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1 **Characterization and source apportionment of black carbon aerosol**
2 **in the Nanjing Jiangbei New Area based on two years of**
3 **measurements from Aethalometer**

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30

31 **Abstract**

32 Black carbon (BC) aerosols were measured using a seven channel Aethalometer at a
33 suburban site of Nanjing, East China from January 2015 to December 2016 to study
34 its temporal variations and quantify the magnitude of BC from fossil fuel (BC_{ff}) and
35 biomass burning (BC_{bb}) sources. The mean BC mass concentration was observed to
36 be 2200 ± 1309 ng/m³ at the sampling site during the entire observation period, with
37 the highest (lowest) concentrations found in the winter (summer and spring). A
38 distinct diurnal variations in BC revealed with two maximum peaks occurred between
39 06:00 and 09:00 local time (LT) and 19:00 and 22:00 LT in all four seasons, was
40 correlated with source emissions, meteorology, and dynamics of the atmospheric
41 boundary layer. A significant seasonality was observed in the absorption Ångström
42 exponent (α) with higher in spring and winter seasons, and lower during the summer.
43 Further, it is evident that the contribution of BC_{ff} (BC_{ff}%) dominated during the
44 observation period with the mean contributions of BC_{ff}% and BC_{bb}% to the BC were
45 found to be ~81% and 19%, respectively. Similar to α , BC_{bb}% also exhibited high
46 fractions in spring and winter suggests an enhanced contribution from biomass
47 burning sources, and a low in summer. The results from both the potential source
48 contribution function (PSCF) and concentration weighted trajectory (CWT) models
49 indicated that the air masses originating from the northern Zhejiang and Anhui
50 provinces were the potential source areas responsible for the high BC concentrations
51 attributed from agricultural waste burning in Nanjing.

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57 **Keywords:** Aethalometer; Black carbon aerosol; Absorption Ångström exponent;
58 Source apportionment; Nanjing.

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61 **1. Introduction**

62 Atmospheric aerosols have an important influence on the radiation balance of the
63 earth-atmosphere system, both at regional and even global scales (Ramanathan and
64 Carmichael, 2008). The carbonaceous aerosol is one of the most important
65 components of aerosol particles, which are mainly composed of organic carbon (OC)
66 and black carbon (BC) (Bisht et al., 2015; Begam et al., 2016; Prasad et al., 2018).
67 Black carbon aerosols are often produced by human activities, resulting from the
68 incomplete combustion of fossil fuels and biomass (Ravi Kiran et al., 2018).
69 Considering the key role that BC plays in atmospheric radiative effects, as well as
70 health effects, BC aerosols have become an important topic in the recent years.
71 However, BC can heat the atmosphere and contribute strikingly to global warming
72 because of its strong absorption of incoming solar radiation from the visible to
73 infrared wavelengths. Some studies have shown that BC may be the second greatest
74 contributor to the greenhouse effect after CO₂ (Babu and Moorthy, 2002; Joshi et al.,
75 2016; Liu et al., 2018). Furthermore, it contributes greatly to indirect radiative forcing
76 by acting as cloud condensation nuclei (Dumka et al., 2010; Joshi et al., 2016). BC
77 can also reside in the atmosphere from 1 week to 10 days in the absence of
78 precipitation (Kumar et al., 2011; Tiwari et al., 2013), thereby facilitating the
79 deterioration of air quality (Liu et al., 2018). Being nearly in the sub-micron size
80 range, BC particles are readily able to penetrate into the human respiratory system and
81 cause adverse effects on health (Babu and Moorthy, 2002).

82 It is well known that a tremendous amount of pollutants have been emitted in
83 East Asia due to rapid industrial development, surging auto possession, and frequent
84 amount of seasonal biomass burning (Pan et al., 2011; Liu et al., 2018; Kolhe et al.,
85 2019). Based on the study of regional and global BC emission inventories,
86 approximately 1/4 of the global annual BC emissions were found to be derived from
87 China (Bond, 2004; Zha et al., 2014), However, the contribution of different source
88 types to BC aerosol remains unclear due to their heterogeneous emission sources and
89 strengths (Pan et al., 2011; Prasad et al., 2018). Since the 1990s, field measurements
90 of BC aerosols in China have been performed in many cities, including those in the

91 Pearl River Delta region (Cheng et al., 2006; Huang et al., 2012; Tan et al., 2016),
92 Yangtze River Delta region (Feng et al., 2014; Zhang et al., 2015; Zhuang et al., 2014),
93 and the Beijing-Tianjin-Hebei region (Ran et al., 2016; Yan et al. 2008). These studies
94 have provided useful information for better comprehending the characteristics of BC
95 aerosols in China. However, most studies relied on the short-term measurements, and
96 few studies have attempted to estimate the contribution of BC sources due to fossil
97 fuel and biomass burning.

98 Nanjing is an economic megacity in the Yangtze River Delta (YRD) region,
99 which is confronted with severe air pollution (An et al., 2014). In this study,
100 measurements of BC aerosol were conducted from January 2015 to December 2016
101 with a seven channel Aethalometer. The data obtained is used to investigate the
102 temporal variations of BC mass concentrations on daily, monthly, and seasonal scales,
103 and identify potential source areas of BC through the Potential Source Contribution
104 Function (PSCF) and Concentration Weighted Trajectory (CWT) statistical models.
105 Further, we also investigated to quantify the BC mass concentration in terms of BC
106 from fossil fuel (BC_{ff}) and biomass burning (BC_{bb}) sources using an Aethalometer
107 model. The results may aid in our understanding of the pollution characteristics of BC
108 and provide useful information with implications for formulating BC emissions
109 inventory over Nanjing in the YRD region, East China.

110 **2. Materials and methods**

111 **2.1. Sampling site**

112 Measurements of BC were performed on the campus of Nanjing University of
113 Information Science and Technology (32°21'N, 118°72'E), a suburban site of Nanjing
114 in the Jiangbei New Area (Fig. 1). The Jiangbei New Area is the only state-level new
115 district in East China located within the Jiangsu Province (Figure S1 of
116 Supplementary Material (SM)). It is located north of the Yangtze River in Nanjing,
117 with a total area of approximately 2451 km², and accounting for 37% of the Nanjing
118 area. It is also an important node of the Yangtze River Economic Belt. The pollution
119 sources in this area are primarily emissions from traffic emissions, urban building and
120 road constructions, and industrial activities (Wang et al., 2016). The Nanjing chemical

121 industry is situated ~3 km northeast of the sampling site; there were also iron and
122 steel plants within 2 km of the site and main roads ~ 500 m near to the sampling site.
123 General meteorological parameters were monitored continuously by the automatic
124 weather station (AWS) installed on the campus, which is approximately 800 m away
125 from the measurement site.

