balance, and assist climate modelling.

58 Aerosol types such as dust, organics, sea salt, and sulfate are predominantly reflective
59 and scatter incoming solar radiation back to space, thus cooling the atmosphere (Bilal et al.,
60 2013). However, other aerosols have more absorbing than scattering properties (Li et al., 2016).
61 The main absorbers in the aerosol mixture are iron oxides from dust, and Black Carbon (BC)
62 released from biomass burning and combustion processes and Brown Carbon (BrC) from organic
63 matter combustion (Wang et al., 2011). Moreover, iron oxides, BC and (BrC) show the greatest
64 absorption from the ultraviolet (UV) to the visible region (Eck et al., 2010; Liakakou et al.,
65 2020), while BC particles display constant absorption across the entire solar region (Bergstrom et
66 al., 2002). A thorough understanding of climate forcing due to aerosol requires knowledge of
67 aerosol concentration, its composition, size, and optical properties such as absorption or
68 scattering. The aerosol size distribution and absorption properties can be used to classify the
69 aerosols over the region (Higurashi and Nakajima, 2002; Lee et al., 2010). These properties vary
70 spatially and temporally (Choi et al., 2009) according to the season, emission sources, and
71 aerosol transportation (Ram et al., 2016).

72 The ground-based Aerosol Robotic Network (AERONET) provides aerosol absorptivity,
73 from the Absorption Ångström Exponent (AAE) at 440–870 nm and Single Scattering Albedo
74 (SSA) at 440 nm data. Complementing this, the Ozone Monitoring Instrument (OMI) on the
75 Aura satellite also provides aerosol absorbing properties such as the Ultra-Violet Aerosol Index
76 (UVAI) and Aerosol Absorption Optical Depth (AAOD) (Eq. 1) calculated in the UV and visible
77 bands (Adesina et al., 2016) (Table 1). Light absorbing particles (e.g., dust, BC, or BrC) in the
78 atmosphere can be determined by single scattering albedo (SSA) and absorbing aerosol optical
79 depth (AAOD) (Shin et al., 2019). The AAOD is the columnar aerosol loading (i.e. AOD) due to
80 light absorption based on the relationship

$$81 \quad \text{AAOD} = (1 - \text{SSA}) \times \text{AOD} \quad (1)$$

82 This is the most important parameter for the evaluation of atmospheric warming due to light
83 absorbing aerosols. Hu et al. (2016) reported that high AAOD levels commonly found over East

84 Asia result mainly from aerosol mixtures comprising desert dust, industrial pollutants, and smoke
85 from biomass burning. The study and classification of absorbing aerosols over the globe based
86 on AERONET and satellite observations is well established (Cazorla et al., 2013; Logan et al.,
87 2013; Logothetis et al., 2020; Kedia et al., 2014; Rupakheti et al., 2019, 2019a; Shen et al.,
88 2019). Dubovik et al. (2002) established the relationship technique, which uses relationships
89 between different optical properties of aerosols derived from AERONET and laboratory
90 measurements, for the classification of global aerosols. Thus the relationship techniques of FMF
91 (Fine Mode Fraction) vs. AE, FMF vs. AAE, FMF vs. SSA, AE vs. UVAI, and FMF vs. UVAI
92 can be used to distinguish the major aerosol types (Tables 1 and 2). Since then, studies have used
93 different relationship techniques, including FMF vs. AE (Eck et al., 2010), FMF vs. AAE (Giles
94 et al., 2011), FMF vs. SSA (Lee et al., 2010; Giles et al., 2012), AE vs. UVAI and FMF vs.
95 UVAI (Bibi et al., 2017) to classify aerosols into dust modes and BC). For example, low values
96 of FMF vs. AE indicate coarse mode dust aerosol (Aloysius et al., 2009); and high values of
97 FMF (> 0.6) and intermediate values of AAE ($1.0 < AAE < 2.00$) indicate BC aerosols (Giles et
98 al., 2011). Similarly, values of SSA ($SSA \leq 0.95$) and high values of FMF also indicate BC
99 aerosols (Lee et al., 2010; Giles et al., 2012). Several studies have used relationship techniques
100 from ground-based instruments alone, including Schmeisser et al. (2017), Jose et al. (2016) who
101 classified absorbing aerosols over Hyderabad, India, Alam et al. (2016) over urban areas of
102 Pakistan, and Gharibzadeh et al., (2018) over Iran. A validation and comparison of the
103 classifications done based on FMF vs. SSA are suggested to be included with the previous
104 studies done by Srivastava et al. (2012) and Tiwari et al. (2015) at different locations in India
105 and Pakistan. Bibi et al. (2016; 2017) classified aerosol types using ground-based and satellite-
106 based aerosol optical properties over Karachi (Pakistan) and the Indo-Gangetic Plain (IGP).

107 However, only a few such studies are available over the Middle-East. Of these, Farahat et
108 al. (2016) reported only a single aerosol type: dust, over the Middle-East and North Africa. Al-
109 Salihi (2018) classified aerosols based on AOD and AE relationship and over only one site,
110 Baghdad, Iraq, reporting four different aerosol types (maritime, dust, urban, and biomass
111 burning). The few aerosol classification studies conducted over Saudi Arabia have used only
112 one site in Saudi Arabia, the Solar Village, as part of larger studies in other regions. For
113 example, Logothetis et al. (2020) classified aerosols into eight types (Fine (highly, moderately,

114 slightly, and non-absorbing), mixed (absorbing and non-absorbing), coarse (absorbing and non-
115 absorbing)) based on FMF, SSA, and AE relationships over Europe, the Middle East, North-
116 Africa and Arabian Peninsula. Kaskaoutis et al. (2007) classified aerosols into four types (clean
117 maritime, biomass burning-urban, desert dust, and mixed) using relationship techniques over
118 four continents and other studies include Chen et al. (2016) and Mao et al. (2019) who also
119 included Saudi Arabia's Solar Village, Riyadh site as part of a larger study. None included the
120 KAUST Campus site in Jeddah, which is situated at the other side of the country (Figure 1), and
121 thus could offer a wider perspective of aerosol properties. Because Saudi Arabia has distinctive
122 geographical and climatic environments, which differentiate it from other countries, accurate
123 classification of aerosols cannot rely on universal classifications. High aerosol concentrations
124 over Saudi Arabia have traditionally been attributed to frequent dust storms (Awad and Mashat,
125 2014; Almazroui et al., 2015; Awad et al., 2015; Kumar et al., 2018; Ali and Assiri, 2016, 2019;
126 Mashat et al., 2019, 2020). However, the booming oil, and gas industry generating
127 unprecedented economic growth, have stimulated, rapid urbanization and industrial

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138 **Table 1** Definition of aerosol optical properties and relationship indicators.

Index designation	Name	Indicator
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AAOD	Aerosol Absorption Optical Depth	Columnar aerosol loading of light absorbing aerosols
AE	Ångström Exponent	Indicates the size of the dominant aerosol particles in the column (AE < 1 specifies dominance of coarse mode and AE > 1 demonstrates the dominance of fine mode aerosol)
AAE	Absorption Ångström Exponent	Measures the spectral dependence of absorption (UV to NIR), which depends on size, shape, and chemical composition of aerosols. Dust and BC (absorbing aerosols) have high values > 2.
FMF	Fine Mode Fraction	Provides quantitative information about the proportion of coarse and fine mode aerosol particles. FMF < 0.40 = coarse mode, $0.4 \leq \text{FMF} \leq 0.6$ = mixture, FMF > 0.60 = fine mode
SSA	Single Scattering Albedo	The ratio of scattering to extinction, and indicates the proportion of absorbing versus scattering aerosol particles SSA > 0.95 = non-absorbing, SSA ≤ 0.95 = absorbing aerosols
UVAI	Ultra-Violet Aerosol Index	A robust index for detecting absorbing aerosols (dust and soot) in the atmosphere. Uses 2 UV wavebands. UVAI > 1.0 = the enhanced presence of UV-absorbing, UVAI = 0.5-1.0 = weak presence of UV-absorbing aerosols.
FMF vs. AE	Low values of both = Dust, Medium values = Mixed (BC and Dust), High values = BC	
FMF vs. AAE	High values of AAE and low values of FMF indicate Dust, Medium values of both indicate Mixed (BC and Dust), Low values of AAE and High values of FMF indicate BC	
FMF vs. SSA	Low values of FMF and Medium values of SSA indicate Dust, Medium values of FMF and Low SSA indicate Mixed (BC and Dust), and High values of both indicate BC.	
FMF vs. UVAI	Low values of FMF and High values of UVAI indicate Dust, Medium FMF and Low UVAI indicate Mixed (BC and Dust), High values of FMF and Low values of UVAI indicate BC	
AE vs. UVAI	Low values of AE and High values of UVAI indicate Dust, Medium AE, and Low UVAI indicate Mixed (BC and Dust), High values of AE, and Low values of UVAI indicate BC.	

