Aerosol data assimilation using data from Fengyun-4A, a next-generation geostationary meteorological satellite

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Author contributions:

Xiaoli Xia: Methodology, Software, Investigation, Run the experiments, Writing Original Draft.

Jinzhong Min: Collect data, choose the dust storm, Methodology, Validation, Supervision.

Feifei Shen: Data quality control, Validation, Formal analysis, Visualization.

Yuanbing Wang: Assimilation system construction, Edit code, Writing.

Dongmei Xu: Processing of observation data, Calculating the BEC, Writing.

Chun Yang: Writing – review & editing

Peng Zhang: Writing – review & editing

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2	next-generation geostationary meteorological satellite
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25 Abstract

The Fengyun-4A (FY-4A) meteorological satellite, a next-generation 26 27 geostationary meteorological satellite, was launched on 11 December 2016. For instance, the Advanced Geosynchronous Radiation Imager 28 (AGRI) aboard FY-4A (AGRI/FY-4A) takes full-disk images at a 15-min 29 interval in 14 spectral bands with the 0.5-4-km resolution. Here we 30 developed data assimilation system based on the Gridpoint Statistical 31 Interpolation (GSI) system in which the Aerosol Optical Depth (AOD) 32 derived from FY-4A data were successfully assimilated for the first time. 33 The capability to assimilate FY-4A Aerosol optical depth (AOD) with an 34 hourly cycling configuration was then evaluated by a dust storm over 35 East Asia during 12-14 May 2019. The analyses initialized Weather 36 Research and Forecasting-Chemistry (WRF-Chem) model forecasts. The 37 system is tested with FY-4 AOD, Himawari-8 AOD in experiments and 38 then the results are compared to the Aerosol Robotic Network 39 (AERONET) AOD observations, which serving as the independent 40 observations. The results indicated that assimilating FY-4 AOD 41 substantially showed much better agreement with observations than those 42 from the control. Furthermore, the Bias and RMSE generally reduced 43 about 20% with forecast range. This study indicates that the aerosol data 44 assimilation using data from FY-4A can be used to improve the 45 performance of forecast model. 46

- 47 **Key words**: Fengyun-4 satellite; aerosol optical depth; data assimilation;
- 48 dust storm

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1. Introduction

- Economic growth has been very fast in China during the last decades 50 (Chandran et al., 2013). The great increases in anthropogenic emissions 51 during the last decades have caused serious air pollution problems in 52 China (Lei et al., 2011; Chan et al., 2008). The poor air quality mainly 53 related to the aerosol particles suspended in the near surface air, which 54 could downgrade air quality and visibility, and damage human health 55 (Pope et al., 2002; Peng et al., 2018). As the aerosols strongly absorb and 56 scatter solar radiation over the ultraviolet, visible light, and infrared 57 spectrum, they exert a significant influence on global climate and weather 58 processes (Rosenfeld et al., 2007; Wang et al., 2012). 59 Aerosol Optical Depth (AOD), a measure of the columnar extinction 60
- of solar radiation and long wave radiation by aerosols, is an important optical parameter in estimation of aerosol concentration, evaluation of the level of atmospheric pollution, and assessment of the climate effect of aerosols (He et al., 2014; Lohmann et al., 2006; Hu et al., 2012; Peng et al., 2015; Dai et al., 2014; Dai et al., 2015). To evaluate the effects of aerosols on climate it is necessary to estimate their spatial and temporal distributions.
 - Recent satellite-based remote sensing of AOD has greatly improved

our understanding of aerosols properties (Adhikary et al., 2008). Because 69 of its spatial and temporal coverage, satellite-based aerosol optical depth 70 is the most practical measurement of aerosol amount for global 71 assessments (Anderson et al., 2005; Hu et al., 2013). The satellite remote 72 sensing method is able to generate results over an extensive area in 73 real-time. The gained results can be updated periodically at a relatively 74 short span (He et al., 2014; Mustard et al., 2017). The uneven distribution 75 of conventional ground observation stations makes it more difficult to 76 study the spatial and temporal distribution characteristics of regional 77 aerosols. Such point-specific data are unable to yield a detailed pattern of 78 aerosol spatiotemporal distribution (Dai et al., 2018; Shen and Min, 2015; 79 Xu et al., 2013). Satellite remote sensing method can become a vital 80 addition to supplement ground-based measurement so as to gather 81 accurate aerosol information effectively and quickly (Stengel et al., 2009; 82 Zou et al., 2011). Meteorological satellite can be divided into 83 polar-orbiting and geostationary satellites by different operating orbits. 84 The polar-orbiting satellites products usually have the high spectral 85 resolution and can monitor the global atmospheric environment. However, 86 it can observe the same region only twice a day and so may miss rapidly 87 developing severe storms (Li & Zou, 2017; Qin et al., 2013; Wang et al., 88 2018). Conversely, geostationary instruments can provide continuous 89 imagery for resolving the evolution of weather phenomena from 90