126 **2.2. Instrumentation**

127 Mass concentrations of BC aerosol were recorded with an Aethalometer (Model
128 AE-33, Magee Scientific Company, USA) at 370 nm, 470 nm, 525 nm, 590 nm, 660
129 nm, 880 nm, and 940 nm. The Aethalometer was placed in an air-conditioned room,
130 on the roof of the Meteorology Building located on the campus of Nanjing University
131 of Information Science and Technology, Nanjing where the temperature inside of the
132 room was maintained at 25°C. The instrument aspirates ambient air through an inlet
133 tube mounted vertically, at an elevation of 2 m from the roof level and ~30 m from the
134 ground level. The particles in the incoming air flow were deposited on the quartz filter
135 tape of the Aethalometer and the BC concentration was determined by measuring the
136 change in transmittance through the quartz filter. The instrument continuously
137 operated at a constant flow rate of 5.0 litres/min (LPM), with a measurement
138 frequency of 5 min. A 2.5 µm sharp cut-point of the cyclone was used at the inlet of
139 the sampling tube to avoid dust and other coarse particles entering the instrument.
140 After passing through the inlet, the sampled air was then dried by a drier to control the
141 relative humidity (RH) to ~40%. Further details of the instrument and its operation
142 have been discussed elsewhere (Hansen et al., 1984). The light absorption measured
143 at 880 nm is considered to represent the effect of BC, since BC is the principal
144 absorber of light at that wavelength, while other known aerosol components have
145 negligible absorption (Bodhaine, 1995). An absorption efficiency value of $16.6 \text{ m}^2 \text{ g}^{-1}$
146 (given by manufacturer) was used to determine the BC mass concentration at 880 nm
147 as used by previous authors (Beegum et al., 2009; Kumar et al., 2011; Begam et al.,
148 2016; Liu et al., 2018; Prasad et al., 2018). Periodical and routine calibration, and
149 flow check of the instrument were carried out to maintain the data reliability and
150 precision.

151 The Aethalometer calculated the absorption coefficients by measuring the
 152 attenuation (ATN) of the incident light due to which the particles impinged on the
 153 tetrafluoroethylene (TFE)-coated glass filter, and the BC concentration was
 154 determined using the wavelength-dependent mass absorption cross section (Ran et al.,
 155 2016). Because the Aethalometer relies on filter sampling, several systematic errors
 156 exist in the measurement of BC. It has been observed that after the complete
 157 advancement of filter tape, the deposition of BC particles started impinging on the
 158 filter tape and the loading apparently increases with the real-time (Hansen et al., 1984;
 159 Begam et al., 2016). Consequently, the relationship between BC loading and ATN
 160 change is non-linear, and this results in the underestimation of BC concentrations,
 161 with is referred to as the “loading effect”(Virkkula et al., 2007; Beegum et al., 2009;
 162 Singh et al., 2015; Begam et al., 2016; Vaishya et al., 2017; Ravi Kiran et al., 2018).
 163 Compensation for the loading effect requires information on the temporal and spatial
 164 variability of aerosol properties. Drinovec et al. (2015) have suggested a real-time
 165 loading effect compensation algorithm that uses a dual-point approach in which the
 166 incoming flow is divided into two streams with different rates and passing through
 167 two sample points. Therefore, two synchronous measurements of ATN may be
 168 obtained, both of which are affected by different loading effects. These techniques
 169 have been incorporated into the new Aethalometer, which is used in this work.

170 **2.3. Methodology**

171 *2.3.1. Calculation of absorption Ångström exponent*

172 The calculation of the absorption Ångström exponent involves a power law
 173 dependence of σ_{abs} on λ . This relationship is commonly expressed as:

$$174 \sigma_{\text{abs}}(\lambda) = K\lambda^{-\alpha}, \quad (1)$$

175 where σ_{abs} is the spectrally dependent absorption coefficient, λ is the light wavelength,
 176 and K is a constant. The absorption Ångström exponent is represented by α , which
 177 may then be calculated from measured σ_{abs} as a function of different wavelengths
 following the formula:

$$\alpha_{\lambda_1-\lambda_2} = -\frac{\ln(\sigma_{abs,\lambda_1} / \sigma_{abs,\lambda_2})}{\ln(\lambda_1 / \lambda_2)}. \quad (1)$$

178 2.3.2. Aethalometer model

179 The Aethalometer model is based on the work of Sandradewi et al. (2008) and
 180 aims to quantify the contribution of fossil fuel (BC_{ff}) and biomass burning (BC_{bb}) to
 181 the total BC mass concentration. It assumes that the aerosol optical absorption
 182 coefficient can be viewed as the sum of fossil fuel and biomass combustion fractions.
 183 The two sources exhibit specific Ångström exponents with respect to aerosol
 184 absorption because BC from fossil fuels has a weak wavelength-dependence, while
 185 BC originating from biomass burning exhibits enhanced absorption towards shorter
 186 wavelengths, and thus the sources of these two components can be modeled
 187 (Drinovec et al., 2015; Herich et al., 2011). Following a study by Kirchstetter et al.
 188 (2004), the absorption Ångström exponents of fossil fuels ($\alpha_{ff} = 1$) and biomass
 189 burning ($\alpha_{bb} = 2$) were used in this study. The quantification of BC from fossil fuels
 190 and biomass burning may then be calculated by solving equations (3–7).

$$\sigma_{abs}(\lambda) = \sigma_{abs}(\lambda)_{ff} + \sigma_{abs}(\lambda)_{bb} \quad (3)$$

191 In equation (3), $\sigma_{abs}(\lambda)_{ff}$ and $\sigma_{abs}(\lambda)_{bb}$ refer to the aerosol absorption coefficient of BC
 192 from the combustion of fossil fuels and biomass burning at a wavelength λ ,
 193 respectively. Furthermore,

$$\frac{\sigma_{abs}(370)_{ff}}{\sigma_{abs}(880)_{ff}} = \left(\frac{370}{880}\right)^{-\alpha_{ff}} \quad (4)$$

$$\frac{\sigma_{abs}(370)_{bb}}{\sigma_{abs}(880)_{bb}} = \left(\frac{370}{880}\right)^{-\alpha_{bb}} \quad (5)$$

$$BC_{ff} = BC \frac{\sigma_{abs}(880)_{ff}}{\sigma_{abs}(880)} \quad (6)$$

$$BC = BC_{ff} + BC_{bb} \quad (7)$$

194 2.3.3. HYSPLIT model

195 Air mass backward trajectories arriving at Nanjing were calculated using the
 196 United States National Oceanic and Atmospheric Administration (NOAA) Air

197 Resource Lab (ARL) Hybrid Single Particle Lagrangian Integrated Trajectory version
 198 4.0 (HYSPLIT-4) model and the meteorological data derived from the Global Data
 199 Assimilation System (GDAS). Four-day backward air trajectories were computed at
 200 500 m above ground level and 1-h time intervals. As this arrival altitude was suitable
 201 for the identification of long-range transport impacts, it ensured that the computation
 202 of the back-trajectories began within the heights of atmospheric boundary layer
 203 (Karaca et al., 2009).

204 2.3.4. PSCF and CWT models

205 The potential source contribution function (PSCF) model is a method of
 206 identifying source areas based on air trajectory analysis. It can localize suspected air
 207 pollutant source areas efficiently by counting air back-trajectory segment endpoints
 208 that terminate within each cell (Cheng et al., 2006; Dimitriou and Kassomenos, 2016).
 209 The PSCF value of a single grid cell ($0.5^\circ \times 0.5^\circ$) is given by the following equation:

$$210 \quad \text{PSCF}_{ij} = \frac{m_{ij}}{n_{ij}} \quad (8)$$

211 where n_{ij} is the total number of trajectory points that fall in the ij th cell and m_{ij} is the
 212 number of trajectory points included in the ij th cell, corresponding to trajectories with
 213 pollutant concentrations higher than a threshold value (Wang et al., 2009; Kong et al.,
 214 2013).