141

142 activities (i.e., cement, petrochemical, fertilizer, water desalination, and electric energy
 143 generation plants). The outcomes are unquantified in terms of human health, as it is known that
 144 different aerosol types vary in their health impacts. Several studies have reported health impacts
 145 from heavy metals and pathogens accompanying dust storms over the Middle East. For example,
 146 Leili et al. (2008) examined total suspended particles and PM₁₀ over the center of Tehran (Iran),
 147 and reported heavy metal contents (Pb, Co, Cd, Cu, and Cr) at levels dangerous enough to cause

148 neurodevelopmental and behavioral defects in children. Other findings of high heavy metal
149 content of airborne dust in the Middle East include Foroushani et al. (2019) in western Iran, and
150 Farahmandkia et al. (2010) in Tehran. Studies, which linked heavy metals in dust with serious
151 human health concerns, include Leili et al. (2008) in Tehran, Jiries et al. (2003) in Amman,
152 Jordan, and Al-Rajhi et al. (1996) in Riyadh. In addition to heavy metals, Gerivani et al. (2011)
153 found that dust storms in Iran can contain and transport viruses, which affect human populations
154 and Saeedi et al. (2012) reported dust particles containing polycyclic aromatic hydrocarbons
155 (PAHs) in Teheran. A potentially beneficial impact of dust particles, is their role in carrying
156 nutrients to the marine ecosystem of the Northern Red Sea (Jeddah) and their contribution to
157 nutrient balance continues largely unexplored (Prakash et al., 2015). The amplified threats of
158 climate change for desert animals world-wide are magnified in Saudi Arabia (Williams et al.,
159 2012). Thus, the classification of aerosols over the Arabian Peninsula for accurate estimates of
160 climate forcing and health impacts is urgent. This study uses available ground-based AERONET
161 sites within Saudi Arabia i.e., the Solar Village and KAUST Campus sites, combined with
162 satellite data from the Ozone Monitoring Instrument (OMI) to classify the predominant types of
163 aerosols over Saudi Arabia.

164 The main contributions of the current research are (a) the long-term period and spatial
165 spread of observations that makes the results more robust and (b) the selection of the most
166 appropriate technique (classification scheme) for the determination of different aerosol types
167 over Arabia, based on comparison with CALIPSO.

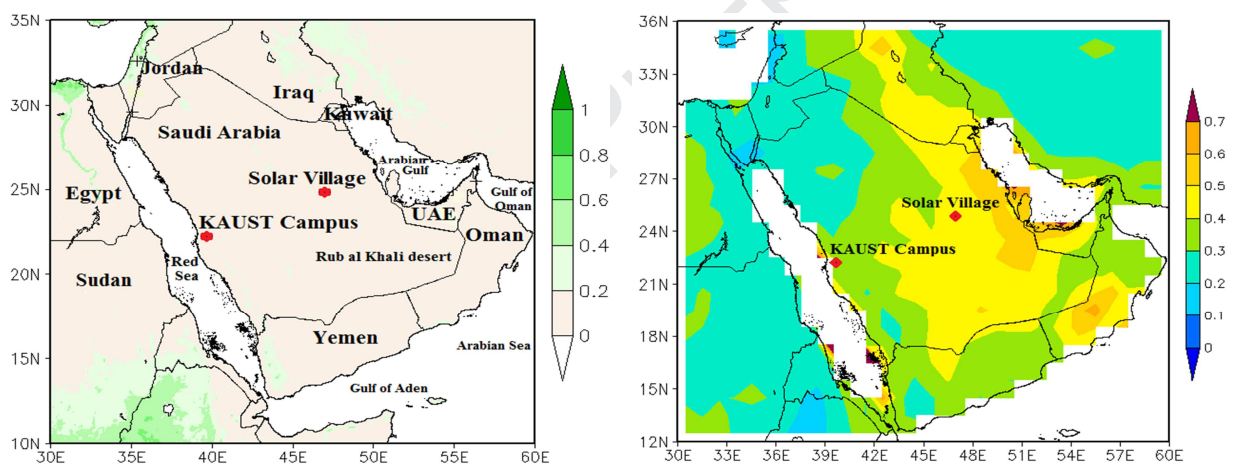
168 **2 Study area and Data-sets**

169 *2.1 Study Area*

170 Saudi Arabia is the largest country in the Middle-East, covering 80% of the Arabian
171 Peninsula with an area of approximately 2,218,000 km² (Figure 1), and is bordered by the
172 Arabian Gulf and the Red Sea. The largest desert, the Rub al Khali or Empty Quarter covers
173 647,500 Km² in the southern part of the country and is a source of frequent dust outbreaks and
174 severe dust storms. The country comprises 13 provinces and a total of 104 cities, of which the 20
175 largest have over 100,000 residents. This study was accomplished over the two AERONET
176 ground stations in Saudi Arabia: Solar Village and KAUST Campus (Figure 1). The Solar
177 Village (24.91° N, 46.41° E and 764 m a.s.l) is approximately 50 km from the north-west

178 periphery of Riyadh. This city is situated on the desert plateau resulting in frequent dust storm
 179 events (Farahat, 2016). The KAUST Campus site (22.30° N, 39.10° E and 11.2 m a.s.l) is
 180 positioned in the village of Thuwal in a rural and coastal site in the Red Sea on the roof of a
 181 building. Climatically, the country has little rainfall, with approximately 100 mm between
 182 autumn and early spring, followed by hot and dry late spring and summer. The Shamal winds
 183 lead to dust events during spring (Mashat et al., 2020) and summer (Notaro et al., 2013, 2015;
 184 Yu et al., 2013, 2015, 2016), with over thus the KAUST Campus and Solar Village sites (Figure
 185 1) observe higher aerosol loadings at these times.

186 Based on the three-month mean readings of temperature and rainfall, Saudi Arabia's
 187 seasons are classified into spring from March to May (MAM), summer from June to August
 188 (JJA), autumn from September to November (SON), and winter from December to February
 189 (DJF) (AMS, 2001).



190 **Figure 1:** Geographical map of the Kingdom of Saudi Arabia. Red asterisks represent the two
 191 ground-based AERONET stations. Whereas, the left panel represents the mean values (2004–
 192 2016) of Normalized Difference Vegetation Index (NDVI) and the right panel represents the
 193 long-term (2004–2016) mean values of aerosol optical depth (AOD) based on the MODIS
 194 Collection 6.1 Deep Blue algorithm.

196 2.2 Data-sets

197 OMI is carried by the Aura satellite, launched in July 2004, and is designed to measure
 198 air quality, the earth's climate, and ozone. It measures sunlight scattered by aerosols with high
 199 spectral resolution from the ultraviolet to visible regions (270–500 nm) and a spatial resolution

200 of 13–24 km (Levelt et al., 2006). The OMI near-UV aerosol retrieval algorithm (OMAERUV)
201 can be used to measure prominent absorbing aerosols such as dust and carbonaceous aerosols
202 (Torres et al., 2007). The OMAERUV algorithm utilizes the near-UV spectral region for
203 estimation of AAOD and UVAI products. Of major interest in the near-UV measurements is the
204 powerful interaction between aerosol absorption and scattering in this spectral region, which
205 facilitates the calculation of aerosol absorption capacity. For this study, OMI Level 2 and Level 3
206 OMAERUV OMI AAOD (500 nm) data were obtained from
207 <https://giovanni.gsfc.nasa.gov/mapss/> and "<https://giovanni.gsfc.nasa.gov/giovanni/>",
208 respectively.

209 The Cloud-Aerosol Lidar and Infrared Pathfinder Satellite Observation (CALIPSO) was
210 launched on 28th April 2006 on the CloudSat satellite to study the roles of aerosols and clouds in
211 earth's air quality, weather and climate. The CALIPSO gives information on aerosol vertical
212 profiles and 3-dimensional information of aerosol properties throughout day and night over the
213 globe (Winker et al., 2003), based on the Cloud-Aerosol Lidar with Orthogonal Polarization
214 (CALIOP) sensor. The aerosol lidar ratio, a key parameter for extinction retrieval, is determined
215 for each aerosol subtype based on measurements, modelling, and the cluster analysis of a
216 multiyear Aerosol Robotic Network (AERONET) dataset (Omar et al., 2005, 2009), they are
217 considered more accurate than other measurements (Su et al., 2020). In version 3 (V3) and
218 earlier, the CALIOP level 2 aerosol classification and lidar ratio selection algorithm defined six
219 aerosol types: clean marine, dust, polluted continental, clean continental, polluted dust, and
220 smoke (Omar et al., 2009). Each type is assigned an extinction-to-backscatter ratio (i.e., lidar
221 ratio) with an associated uncertainty that defines the limits of its expected natural variability
222 (<https://www.atmos-meas-tech.net/11/6107/2018/amt-11-6107-2018.pdf>). This study used the
223 Level 2 CALIPSO version 4.10 aerosol-type profiles for aerosol classification. These images
224 were downloaded from [https://www-](https://www-calipso.larc.nasa.gov/products/lidar/browse_images/std_v4_index.php)
225 [calipso.larc.nasa.gov/products/lidar/browse_images/std_v4_index.php](https://www-calipso.larc.nasa.gov/products/lidar/browse_images/std_v4_index.php). The temporal resolution
226 of 16 days for CALIPSO makes it unsuitable for continuous monitoring, but due to its accuracy,
227 it is used in this study for validation.

228 The AERONET is NASA's ground-based aerosol network, which has more than 700
229 stations over the globe (Holben et al., 1998). The data are commonly used for validating satellite-

230 based aerosol retrievals. In this study, version-3 Level 2.0 (cloud-screened and quality-assured)
 231 daily averaged direct sun products (FMF_{440nm} and AE_{440-870nm}) and sky irradiance products
 232 (AAOD_{440nm}, SSA_{440nm}, and AAE_{440-870nm}) were obtained from <https://aeronet.gsfc.nasa.gov/> for
 233 the period 2004–2016 (Table 2).

234 **Table 2** Observation and the total number of datasets at each AERONET site.