215 To overcome limitations with the PSCF, we further performed the concentration
 216 weighted trajectory (CWT) method to identify the source strengths of BC in Nanjing.
 217 In the CWT method, we used a plugin called TrajStat: GIS based software to elaborate
 218 the analysis which can be found elsewhere (Wang et al., 2009). In the CWT method,
 219 each grid cell is assigned a weighted concentration by averaging the sample
 220 concentrations that have associated trajectories crossing that grid cell represented as
 221 follows:

$$222 \quad \text{CWT} = \frac{1}{\sum_{i=1}^M T_{ijl}} \sum_{i=1}^M C_i T_{ijl} \quad (9)$$

223 where CWT is the weighted concentration of BC in trajectory in the ij th cell, l and M

224 denote the trajectory indices and total number of trajectories, respectively. C_l is the
225 concentration of BC observed at receptor site on arrival of trajectory l and T_{ijl} is the
226 residential time of trajectory l in the ij th cell. Higher the value of CWT represents the
227 magnitude of source strength associated with BC concentration observed at receptor
228 site.

229 **3. Results and discussion**

230 **3.1. General meteorology**

231 Figure 2 shows the monthly variations of temperature (T), RH, precipitation (PP),
232 and wind speed (WS) measured from the AWS during the years between 2015 and
233 2016 in the northern suburb of Nanjing. The monthly mean RH found higher than
234 60% during the observation period, and the highest value of 87% was occurred in
235 October 2016. Whereas, total annual precipitation was measured of ~1350 mm, with
236 an occurrence of maximum precipitation (~900 mm) was observed during summer
237 (JJA months) and autumn (SON) seasons, which accounted for more than 70% of the
238 total precipitation. An annual precipitation of 1687 mm noticed in 2016 which is at an
239 increase of 65% compared to 2015. However, the ambient air temperature presented a
240 distinct seasonal pattern, with the highest values (24–29°C) in summer and the lowest
241 values (3–7°C) in winter (DJF months). A weak seasonal variation in WS was
242 depicted in Nanjing, but the values were higher (1.8–2.4 m/s) during spring (MAM)
243 and lower in winter (1.5–2.0 m/s).

244 **3.2. Temporal variations in BC concentrations**

245 The daily and monthly variations in the mass concentrations of BC from 2015 to
246 2016 are shown in Figure 3a. The daily averaged BC concentration measured at
247 Nanjing varied from 150 ng/m³ to 10423 ng/m³, with a mean value of 2200 ± 1309
248 ng/m³ during the entire study period. In general, higher BC concentrations were
249 mainly observed during December and January, with the monthly mean values
250 fluctuating between 2796 ± 1646 ng/m³ and 3470 ± 2450 ng/m³. The concentrations
251 were relatively lower during other months, particularly in April and May, with the
252 mean values were ranged between 933 ± 536 ng/m³ and 2778 ± 1486 ng/m³. The
253 maximum monthly mean concentrations of BC occurred in January of 2015, likely

254 caused by the occurrence of several severe pollution days during January 3–5, January
255 10, and January 24, when the daily averaged BC concentrations exceeded 5000 ng/m^3 .
256 The inter-annual variations in BC concentrations were observed in Nanjing. In 2015,
257 the annual mean concentration of BC was $2654 \pm 1364 \text{ ng/m}^3$, while a lower mean
258 value of $1866 \pm 1160 \text{ ng/m}^3$ was found in 2016. This decrease was mainly due to the
259 occurrence of heavy precipitation in 2016 (Fig. 2), which indicates that the
260 meteorological conditions played a crucial role in the variation of BC.

261 BC concentrations in Nanjing exhibited a substantial seasonal heterogeneity during
262 the measurement period (Fig. 3b). The mean concentrations of BC by season in 2015
263 were found to be: spring = $2146 \pm 938 \text{ ng/m}^3$, summer = $2427 \pm 1217 \text{ ng/m}^3$, autumn =
264 $2647 \pm 1123 \text{ ng/m}^3$, and winter = $3454 \pm 1790 \text{ ng/m}^3$. Mean while, BC concentrations
265 for each season in 2016 were noticed as: spring = $1640 \pm 1140 \text{ ng/m}^3$, summer = 1461
266 $\pm 744 \text{ ng/m}^3$, autumn = $1586 \pm 773 \text{ ng/m}^3$, and winter = $2852 \pm 1297 \text{ ng/m}^3$. The
267 concentrations of BC were remained lower during spring and summer seasons likely
268 related to the strong convective activity and heavy summer precipitation phenomena
269 over the region (Kang et al., 2016). High BC concentrations in winter were mainly
270 attributable to the severe emissions of anthropogenic pollutants produced more from
271 the heating sources and the stability of the atmospheric boundary layer (ABL) due to
272 lower temperature, which favored the accumulation of pollutants in the surface (Cao
273 et al., 2009). Additionally, as the winters were dominated by westerly winds, air
274 masses from inland areas may be expected to have been anthropogenically perturbed
275 during their migration before reaching the sampling site, resulting in an increase in the
276 BC concentration (Feng et al., 2014; Prasad et al., 2018). Similar results have also
277 been observed for Shanghai (Zha et al., 2014), Hefei (Zhang et al., 2015), and Xi'an
278 (Cao et al., 2009) over China; and Dhanbad (Singh et al., 2015), Kadapa (Begam et al.,
279 2016), Gadanki (Ravi Kiran et al., 2018), and Pune (Kolhe et al., 2019) over India.

280 The variations of atmospheric aerosols are largely influenced by regional
281 environmental conditions, including rural and urban settings. For example, the
282 characteristics of aerosols may vary greatly due to different sources and atmospheric
283 conditions. To facilitate a more holistic understanding of the distribution

284 characteristics of BC in different areas, Table 1 lists the observed BC concentrations
285 from other sites. Overall, BC concentrations at urban sites were higher than those at
286 rural, suburban, and remote sites. The BC concentration in the northern suburb of
287 Nanjing was much lower than those at urban sites, such as Shenzhen (Huang et al.,
288 2012) and Shenyang (Wang et al., 2011) in China, and at Delhi in India (Tiwari et al.,
289 2013). However, it was comparable with urban Shanghai (Zha et al., 2014), suburban
290 Beijing (Yan et al., 2008), and urban Pune in India (Kolhe et al., 2019). Furthermore,
291 it was higher than the measurements over the outskirts of Lanzhou in China (Zhang et
292 al., 2011), urban El Paso in the USA (Raysoni et al., 2011), urban Pune (Kolhe et al.,
293 2019), and rural Évora in Portugal (Pereira et al., 2012), and lower than urban
294 Granada (Valenzuela et al., 2017).

295 **3.3. Diurnal variations in BC concentration**

296 Figure 4 shows the diurnal variations of hourly averaged BC measured in the
297 northern suburban area of Nanjing. The diurnal pattern of BC concentration exhibited
298 a bimodal distribution with two peaks evident clearly in all the four seasons, which is
299 similar to that observed in all sites listed in Table 1. The BC mass concentration was
300 found maximum in the morning from 06:00 to 09:00 local time (LT). Then the mean
301 value of BC began to decrease after 9:00 LT and reached its lowest values in the
302 afternoon hours between 13:00 and 15:00 LT. The lowest mean BC concentration was
303 approximately $3/5^{\text{th}}$ of the value observed in the morning. After 15:00 LT, the mean
304 concentration of BC began to rise in the evening and remained elevated during the
305 hours 19:00–22:00 LT, when it then decreased again until 04:00 LT of the following
306 morning.

307 The highest BC concentrations were occurred in the morning and late evening hours
308 mainly attributable to the severe emissions of pollutants from transportation sources
309 during the morning and evening rush hours of traffic (Kumar et al., 2011; Zhang et al.,
310 2015; Begam et al., 2016; Prasad et al., 2018). Meanwhile, the concentrations are also
311 dependent upon the dynamics of the ABL height, which is usually stable during the
312 early morning and late nighttime (Stull, 1998). The low value of BC in the afternoon
313 might be related to strong atmospheric mixing and dilution of pollutants due to the

314 elevated ABL height, which increases with enhanced solar radiation (Kompalli et al.,
315 2014). After sunset, BC concentrations began to rise and showed another moderate
316 peak at midnight, which is due to vehicular exhausts and anthropogenic emissions.
317 Besides, the decrease in the surface temperature can cause temperature inversion, and
318 hence the formation of a stable nocturnal ABL in which pollutants are bound in the
319 shallow ABL (Aruna et al., 2013). It is interesting to note that the peak and valley of
320 BC concentration during winter occurred at 09:00 and 15:00 LT, respectively, which is
321 2–3 h later than during the other three seasons. This could be explained by the delay
322 of local human activities because of the cold weather (Cao et al., 2009).

323 **3.4. Source apportionment of BC**

324 The absorption Ångström exponent (α) has typically been used to describe the
325 spectral dependence of light absorption for aerosols derived from different sources
326 and with different absorption characteristics to study and understand its spectral
327 dependence. For BC, α was less than 1.0 when aerosols produced from the
328 combustion of fossil fuels; while aerosols emitted from biomass burning, it showed
329 stronger spectral dependence in absorption characteristics with high values (~ 1.8) and
330 those mixed with dust exhibited much more higher α (>2). This clearly shows that the
331 values of α derived in this study are consistent and close to that found in the works of
332 Kirchstetter et al. (2004), hence followed the same.

333 The monthly mean variations in α over Nanjing is illustrated in Figure 5a. It is
334 found that α showed greater than 1.0 during the entire observation period, without any
335 obvious seasonal change. The monthly mean of α in 2015 exhibited a wide range of
336 fluctuations, with a maximum value of 1.45 occurred in January and minimum (0.99)
337 in July. The variation in monthly mean α in 2016 was mild, with a range of 1.14–1.30.
338 The annual mean α in 2015 and 2016 were found 1.26 and 1.18, respectively, which
339 indicates that BC aerosols over Nanjing mainly originated from fossil fuel combustion.
340 From the seasonal frequency distribution of α shown in Fig. 5a, it can be seen that the
341 spring and winter seasons were characterized by high α levels with a median value of
342 1.25, while the median value decreased to 1.19 during autumn. However, the

343 percentage of high values ($\alpha > 1.6$) during autumn accounted for 11% of the total,
344 signifies the frequent occurrence of substantial biomass burning events to clear the
345 agricultural waste from the farmlands in Nanjing. The α level in summer was slightly
346 lower than during other seasons, with a median value of 1.1. However, α ranged from
347 0.8 to 1.8, implies that the effect of biomass burning on BC concentrations during
348 summer and autumn seasons cannot be ignored relative to the BC sources from fossil
349 fuel combustion which is at a greater extent as observed in this study.

350 In order to have more intuitively comprehend for the relative contributions of
351 fossil fuel (ff) and biomass burning (bb) to BC concentration in different seasons, the
352 monthly mean values of BC_{ff} , BC_{bb} , and their percentage contributions were
353 calculated using the Aethalometer model (Sandradewi et al., 2008). It is evident that
354 the mean contributions of BC_{ff} ($BC_{ff}\%$) and BC_{bb} ($BC_{bb}\%$) to the total BC were found
355 as ~81% and ~19%, respectively (Figure 5b). The maximum contribution from the
356 $BC_{ff}\%$ occurred in July 2015 (85%) and the minimum during January 2015 (73%).
357 Conversely, the contribution of BC_{bb} reached its maximum in January 2015 (27%)
358 and minimum in July 2015 (15%). However, BC_{ff} and BC_{bb} exhibited similar trends,
359 with higher values during winter. Although, the values of BC_{bb} were similar (e.g., in
360 December 2016, June 2015, and October 2015) and, the $BC_{bb}\%$ in December 2016
361 was much lower than in the other two months due to the high BC concentration in
362 winter.

363 Similar to the seasonal pattern of α , the frequency distribution of $BC_{bb}\%$ also
364 exhibited noticeable seasonality (Figure. 5c). The overall contribution of $BC_{bb}\%$ was
365 comparatively lower during summer and the values mostly ranged between 10% and
366 17%, indicates that fewer biomass burning events occurred at the observation site or
367 that there was less contribution of such sources produced or transported from other
368 areas. However, the $BC_{bb}\%$ was centralized around 20% during spring and winter
369 seasons, suggesting that the contribution of biomass burning increased except, for the
370 existing fossil fuel emissions.

371 **3.5. Analysis of BC transport from the CWT model**

372 Figure 7 presents the results obtained from the calculation of CWT model using

373 the HYSPLIT trajectories and GDAS meteorological achieved datasets for BC
374 concentration in different seasons of Nanjing during the entire study period. The
375 seasonal variation of airflow trajectory was obvious, and the potential source areas
376 and their contributions were distinct in different seasons. Despite a large number of
377 air mass trajectories originated from the East China sea due to the transport of ship
378 emissions during spring, summer and fall seasons, their contribution to BC was small
379 ($< 2500 \text{ ng/m}^3$). However, the larger value of CWT with more contribution to BC
380 ($2500\text{-}5500 \text{ ng/m}^3$) in spring mainly distributed in the central parts of Anhui and
381 western Zhejiang Provinces. In addition, the airflow from northern Fujian to eastern
382 Jiangxi also contributed about 5000 ng/m^3 to BC. The larger value of CWT in summer
383 occurred in the junction area between western Zhejiang and southern Anhui and
384 contributed between 2500 and 3500 ng/m^3 to BC mass concentration. The BC
385 concentration was lower in summer because of the greater influence of ocean air
386 masses resulting in a larger amount of precipitation with high wind speed, which is
387 also conducive to the diffusion of pollutants. In fall, the high value of CWT
388 distributed from northern Jiangxi to southern Anhui Provinces, and northeastern
389 Henan to western Shandong Provinces, which contributed more than 4000 ng/m^3
390 to the BC concentration. This is possibly attributed to the anthropogenic activities such
391 as biomass combustion that occurred to clear the agricultural waste burning after
392 harvesting in these areas during the fall season. The high CWT values in winter were
393 mainly distributed in provinces of Fujian, Zhejiang, Jiangxi, Hubei and most areas of
394 Anhui, contributing about $4500\text{-}6500 \text{ ng/m}^3$ to the BC mass concentration in Nanjing.
395 The high CWT values with a large amount of BC ($3000\text{-}4500 \text{ ng/m}^3$) were also
396 distributed in parts of Henan and Shandong provinces possibly related to the biomass
397 burning and heating in the north of China.