Location	Observation	Total	
		Direct products AE/FMF	Inversion products SSA/AAE/AAOD
Solar Village	2004–May 2013	2463/2549	1001
KAUST Campus	Feb 2012–2016	1580/1393	1285

235

236 **2.3 Research Methodology**

237 The methodology of the present study is as follows: Mean seasonal and annual spatial
 238 distributions of OMI-AAOD were calculated from daily observations for the period 2004–2016.
 239 AERONET AAOD at 440 nm was interpolated to AAOD at 500 nm using the Ångström
 240 Exponent (Equation 1):

$$AAOD_{500\text{ nm}} = AAOD_{440\text{ nm}} \times \left(\frac{500}{440}\right)^{-AAE_{440-870}} \quad (2)$$

241 Monthly and seasonal temporal analyses were performed for the AERONET data (AAOD, AE,
 242 AAE, FMF, and SSA) and OMI data (AAOD and UVAI). For validation purposes, to obtain
 243 collocated OMI-AAOD with AERONET, the OMI AAOD values were averaged for a spatial
 244 window of 1×1 pixel centered over the Solar Village and KAUST Campus sites, and AERONET
 245 values were averaged for +/- 30mins of the overpass time of OMI. Similarly, OMI UVAI
 246 collocated retrievals were obtained which were used for the classification of aerosol types. In the
 247 present study, a total of five relationships such as FMF vs. AE (Logothetis et al., 2020), FMF vs.
 248 SSA (Logothetis et al., 2020; Lee et al., 2010), AE vs. UVAI and FMF vs. UVAI (Bibi et al.,
 249 2017) were used. Besides, the FMF vs. AAE relationship is modified based on several previously
 250 published studies (Lee et al., 2010; Bibi et al., 2017; Rupakheti et al., 2019; Logothetis et al.,
 251 2020). The above relationships classified aerosols into three main categories (Table 3), namely
 252 (1) dust, (2) mixed dust and black carbon (BC), and (3) BC (nearly exclusively attributed to
 253 fossil-fuel emissions, industrial and traffic). The remaining data points, which do not fall within

254 the classification thresholds, are denoted as other aerosol types. Finally, the identified aerosol
 255 types were confirmed by comparison with satellite aerosol products from CALIPSO datasets.

256 **Table 3** Classification of aerosol types over Saudi Arabia using threshold values taken from
 257 previous studies.

Aerosol Types	FMF vs AE		FMF vs AAE		FMF vs SSA	
Dust	FMF<0.4	AE<0.6	FMF<0.4	AAE>2.0	FMF<0.4	SSA≤0.95
Mixed (BC & Dust)	0.4≤FMF≤0.6	0.6≤AE≤1.2	0.4≤FMF≤0.6	1.0<AAE<2.0	0.4≤FMF≤0.6	SSA≤0.95
BC	FMF>0.6	AE> 1.2	FMF>0.6	1.0<AAE<2.0	FMF>0.6	SSA≤0.95
	AE vs UVAI			FMF vs UVAI		
Dust	0.0<AE<0.4		UVAI>1.57		0.1<FMF<0.3	
Mixed (BC & Dust)	0.0<AE<1.0		0.5<UVAI<1.55		0.1<FMF<0.55	
BC	1.0<AE>1.55		0.5<UVAI<1.52		0.55<FMF>1.0	
						0.5<UVAI<1.50

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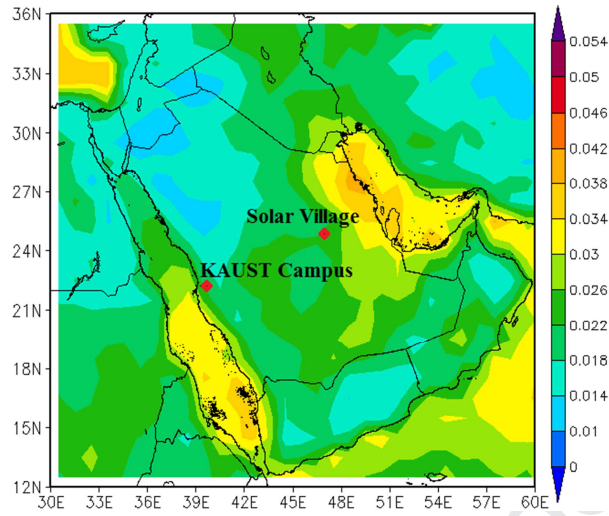
259 **3 Results and Discussion**

260 **3.1 Spatial distribution of OMI-based AAOD**

261 Figures 2 and 3 show the mean annual and seasonal Level-3 OMI-AAOD retrievals at 500 nm of
 262 over Saudi Arabia for 2004–2016. Figure 2 shows high AAOD values (greater than 0.03) over
 263 the Eastern provinces, moderate AAOD (0.018 to 0.03) over most parts of the country, and low
 264 AAOD (0.01 to 0.018) in the North-Western region. These AAOD values are less than 10% of
 265 the columnar AOD values, which suggests that absorbing aerosols are much fewer than
 266 scattering aerosols over Saudi Arabia. High AAOD is mainly distributed near the sources of dust,
 267 BC, and OC (Islam et al., 2019; Kang et al., 2017). The seasonal distributions (Figure 3) show
 268 the highest AAOD (greater than 0.03) in spring, and over the Eastern and Southern provinces,
 269 followed by summer, winter, and autumn. This is because dust storms originate in the Sahara
 270 Desert due to depressions passing eastwards over the Mediterranean Sea, and strong ground
 271 heating produces turbulence, local pressure gradients, and the Shamal (wind) pattern (Shao,
 272 2008; Prakash et al., 2015; Mashat et al., 2019). April to May (spring) experiences by peak
 273 dustiness over Eastern regions, and May-June over Southern and Central regions of Saudi Arabia
 274 (Sabbah and Hasan, 2008; Yu et al., 2013). However, an anticyclonic pattern is developed in
 275 autumn leading to weak dust activity resulting lowest columnar AAOD over Saudi Arabia (Kang
 276 et al., 2017; Mashat et al., 2019).

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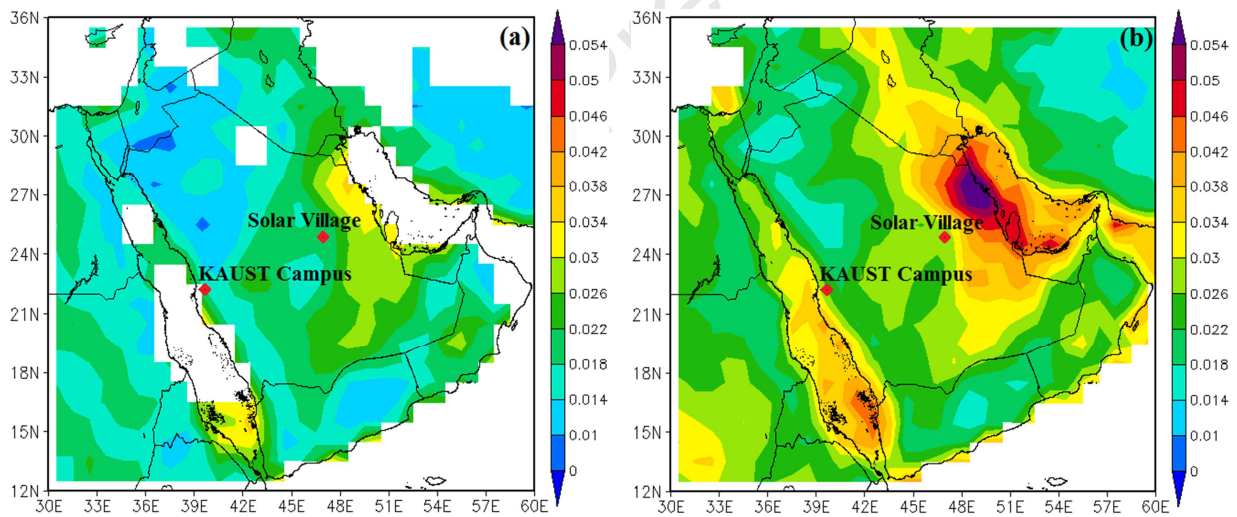


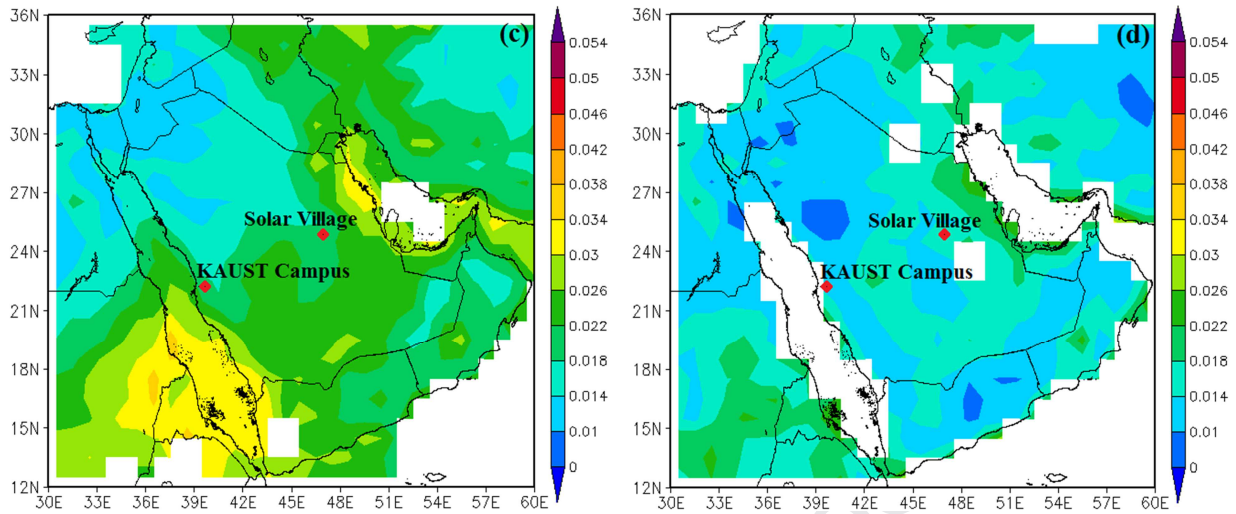
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Figure 2: Annual mean Aerosol Absorption Optical Depth (AAOD) obtained from the OMI instrument over Saudi Arabia averaged over the period 2004–2016.