398 **3.6. BC pollution events – A case study**

399 To further understand the impacts of different sources on local BC concentrations,
400 two BC pollution events were analyzed in this study. BC pollution events were
401 defined as the periods when the daily averaged BC concentration exceeded the
402 seasonal geometric mean concentration plus 1 time the standard deviation of the mean

403 (Cao et al., 2009). The selected BC events occurred during October 13–17 (event I)
404 and December 20–23 of 2015(event II). Both the temperature and RH showed
405 obvious diurnal variations during the event I (Fig. 6a), wherein the temperature
406 reached its minimum before daybreak and its maximum in the afternoon, which was
407 opposite to the pattern observed for RH. The time series of BC increased and
408 decreased by a wide margin, with an average value of ~ 4800 ng/m³ during event I.
409 The hourly BC concentration sharply increased from 23:00 LT on October 13, and
410 reached its maximum of 12063 ng/m³ at 01:00 LT on October 14, before decreasing to
411 5530 ng/m³ at 03:00 LT. Thereafter, the BC concentration continued to rise, reaching
412 another maximum at 08:00 LT on the same day. Similar pattern was also observed
413 from 23:00 LT on October 14 to 05:00 LT on the 15th October and from 21:00 LT on
414 October 15 to 06:00 LT on the 16th October. In the event I, the contribution of BC_{bb}
415 was nearly agreement with the variation of BC, with an average of 26%. This value
416 was apparently higher than the mean BC_{bb}% of 19% during the observation period,
417 which illustrates that biomass combustion made an important contribution to the high
418 concentration of BC.

419 During event II (Fig. 6b), there was no evident regularity in T or RH. Two weak
420 rainfall events occurred from 04:00–20:00 LT on December 20 and from 03:00–19:00
421 LT on the 22nd December and the cumulative precipitation values were 0.4 mm and
422 0.3 mm, respectively. Accordingly, the RH remained high for the duration of the event,
423 with a mean value of 84%. Weak precipitation does not have an effective scavenging
424 effect on pollutants, and the high RH in the near-surface layer was conducive to the
425 accumulation of water vapor and pollutants. The mean concentration of BC was
426 ~ 5230 ng/m³ and was relatively stable during event II. An abrupt increase could be
427 seen between 05:00 and 11:00 LT on December 21, wherein the highest BC
428 concentration occurred at 09:00 LT on the 21st December, with a value of 12050
429 ng/m³. The BC_{bb}% in event II was lower than in the event I. The averaged values of
430 BC_{bb}% and BC_{ff}% were found to be 16% and 84%, respectively, which indicates that
431 BC concentration during event II was mainly derived from the combustion of fossil
432 fuels. It is further evident from the PSCF results (Figure S2 of SM) conducted during

433 both BC pollution events revealed that the potential sources originating from the
434 northern Zhejiang, Jiangsu, and Anhui Provinces contributing more to the BC
435 concentration observed at Nanjing.

436 **4. Conclusions**

437 Two years (2015-2016) of continuous and real-time data of BC mass
438 concentrations were measured with a seven channel Aethalometer (AE-33) at a
439 suburban site over Nanjing in the YRD region, East China. Along with the
440 meteorology datasets, the observed BC concentrations have been used to analyze its
441 temporal evolution, identify potential emission source areas and strengths (PSCF and
442 CWT models), and also quantify their emission strengths of BC through
443 apportionment model. During the entire study period, the mean BC concentration was
444 found of 2200 ± 1309 ng/m³ and varied within the range 150–10423 ng/m³. The
445 monthly and seasonal variations of BC concentrations revealed high values during the
446 winter, moderate in the autumn, and low in the spring and summer seasons. The BC
447 mass concentration was found to exhibit a well-defined diurnal variation with two
448 peaks between 06:00 and 09:00 LT and 19:00 and 22:00 LT in all the four seasons.
449 This diurnal pattern is strongly associated with the dynamical behavior of the ABL
450 height, meteorological conditions, and local anthropogenic emissions. The mean
451 absorption Ångström exponent (α) during the entire study period was found to be
452 ~ 1.21 , with higher values being observed in spring and winter, and lower values found
453 in summer. The contribution of BC_{ff} (BC_{ff} %) was dominated during the entire
454 observation period, with the mean contributions of BC_{ff} % and BC_{bb} % to the total BC
455 were approximately 81% and 19%, respectively. Similar to α , BC_{bb} % also exhibited
456 higher levels in spring and winter, and a lower level in summer, suggests an enhanced
457 contribution from biomass burning sources except, for the existing fossil fuel
458 emissions during spring and winter seasons. The results of the CWT model obtained
459 for all seasons revealed that the northern Zhejiang and Anhui provinces were the
460 important source regions for the transport of BC resulted in the high BC
461 concentrations. This is consistent with the PSCF method conducted during the BC
462 pollution events observed at Nanjing.

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472 **Conflict of interests**

473 The authors declare no conflict of interest.

474

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Table 1. Comparison of black carbon (BC) mass concentrations at different sampling sites.

location	Environment type	Study period	BC concentration (ng/m ³)	Citation
Nanjing (32°21'N,118°72'E)	Suburban	2015-2016	2200 ± 1309	Present study
Nanjing (32.05°N,118.78°E)	Urban	2012	4157 ± 2626	Zhuang et al. (2014)
Shanghai (31.18°N,121.30°E)	Urban	2011-2012	2330	Zha et al. (2014)
Hefei	Urban	2012-2013	3500 ± 2500	Zhang et al. (2015)
Shenzhen (22.60°N,113.97°E)	Urban	2009	6000	Huang et al. (2012)
Xianghe (39.80°N,116.96°E)	Rural	2013-2014	5390 ± 4440	Ran et al. (2016)
Beijing (40.39°N,117.07°E)	Suburban	2003-2005	2120 ± 1620	Yan et al. (2008)
Shenyang (41.77°N, 123.50°E)	Urban	2008-2009	6140	Wang et al. (2011a)
Lanzhou (35.57°N, 104.08°E)	Remote	2007-2009	1568	Zhang et al. (2011)
Delhi (28.35°N, 77.12°E)	Urban	2011	6700 ± 5700	Tiwari et al. (2013)
Kadapa (14.47°N, 78.82 °E)	Rural	2011-2012	2200 ± 780	Begam et al. (2016)
Dhanbad (23.47°N, 86.30 °E)	Urban	2012	6300 ± 2700	Singh et al. (2015)
Pune (18.53°N, 73.8°E)	Urban	2015-2016	1924 ± 750	Kolhe et al. (2019)
Gadanki	Rural	2008-2017	2200	Ravi Kiran et al. (2018)
Vijayawada	Suburban	2016	3440±2070	Prasad et al. (2018)
Granada (37.16°N, 3.58 °W)	Urban	2005-2008	3000 ± 1500	Lyamani et al. (2011)
El Paso	Urban	2008	230 - 2480	Raysoni et al. (2011)
Évora (38.57°N, 7.91 °W)	Rural	2007-2009	1300 ± 1200	Pereira et al. (2012)
Granada (37.168°N, 3.608°W)	Urban	2005-2012	2900 ± 900	Valenzuela et al. (2017)