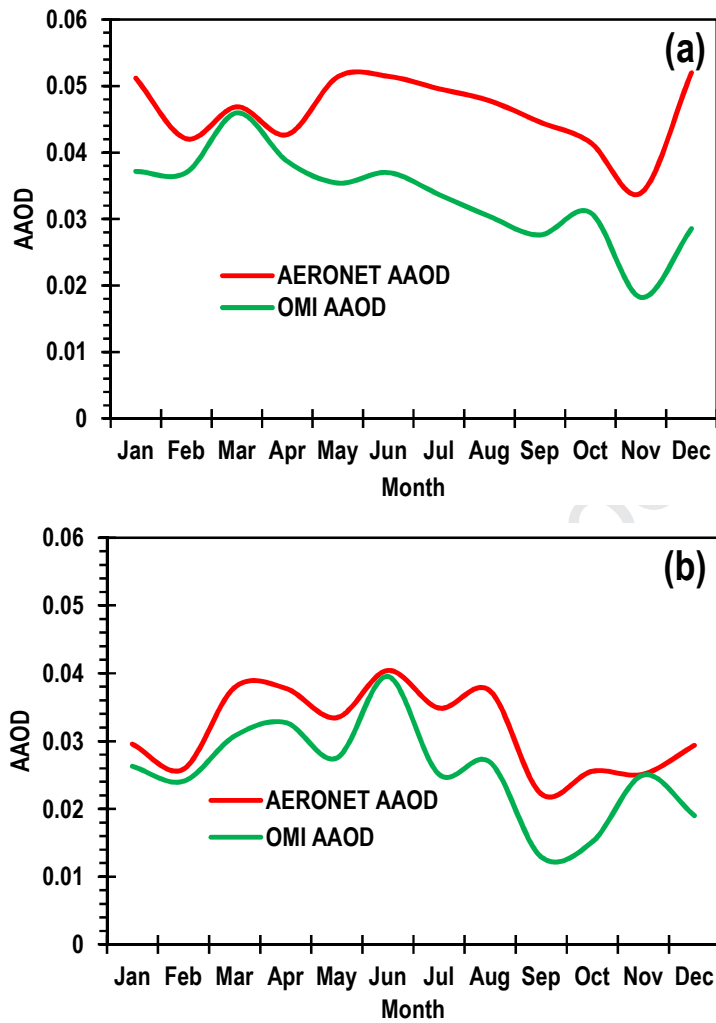
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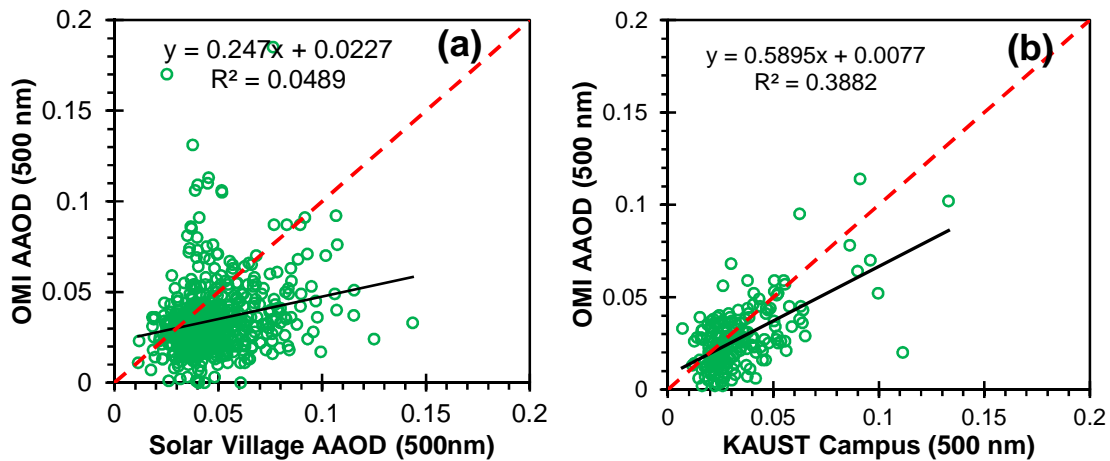


282 **Figure 3:** Mean seasonal spatial distribution of Aerosol Absorption Optical Depth (AAOD) for
 283 (a) Winter, (b) Spring, (c) Summer, and (d) Autumn obtained from OMI instrument over Saudi
 284 Arabia averaged over the period 2004–2016.

285 Figure 4 shows a more detailed annual cycle of both AERONET- and OMI-based Level-
 286 2 AAOD over the Solar Village (2004–2013) and KAUST Campus (2012–2016) sites. Higher
 287 values of AERONET and OMI AAOD retrievals over Solar Village in the east of the Peninsula
 288 indicate more absorbing aerosols present than over the KAUST Campus site. This may be due to
 289 a larger number of dust storm events compared to the KAUST region, as reported by Butt et al.,
 290 2017 using ground-based meteorological data. Results showed that OMI-AAOD retrievals
 291 followed the same temporal pattern as the AERONET-AAOD measurements (Figure 4). Figure 5
 292 shows significant underestimations for both low and high aerosol loadings as indicated by lower
 293 values of the slope, which suggested the inappropriate use of the aerosol model as well as error
 294 in the estimated surface reflectance. The underestimation during the low aerosol loadings is
 295 caused by the overestimation in the estimated in the surface reflectance (Bilal et al., 2013; Bilal
 296 and Nichol, 2015). Error in these parameters might be responsible for the underestimation in the
 297 AAOD retrieved by the OMAERUV algorithm. This may suggest that improvements in the OMI
 298 algorithm (OMAERUV) are required for a better estimation of AAOD over bright desert
 299 surfaces.



300 **Figure 4:** Annual cycle of Aerosol Absorption Optical Depth (AAOD) obtained from
301 AERONET and Level 2 OMI instrument over the (a) Solar Village (2004–2013) and (b) KAUST
302 Campus (2012–2016).
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304

305 **Figure 5:** Scatterplot between OMI AAOD and AERONET AAOD) over the (a) Solar Village

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(2004–2013) and (b) KAUST Campus (2012–2016).

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Figure 6 represents the annual cycle of AERONET- and OMI-based aerosol optical properties over the Solar Village and KAUST Campus sites for the period 2004–2016. These properties describe both aerosol size and absorptivity, including Ångström Exponent (AE), Absorption Ångström Exponent (AAE), Fine Mode Fraction (FMF), Single Scattering Albedo (SSA), and Aerosol Index (UVAI). The AE indicates the size of the dominant aerosol particles in the column, where small values of AE (< 1) indicate the dominance of coarse mode aerosols and large values of AE (> 1) demonstrate the dominance of fine mode aerosol such as BC, sulfate, and organic carbon released from manmade activities (Eck et al., 1999). The annual values of AE (Table 4) suggest coarse mode aerosols over both AERONET sites (Solar Village: 0.48, KAUST Campus: 0.64), as well as in all seasons. AE reaches its minimum in May (Solar Village: 0.20, KAUST Campus: 0.35) and maximum in November (Solar Village: 0.86, KAUST Campus: 0.97) (Figure 6). These results suggest substantially more coarse mode aerosols in spring compared with other seasons. Trend analysis showed no significant increasing or decreasing trends in AE over either site (Table 4).

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The Absorption Ångström Exponent (AAE) indicates the absorption contrast in relation to wavelength, which depends on particle size, shape, and chemical composition of the absorbing aerosols, which have a unique value (Russel et al., 2010; Li et al., 2016). For example, values of $AAE < 2$ and $AAE > 2$ indicate the fine mode and coarse mode absorbing aerosols respectively. Annual average values of AAE suggest coarse mode absorbing aerosols over both AERONET

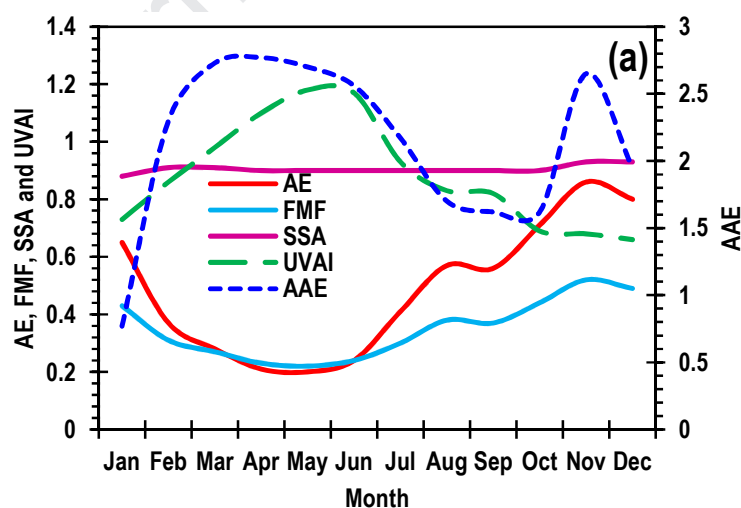
326 sites (Solar Village: 2.19, KAUST Campus: 2.27) (Table 4). However, significant variations
327 were observed at monthly scales, with coarse mode absorbing aerosols observed in spring, but
328 fine mode absorbing aerosols in winter over both sites (Figure 6). Overall, no significant
329 decreasing/increasing trends in AAE were observed over either site (Table 4).