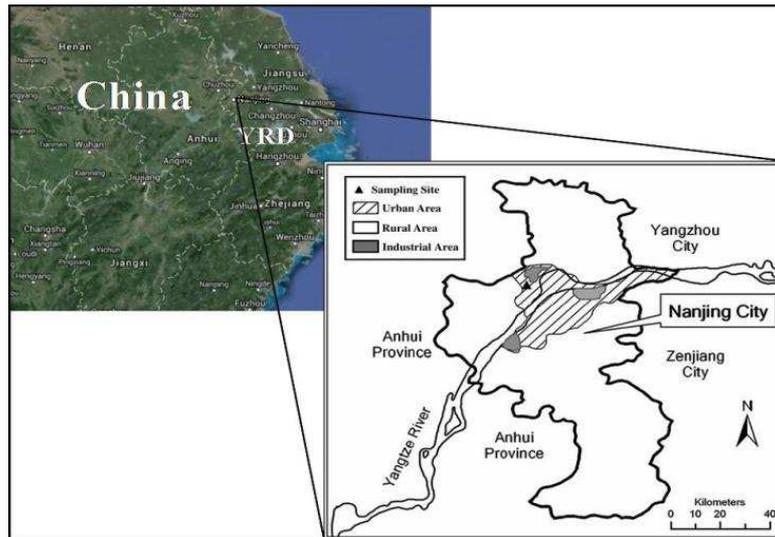


Figure 1. Geographical position of Nanjing city in the YRD region over Southeast China.

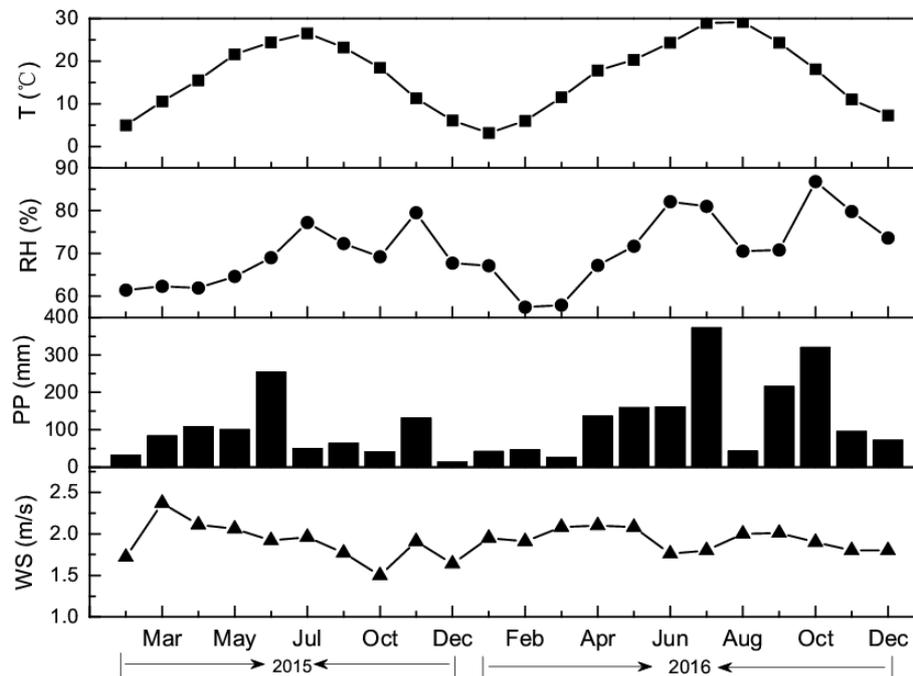


Figure 2. Monthly variations in temperature (T), relative humidity (RH), precipitation (PP) and wind speed (WS) during 2015-2016.

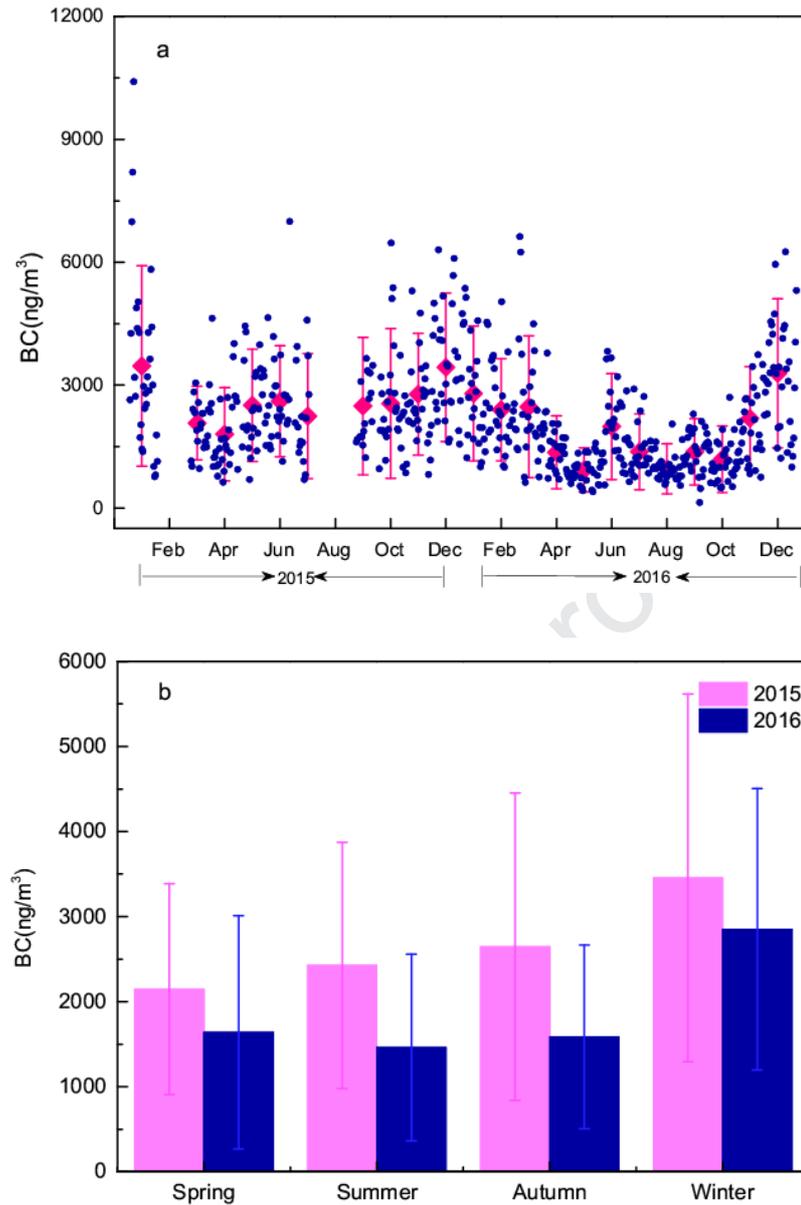


Figure 3. Daily (blue) and monthly (pink) variations of BC mass concentrations (a) and seasonal variations of BC mass concentrations (b) observed in Nanjing from January of 2015 to December of 2016.