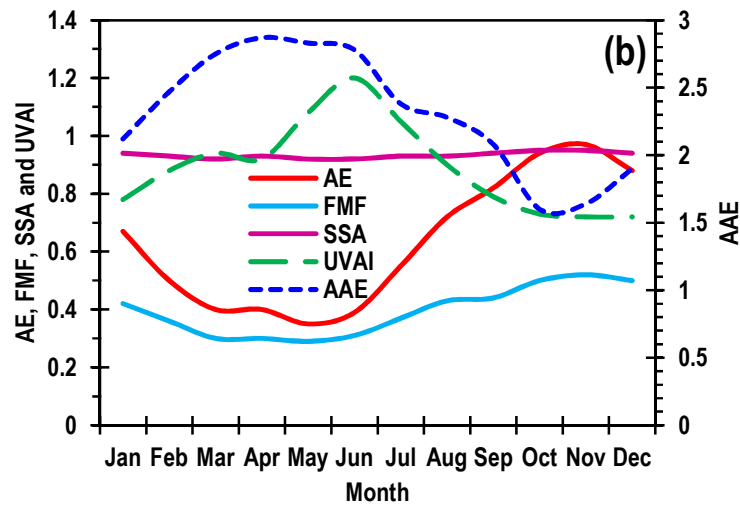
330 The Fine Mode Fraction (FMF) provides quantitative information about the proportion of
331 coarse and fine mode aerosol particles, varying from 0 (coarse mode aerosols) to 1 (fine mode
332 aerosols). According to Lee et al. (2010) and Logothetis et al. (2020), $FMF < 0.40$ represents
333 coarse mode aerosols, $0.40 \leq FMF \leq 0.60$ represents mixed type (coarse and fine mode) aerosols
334 and $FMF > 0.60$ represents fine mode aerosols. The annual average value of FMF (0.34) in Solar
335 Village is lower than that in KAUST Campus (0.40) and corresponds to more coarse mode
336 aerosols over Solar Village. The seasonal average value of FMF (spring: 0.24, summer: 0.31,
337 autumn: 0.44, winter: 0.38) in Solar Village is lower compared to that in KAUST Campus
338 (spring: 0.30, summer: 0.37, autumn: 0.48, winter: 0.43), which suggest more coarse-mode
339 aerosols over Solar Village (Table 4). Annual values of FMF were lower at SV than at KAUST
340 Campus, due to lower levels in spring and early summer. These lower levels of FMF at Solar
341 Village indicated more coarse mode aerosols, compared to mixed aerosols for these months at
342 KAUST Campus. Therefore, overall annually, Solar Village experiences more coarse mode
343 aerosols than KAUST Campus (Figure 6). Overall, no significant decreasing/increasing trends in
344 FMF were observed over the years 2004–2016 (Table 4).

345 The Single Scattering Albedo (SSA) is the ratio of scattering to extinction and indicates
346 the proportion of absorbing versus scattering aerosol particles. The value of $SSA > 0.95$
347 describes non-absorbing aerosols, $0.90 \leq SSA \leq 0.95$ indicates weakly absorbing aerosols, $0.85 <$
348 $SSA < 0.90$ for moderately absorbing aerosols, and $SSA < 0.85$ belongs to highly absorbing
349 aerosols (Lee et al., 2010; Russel et al., 2010; Gyawali et al., 2012; Shin et al., 2019). The annual
350 average values of SSA within the range of 0.85–0.95 suggest the presence of weakly absorbing
351 aerosols over both AERONET sites (Solar Village: 0.90, KAUST Campus: 0.93) (Table 4). At
352 seasonal scale, weakly absorbing aerosols were observed during all seasons except during
353 summer over the Solar Village site. Weakly absorbing aerosols indicate both dust and organic
354 carbon, the latter being a complex mixture of chemical compounds generated from fossil fuel
355 and biofuel burning as well as from natural biogenic emissions. The absorption or scattering

356 property of dust grains depends on their size and composition, whether predominantly
 357 silicate or graphite, thus these results are compatible with the results of FMF and AE,
 358 which suggest coarse aerosols to be dominant at both sites. Overall, no significant
 359 decreasing/increasing trends in SSA were observed (Table 4).

360 The Ultra-Violet Aerosol Index (UVAI) is a well-known index for detecting the
 361 absorbing aerosols (dust and biomass burning) in the atmosphere. It uses the UV spectrum to
 362 distinguish absorbing from non-absorbing aerosols (Graaf et al., 2005). The threshold UVAI >
 363 0.5 is useful to identify absorbing aerosols (Torres et al., 2009). The value of UVAI > 1.0 shows
 364 the enhanced presence of UV-absorbing aerosols (e.g., dust or smoke or biomass burning), and
 365 $0.5 < \text{UVAI} < 1.0$ indicates the weak presence of UV-absorbing aerosols (Washington et al.,
 366 2003). The observed values of UVAI suggest the weak presence of UV-absorbing aerosols over
 367 both AERONET sites except in spring at Solar Village and summer at KAUST Campus (Table
 368 4) and this is confirmed by the monthly values (Figure 6). These findings support Kaskaoutis et
 369 al. (2010) report of dust particles as indicated by the AI (0.5 to 0.6) over the South Greek sea
 370 regions. Overall, a significant increasing trend in UVAI was observed at KAUST Campus (Table
 371 4).





372 **Figure 6:** Annual cycle of averaged Aerosol Optical Properties obtained from AERONET (AE,
 373 AAE, FMF, and SSA), and OMI (UVAI) over the (a) Solar Village site (2004–2013) and (b)
 374 KAUST Campus site (2012–2016).

375

376 As can be seen from the above aerosol descriptions in this section 3.2, the analysis based
 377 on the individual parameters describe, whether fine or coarse mode (AE and FMF), whether fine
 378 or coarse mode absorbing aerosols (AAE), and whether absorbing or non-absorbing aerosols
 379 (SSA and UVAI). However, these individual parameters cannot identify the exact nature of the
 380 aerosol types such as dust or BC or mixed. Therefore, section 3.3 evaluates the combination of
 381 these parameters to classify aerosols into specific types.

382

383

384 **Table 4** Mean seasonal and annual variability of aerosol optical properties (AE, AAE, FMF, SSA, and UVAI) with their trends over
 385 the Solar Village and KAUST Campus sites for the period 2004–2016.

Parameters	Solar Village						KAUST Campus					
	Winter	Spring	Summer	Autumn	Annual	Trends	Winter	Spring	Summer	Autumn	Annual	Trends
AE	0.54±0.17	0.23±0.12	0.41±0.17	0.70±0.18	0.48±0.25	-0.012	0.68±0.20	0.39±0.08	0.55±0.17	0.91±0.12	0.64±0.24	0.023
AAE	2.14±0.93	2.72±0.36	2.18±0.70	1.70±0.66	2.19±0.71	-0.044	2.19±0.40	2.80±0.28	2.48±0.48	1.77 ± 0.47	2.27±0.57	0.067
FMF	0.38±0.08	0.24±0.07	0.31±0.08	0.44±0.08	0.34±0.11	-0.008	0.43±0.07	0.30±0.04	0.37±0.07	0.48±0.05	0.40±0.09	0.007
SSA	0.900±0.03	0.902±0.01	0.897±0.01	0.903±0.20	0.90±0.02	-0.001	0.94±0.01	0.92±0.01	0.93±0.01	0.95±0.01	0.93±0.01	0.002
UVAI	0.78±0.17	1.09±0.20	0.97±0.20	0.73±0.12	0.88±0.23	0.012	0.79±0.15	0.98±0.20	1.05±0.24	0.75±0.11	0.89±0.22	0.015*

386

387 *3.3 Classification of aerosols*

388 The relationships of different parameters, namely FMF vs. AE, FMF vs. AAE, FMF vs. SSA, AE vs. UVAI, and FMF vs.
 389 UVAI were used to classify aerosols into three types: Dust, Mixed (Dust and BC), and BC (Figure 7–10). Results based on FMF vs.
 390 AE, FMF vs. AAE, and FMF vs. SSA demonstrated the dominance of Dust type aerosols followed by Mixed (BC and Dust), then BC
 391 over both sites, (Figures 7–10). Thus on Figures 7 and 8 (a–c), the relationships showing the dominance of Dust type aerosols were
 392 FMF vs. SSA (Solar Village: 84.04%, KAUST Campus: 50.50%) followed by FMF vs. AE (Solar Village: 63.88%, KAUST
 393 Campus: 48.38%), and FMF vs. AAE (Solar Village: 58.10%, KAUST Campus: 46.45%). The results support observations of
 394 frequent dust storms (Kaskaoutis et al, 2007), as well as several studies reporting more dusty days over the Solar Village as compared
 395 to the KAUST Campus site (Yu et al., 2013; Butt et al., 2017), as many dust storms emanate from the desert of Iraq, North-East of
 396 Saudi Arabia, and Southern Iran, directly influencing to the Solar Village site (Prospero et al., 2003; Farahat et al., 2016). Therefore,
 397 dust aerosols are persistently prevalent over the Solar Village site. However, when we consider the relationships of AE vs. UVAI and

398 FMF vs. UVAI, these suggest the dominant aerosol type to be Mixed (BC and Dust) followed by Dust, and then BC (Figures 7 and 8
399 (d–e)). Since very

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400 low levels of Dust are indicated by the two relationships AE (Dust: $0.0 < AE < 0.4$) vs. UVAI
401 (Dust: > 1.57) and FMF (Dust: $0.1 < FMF < 0.3$) vs. UVAI (Dust: > 1.57) over both sites, whereas
402 the other three relationships show high dust levels, supported by the single parameters and by
403 many other studies and reports, these results suggest that AE vs. UVAI and FMF vs. UVAI
404 relationships cannot provide a meaningful aerosol types classification. This may be due to
405 underestimation by the OMAERUV algorithm-based UVAI data (see Figures 4 and 5), as the
406 UVAI alone suggested the dominance of absorbing aerosols over both sites. Therefore, the study
407 indicates that the OMAERUV algorithm may need improvement for better estimating the OMI
408 UVAI over bright-reflecting surfaces.

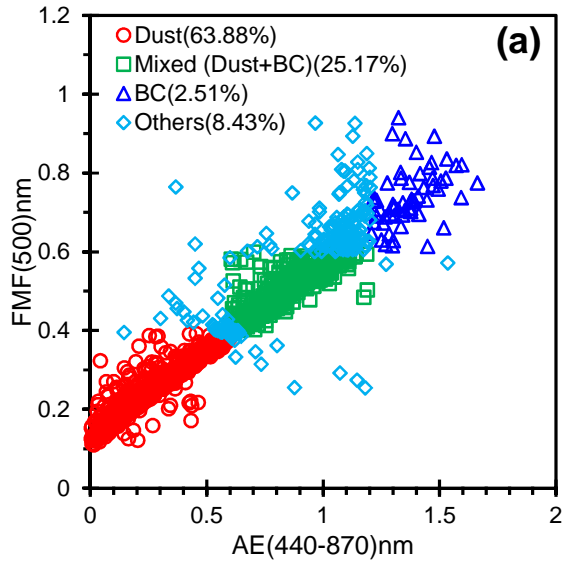
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410 The value of $FMF < 0.6$ demonstrates coarse-mode dominated aerosols, which were
411 associated with a mixture of different types of aerosols (Wu et al., 2015). Pérez-Ramírez et al.,
412 (2015), noted a similar finding over Granada, Spain. This interpretation is supported by Wu et al.
413 2015, who used a value of $FMF < 0.6$ to identify coarse-mode dominated aerosols which are
414 associated with a mixture of different types of aerosols, and a similar finding was reported by
415 Pérez-Ramírez et al., (2015) over Granada, Spain. The FMF ($0.4 \leq FMF \leq 0.6$) vs. AE ($0.6 \leq AE$
416 ≤ 1.2), FMF ($0.4 \leq FMF \leq 0.6$) vs. AAE ($1.0 < AAE < 2.0$), and FMF ($0.4 \leq FMF \leq 0.6$)
417 vs. SSA (≤ 0.95) thresholds represent Mixed (Dust and BC) type aerosols. These Mixed type
418 aerosols are best represented by FMF vs. AE (Solar Village: 25.17%, KAUST Campus: 41.21%),
419 with the FMF vs. SSA and FMF vs. AAE giving lower percentages (Figures 7 and 8 (a–c)).
420 Finally, Figures 7 and 8 (a–c) show a small percentages of BC aerosols based on the above-
421 mentioned three relationships over the both sites. Higher FMF (> 0.6) values indicate fine mode
422 aerosols, which correspond to BC, which may be due to local industrial activities (cement,
423 petrochemical, and fertilizer), water desalination plants, and electric energy generation (Farahat
424 et al., 2016). Some aerosol types were not classified namely 'Other' type (Figures 7–10). These
425 were best represented by FMF vs. AE (31.68%) over the Solar Village and by FMF vs. AAE
426 (29.97%) over the KAUST Campus site (Figures 7 and 8 (a–c)). These aerosols may be formed
427 due to the mixing of natural and anthropogenic aerosols atmospheric water vapor over the study
428 area (Kaskaoutis et al., 2011).

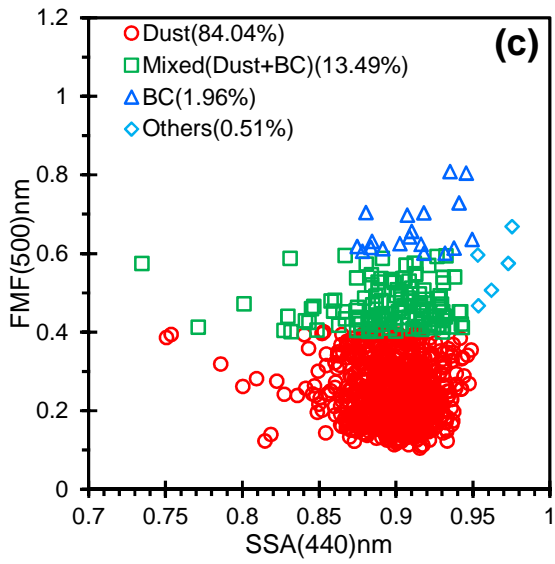
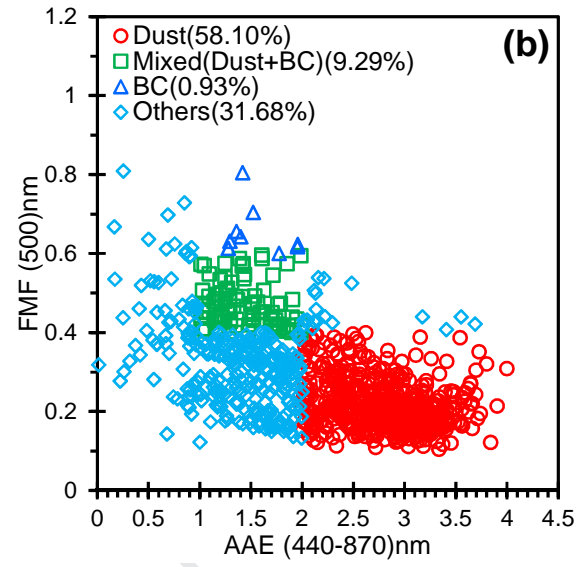
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430 The seasonal distribution of aerosol types from the three relationships FMF vs. AE, FMF
431 vs. AAE, and FMF vs. SSA confirm that Dust is the dominant aerosol type during all seasons
432 over both sites, and this reaches its maximum in spring followed by summer, winter, and autumn
433 (Figures 9 and 10 (a–c)). This confirms the findings of previous satellite- and station-based
434 studies which reported high dust levels during the peak dust storm season of spring and early
435 summer (Sabbah and Hasan, 2008; Yu et al., 2013; Farahat et al., 2016; Albugami et al., 2019).
436 Conversely, the lowest FMF values (< 0.3) were noted in spring and summer over both sites,
437 which support the findings of Kaskaoutis et al. (2007) and Wu et al. (2015). The above-
438 mentioned three relationships indicate Mixed (Dust and BC) type aerosols during all seasons,
439 with highest levels in autumn followed by winter, summer, and spring (Figures 9 and 10 (a–c)).
440 The possible reasons for this decline in dust storm events as well as washing out by the higher
441 rainfall during autumn to winter (Kaskaoutis et al., 2007; Farahat et al., 2016), resulting in FMF
442 values become a little higher varies from 0.46 to 0.50 indicate coarse-mode dominated aerosols
443 (Dust), which correspond to Mixed (Dust and BC) over the study area. The value of $FMF < 0.6$
444 demonstrates coarse-mode dominated aerosols, which were associated with a mixture of different
445 types of aerosols (Lee et al., 2010; Wu et al., 2015). Figures 9 and 10 also show BC aerosols to
446 be dominant during autumn and winter, which is attributed to local anthropogenic activities of
447 urban/industrial and biofuel emission. Consequently, the FMF values are increased to above 0.6,
448 and this is correlated with the observed fine-mode BC particles over the study area. These
449 findings also supported by Gautam et al. (2007), Wu et al. (2015), and Lee et al. (2010).