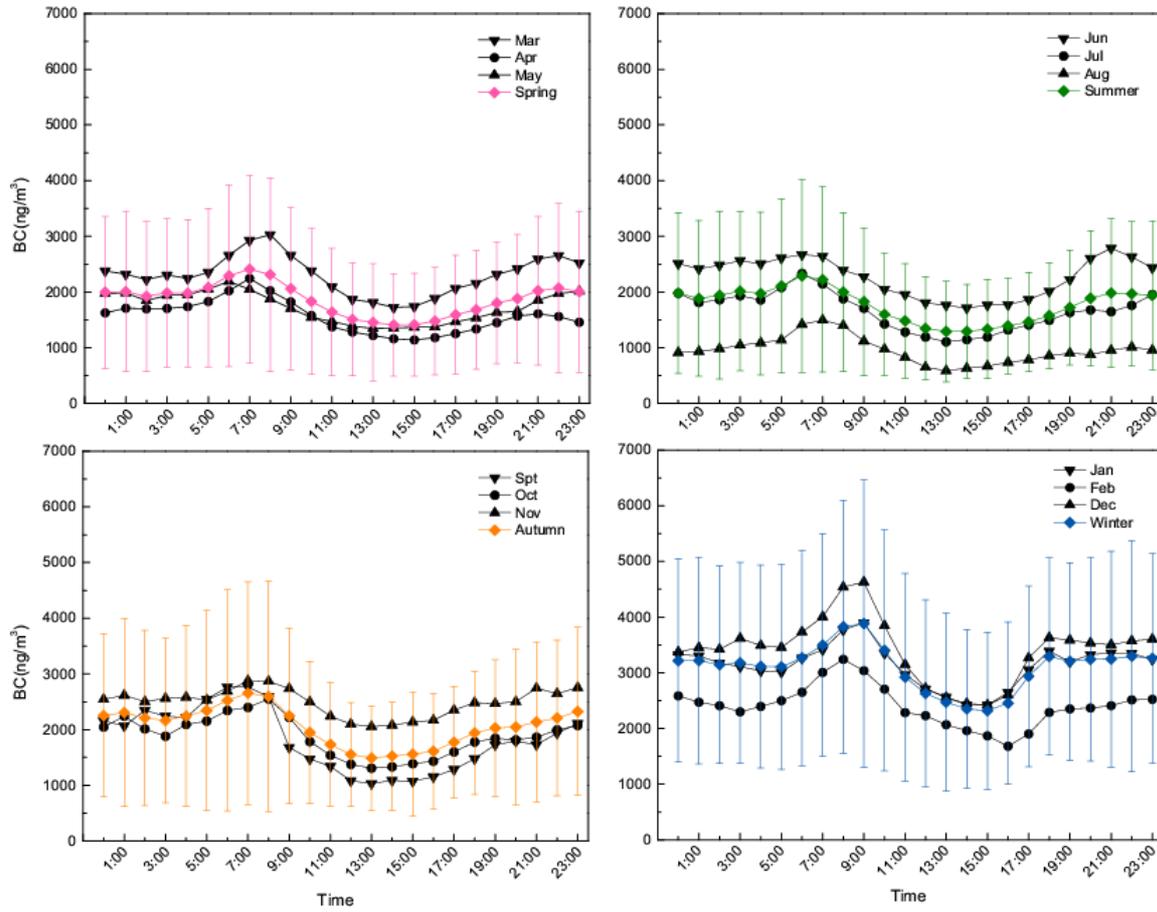


Figure 4. Diurnal variations of BC concentrations for different months and seasons during the entire study period. The solid lines corresponding to the dots denote the standard deviation.

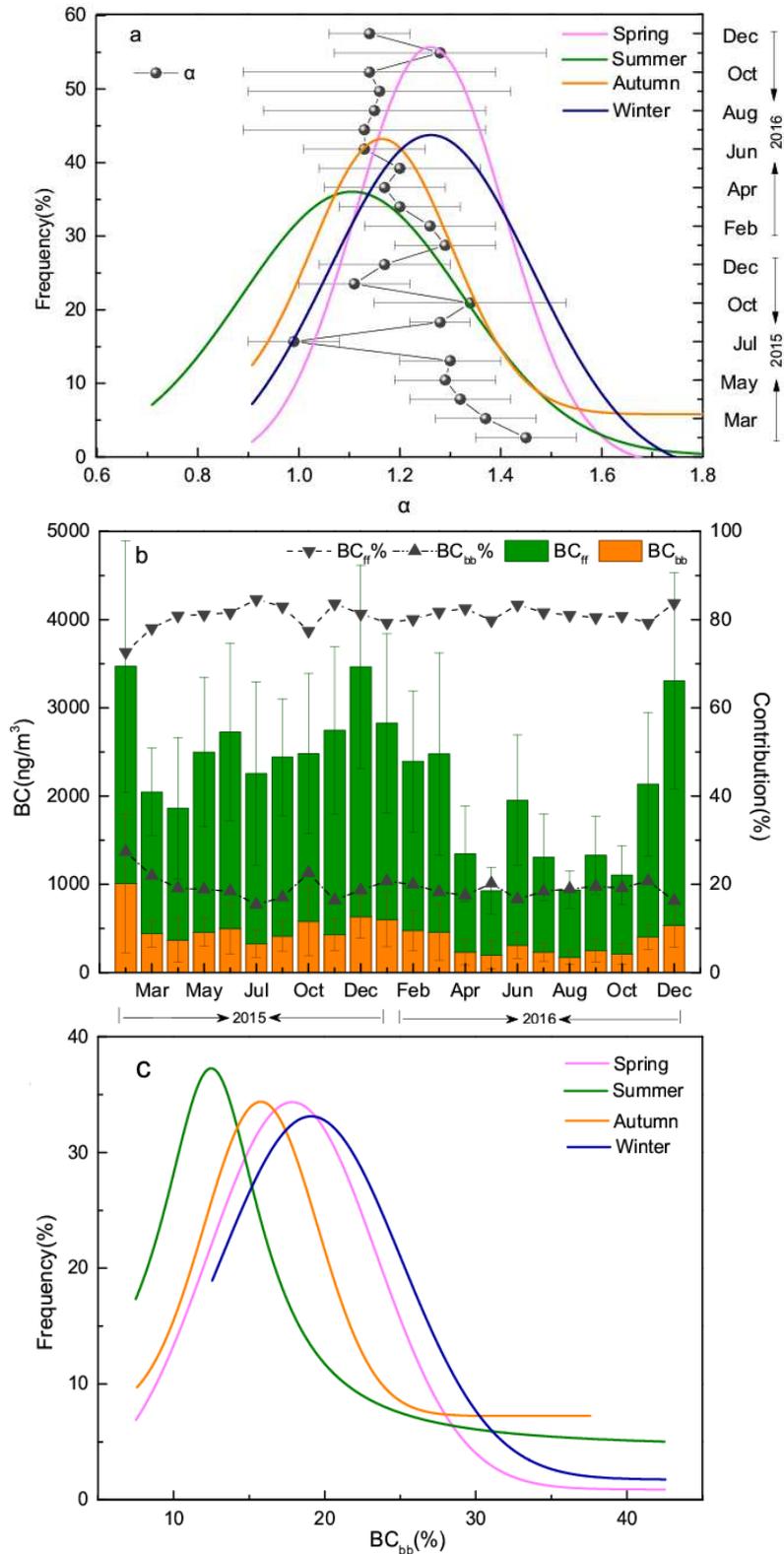


Figure 5. Monthly variations in the absorption Ångström exponent (α) and its frequency distributions for different seasons (a); monthly concentrations of BC from fossil fuels (BC_{ff}) and biomass burning (BC_{bb}) and their contributions in total BC (b); and seasonal frequency distribution of $\text{BC}_{\text{bb}}\%$ (c) observed over Nanjing.