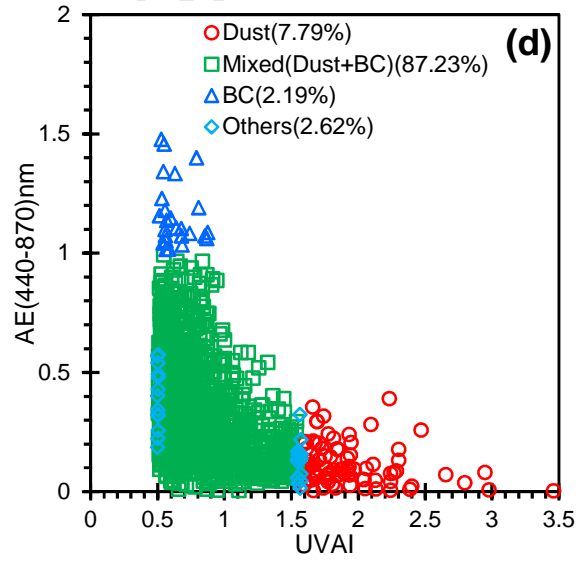
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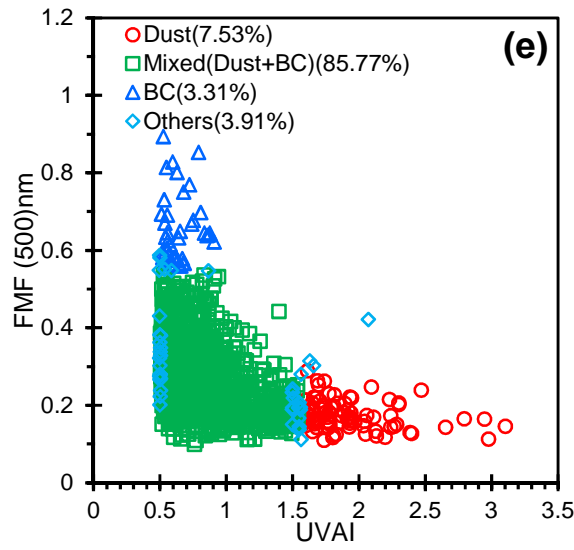


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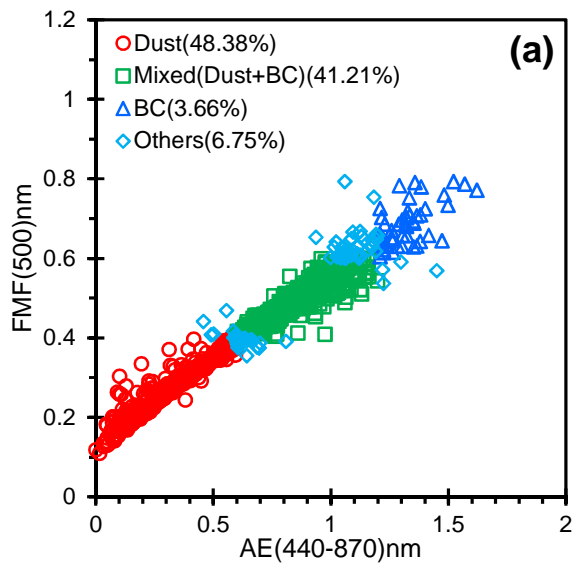
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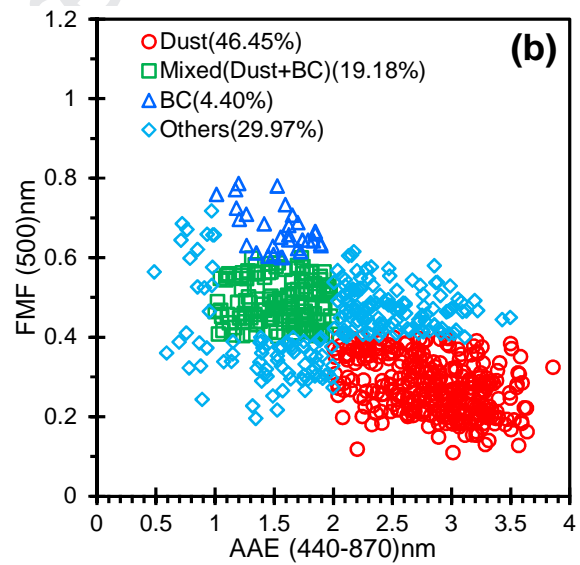
Figure 7: Aerosol classification over the Solar Village site during the period 2004-2013 for (a)

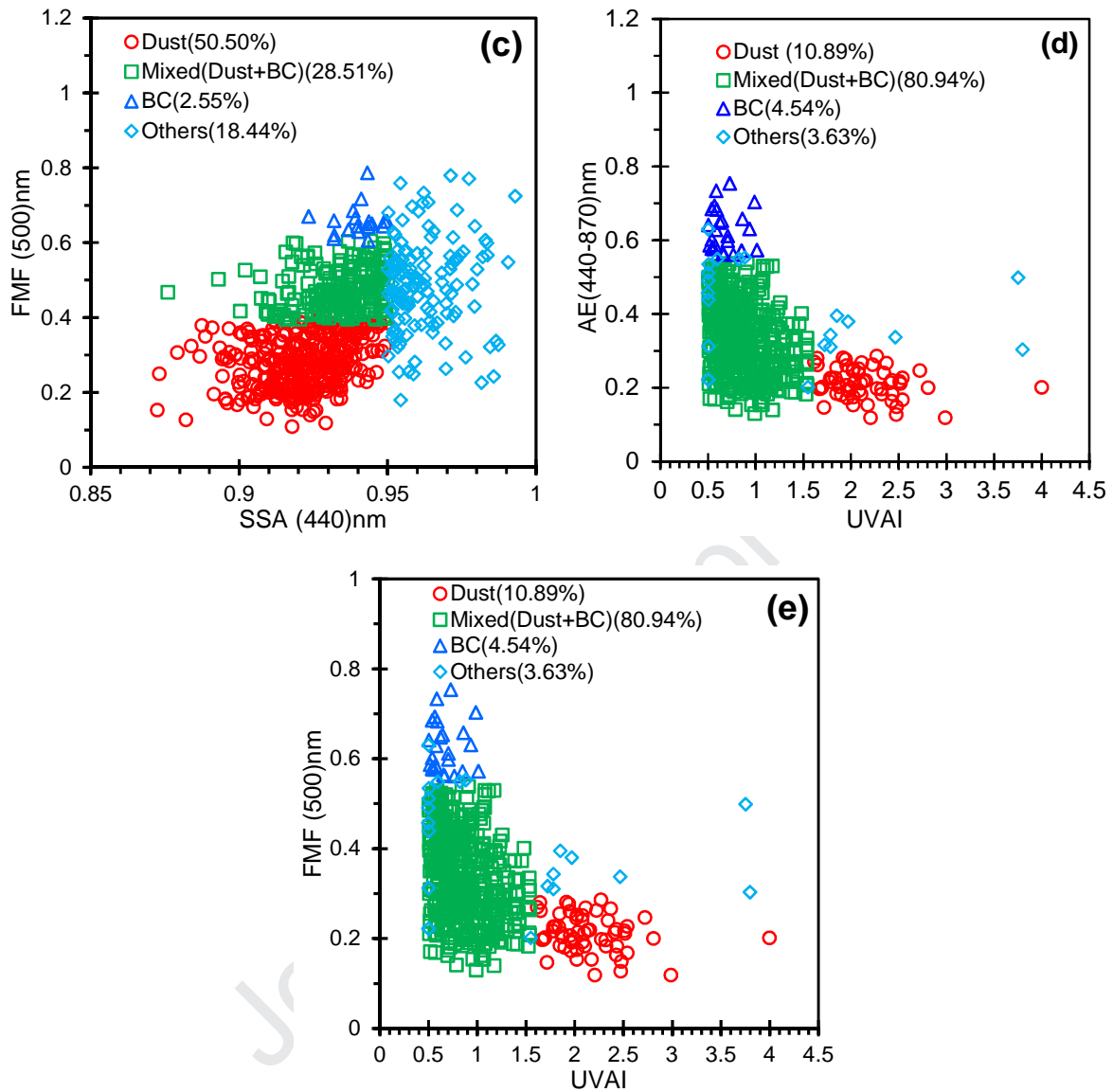
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FMF vs. AE, (b) FMF vs. AAE, (c) FMF vs. SSA, (d) AE vs. UVAI, and (e) FMF vs. UVAI.



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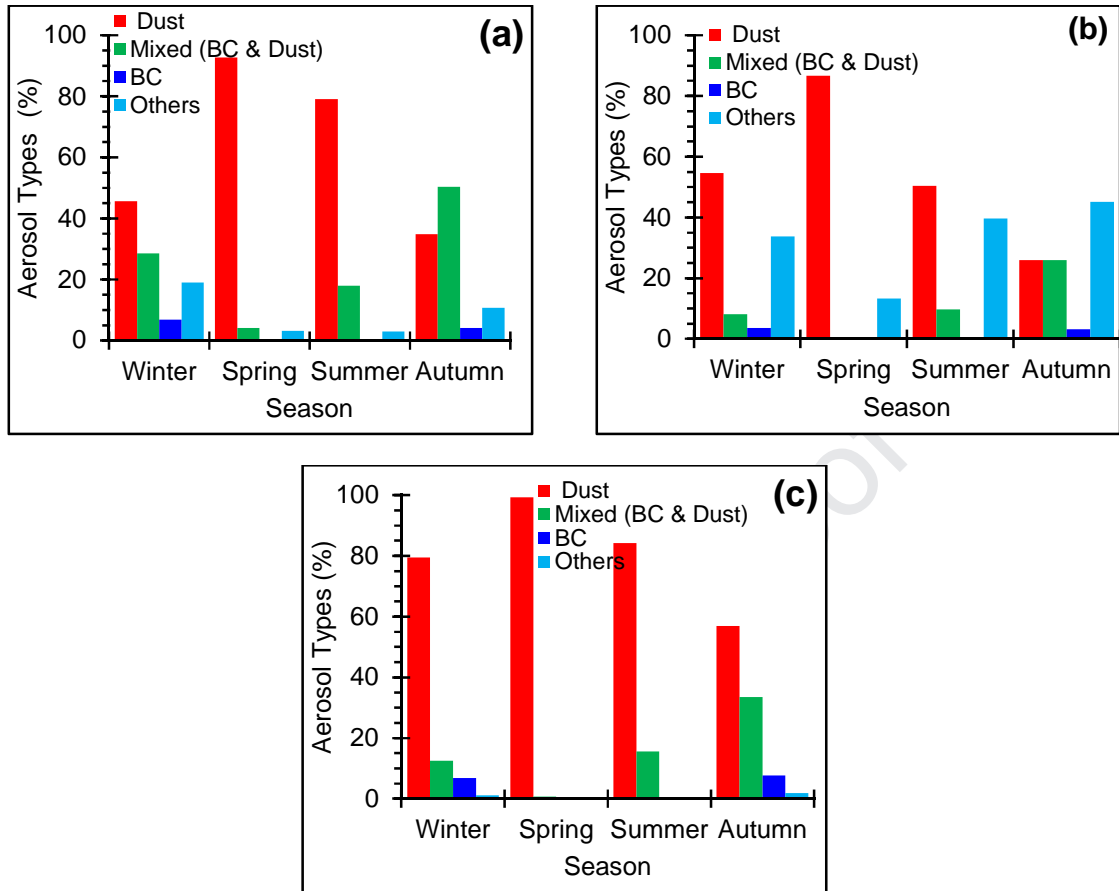
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459 **Figure 8:** Aerosol classification over KAUST Campus site during the period 2012-2016 for (a)