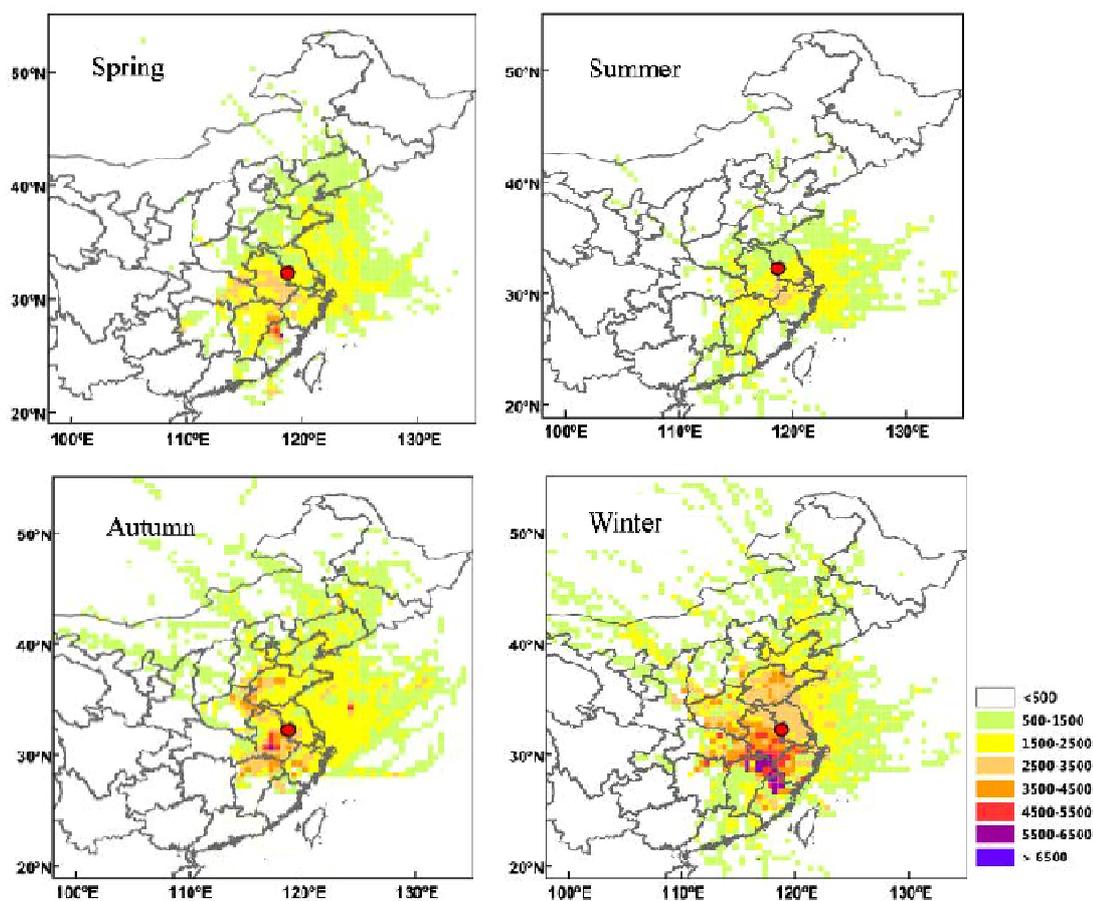


Figure 6. CWT model results obtained for different seasons presenting the long-range transport of BC concentration. The location of the measurement site is shown with the solid red circle.

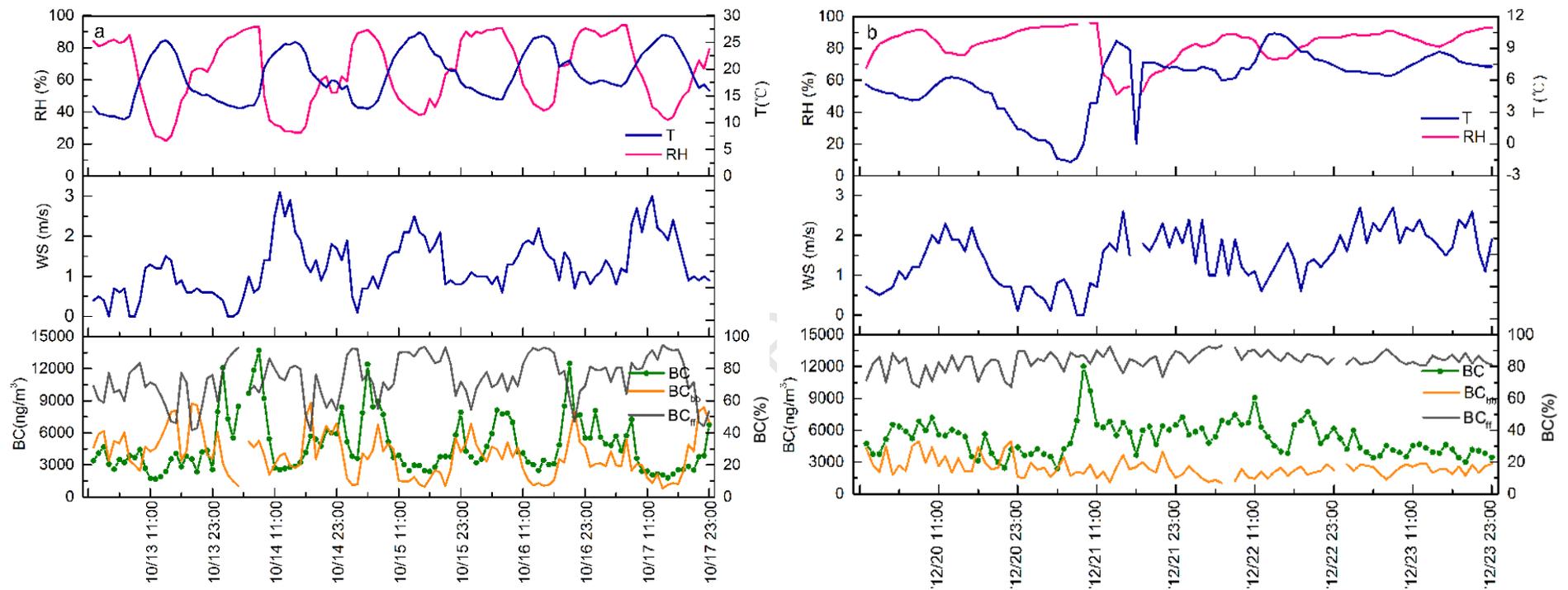


Figure 7. Time series of relative humidity (RH), temperature (T), wind speed (WS), BC mass concentration and the contributions of BC_{ff} (%) and BC_{bb} (%) during event I (a) and event II (b) observed at Nanjing.

HIGHLIGHTS

1. Both seasonal and diurnal variations showed significant changes of BC mass concentration.
2. A strong BC aerosol absorption observed during spring and winter seasons attributed to fossil fuel and biomass combustion.
3. Fossil fuel emission was the major source of BC in Nanjing contributing 81% as evidenced from the source apportionment model.
4. CWT analysis revealed major potential sources of BC originating from Anhui and Zhejiang Provinces.