460 FMF vs. AE, (b) FMF vs. AAE, (c) FMF vs. SSA, (d) AE vs. UVAI, and (e) FMF vs. UVAI.

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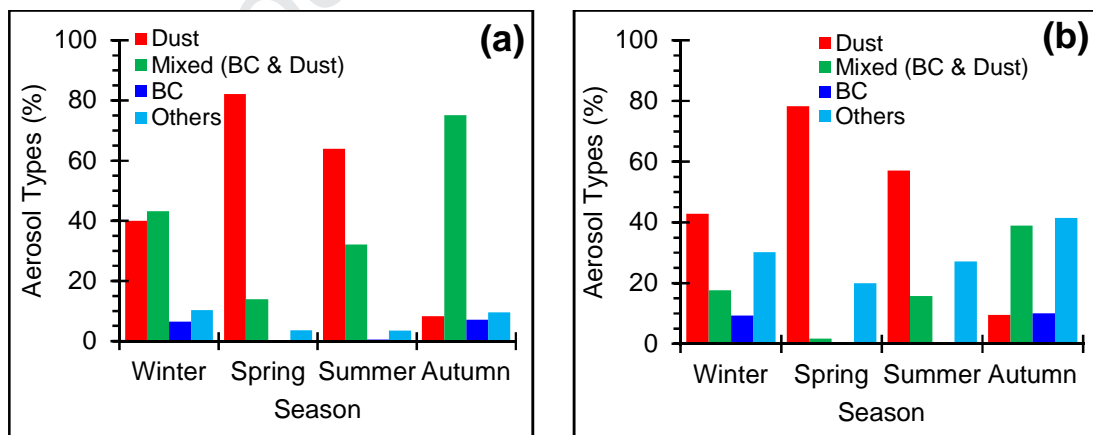
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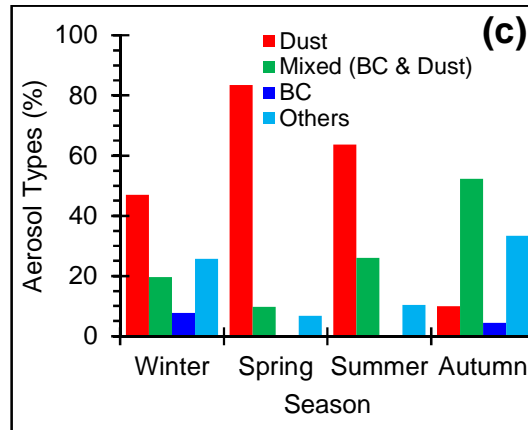
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Figure 9: Seasonal aerosols classification over the Solar Village site during the period 2004-2013 for (a) FMF vs. AE, (b) FMF vs. AAE, (c) FMF vs. SSA.

465





466

467 **Figure 10:** Seasonal aerosols classification over KAUST Campus site, Saudi Arabia during the
 468 period 2012–2016 for (a) FMF vs. AE, (b) FMF vs. AAE, (c) FMF vs. SSA.

469

470 3.4 Validation of Classified Aerosol Types

471 Classified aerosol types were validated against CALIPSO daytime aerosol type profiles.
 472 A similar approach was used in previous studies (Bibi et al., 2016, Bibi et al., 2017; Rupakheti et
 473 al., 2019), as no other data are available for validation. The CALIPSO (daytime) aerosol type
 474 profiles were downloaded for specific dates, according to the availability of AERONET data,
 475 including 24-Jul-2007, 11-Jul-2010, 06-Mar-2011, 29-Mar-2011, and 21-Apr-2008 for the Solar
 476 Village site, and 23-May-2012, 25-Jan-2013, 10-Feb-2013, 5-Mar-2013, and 02-Jun-2013 for the
 477 KAUST Campus site. Results from CALIPSO showed the dominance of Dust aerosol types
 478 reaching up to 5 km from the surface over both sites (Table 5 and Figures S1–S2). Results also
 479 showed the presence of other mixed aerosols (i.e. dust plumes and biomass burning mixed and
 480 forming polluted dust) over the study area (Table 5 and Figure S2–S3). The results showed a
 481 good agreement between FMF vs. AE, FMF vs. AAE, and FMF vs. SSA and the CALIPSO
 482 classified aerosol types over the region. Therefore, this study recommends use of these three
 483 relationship techniques for aerosol classification over Saudi Arabia and other regions with
 484 similar atmospheric and land surface characteristics.

485 **Table 5** Aerosol classifications based on FMF vs. (AE, AAE, and SSA) and CALIPSO.

Date	FMF vs. AE	Types	FMF vs. AAE	Types	FMF vs. SSA	Types	CALIPSO			
Site: Solar Village										
24-Jul-2007	0.24	0.26	Dust	0.24	2.43	Dust	0.24	0.89	Dust	Dust
21-Apr-2008	0.15	0.04	Dust	0.15	3.23	Dust	0.15	0.92	Dust	Dust
11-Jul-2010	0.32	0.46	Dust	0.32	2.70	Dust	0.32	0.93	Dust	Dust

06-Mar-2011	0.25	0.36	Dust	0.25	3.45	Dust	0.25	0.88	Dust	Dust
29-Mar-2011	0.13	0.01	Dust	0.13	2.92	Dust	0.13	0.92	Dust	Dust
Site: KAUST Campus										
23-May-2012	0.29	0.23	Dust	0.29	2.89	Dust	0.29	0.90	Dust	Dust
25-Jan-2013	0.48	0.91	Mixed	0.48	1.41	Mixed	0.48	0.93	Mixed	Mixed
10-Feb-2013	0.37	0.55	Dust	0.37	2.01	Dust	0.37	0.93	Dust	Dust
5-Mar-2013	0.37	0.56	Dust	0.37	2.49	Dust	0.37	0.94	Dust	Dust
02-Jun-2013	0.19	0.10	Dust	0.19	3.15	Dust	0.19	0.90	Dust	Dust

486

487 **4. Conclusion**

488 In this paper, aerosol types over Saudi Arabia were classified using the aerosol property
489 relationships technique and data from OMI (AAOD, UVAI) and AERONET (AAOD, AE, AAE,
490 FMF, SSA). Based on the three relationships FMF vs. AE, FMF vs. AAE, and FMF vs. SSA, the
491 study found dust to be the most common and abundant aerosol type at both annual and seasonal
492 scales, and this was expected, due to the frequent dust storm activity over the study area. Notable
493 temporal variations in aerosol type were observed and attributed to seasonal climatic changes,
494 especially the greater percentage of Dust aerosol types in spring due to depressions passing
495 eastwards over the Sahara Desert, a major dust source. Local dust sources are also more
496 significant during the hot and dry seasons of spring to early summer; spatial variations are also
497 significant, with high AAOD values over the Eastern and Southern provinces, due mainly to
498 local dust sources, and lower AAOD over the Northern Province. Besides Dust, significant
499 amounts of BC and Mixed (Dust and BC) aerosols were observed, though in lesser amounts than
500 Dust, which are attributed to increasing industrial activities (cement, petrochemicals and
501 fertilizers), water desalination plants, infrastructure, and electric energy generation. These release
502 absorbing and fine particles, which often become mixed with dust. Significant underestimation in
503 OMI UVAI and AAOD products was observed, suggesting that significant improvements are
504 required for the OMI OMAERUV algorithm for better estimation of AAOD over bright desert
505 surfaces. Consequently the study found that the relationships FMF vs. UVAI and AE vs. UVAI
506 (with UVAI derived from OMI) relationships misclassified aerosol types over the study area,
507 therefore the relationships FMF vs. AE, FMF vs. AAE, and FMF vs. SSA are recommended for
508 aerosol classification over Saudi Arabia and areas with similar land and atmospheric
509 characteristics. Validation of the classified aerosol types against CALIPSO data showed that the
510 recommended aerosol classification relationships (FMF vs. AE, FMF vs. AAE, and FMF vs. SSA)
511 are robust and effective for aerosol classification over Saudi Arabia. In view of increased

512 knowledge of the harmful health effects of dust-borne synthetic compounds, the aerosol
513 relationships for identifying the specific aerosol types described here, should be of benefit in
514 future air quality control programs, as well as in global studies of climatic forcing due to aerosol
515 in arid regions.

516

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527

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Highlights

- AERONET, OMI, and CALIPSO datasets were used for classifying aerosols
- OMI AAOD shows the dominance of absorbing aerosols with high seasonal variability
- Dust, then mixed black carbon and dust dominated over the study area, Saudi Arabia
- Mixed aerosol types suggest increasing fossil fuel and biogenic emissions
- FMF vs. (AE, AAE, and SSA) are the best techniques for classifying aerosols