mesoscale to the convective scale in the observed domain, due to their 91 fixed position relative to Earth's surface and high spatial and temporal 92 resolution (Montmerle et al., 2007; Wang et al., 2018; Xia et al., 2019b). 93 The Fengyun-4A (FY-4A) meteorological satellite. 94 a next-generation geostationary meteorological satellite, was launched on 95 11 December 2016. For instance, the Advanced Geosynchronous 96 Radiation Imager (AGRI) aboard FY-4A (AGRI/FY-4A) takes full-disk 97 images at a 15-min interval in 14 spectral bands with the 0.5-4-km 98 resolution. In addition, a full-disc image of aerosol optical depth (AOD) 99 100 offered by FY-4A covers a large area (80.6°N–80.6°S, 24.1°E–174.7°W) (Yang et al., 2017; Min et al., 2017; Wang et al., 2019). 101 One method to improve model forecasts of aerosols is data 102 assimilation (DA), which combines observations with numerical model 103 output and can reduce uncertainties of initial aerosol fields. Data 104 assimilation, which an approach first used in generating initial condition 105 for the numerical weather prediction, offers a means to reduce the 106 uncertainties in the estimates of aerosol distributions (Adhikary et al., 107 2008). Aerosol data assimilation first introduced since 2000 as aerosol 108 and chemical transport models have improved and more aerosol 109 observations have become available. Since then, there has been rapid 110 development in aerosol forecasting systems at many of the world's 111 numerical weather prediction (NWP) centers. Most researchers have 112

assimilated aerosol observations from satellite (Lynch et al., 2016; 113 Sekiyama et al., 2016; Zhang et al., 2011; Liu et al., 2011; Peng et al., 114 2015; Yumimoto et al., 2016; Fukuda et al., 2013; Saide et al., 2013; 115 Peng et al., 2016; Xia et al., 2019a) to chemical transport models. Both 116 these models showed that the assimilation greatly improved analysis 117 when comparing with independent observations. However, some 118 researchers indicated that AOD derived from polar-orbiting satellites 119 have limited spatial coverage and observation frequencies. While AOD 120 derived from geostationary satellites can overcome these limitations and 121 they also showed positive impact of geostationary data on air quality 122 prediction and implied that future geostationary missions also have 123 potential to greatly contribute to that (Lahoz et al., 2011; Yumimoto, 124 2013; Zoogman et al., 2014; Wang et al., 2004; Saide et al., 2014). 125 In this study, we construct an aerosol data assimilation experiment 126 with AOD derived from Fengyun-4 AGRI data and an aerosol data 127 assimilation system (Xia et al., 2019b) based upon GSI. To evaluate the 128 effectiveness in precipitation forecast, this system is applied and tested in 129 a heavy dust storm event occurred over East Asia during 12-14 May 2019. 130 This study represents the first attempt to assimilate FY-4A aerosol 131 observations with a rapid update data assimilation system and then it will 132 be investigated for the analyses and forecasts of the severe dust storm 133 event that occurred over East Asia during 12–14 May 2019. Details are 134

present in the rest of the paper. The next section describes the Fengyun-4
AGRI AOD data. while section 3 details the severe dust storm event and
the aerosol DA system setup for FY-4A assimilation such as the quality
control, observation error statistics, and background error covariance in
GSI systems. Results are presented in sections 4, and a brief discussion
regarding meteorological impact on the aerosol forecasts is provided in
sections 5.

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2. Observation Data

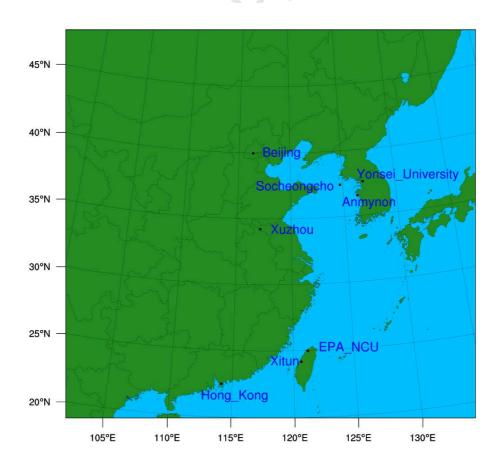
Fengyun-4 introduces a new generation Chinese 144 geostationary meteorological satellites, with the first FY-4A launched on 145 11 December 2016. The FY-4A satellite was equipped with four 146 advanced optical instruments aboard, including Advanced an 147 Radiation Geosynchronous Imager (AGRI), Geostationary 148 Interferometric Infrared Sounder (GIIRS), a Lightning Mapping Imager 149 (LMI) and Solar X-EUV Imaging Telescope (SXEIT). AGRI has 14 150 spectral bands from visible to infrared (0.45 -13.8 μ m) with high spatial 151 (1 km for visible at nadir, 2 km for near-infrared, and 4 km for remaining 152 infrared) and temporal (full-disk images at the 15-min interval) 153 resolutions. In addition, a full-disc image offered by FY-4A covers a 154 large area (80.6°N-80.6°S, 24.1°E-174.7°W). FY-4 represents an 155 exciting expansion in Chinese geostationary remote sensing capabilities 156

and detailed descriptions of FY-4A along with its products are given by Yang et al. (2017) and Min et al. (2017).

In this study, the level 2 total AOD retrievals over land and sea 159 provided by National Satellite Meteorological Center with a spatial 160 resolution of 4 km from FY-4A sensors on board the AGRI satellites 161 were used. We estimated observation errors to be the retrieval uncertainty 162 attached to the FY-4A AOD data plus a standard deviation calculated as 163 the representative error in the regridding (Zhang et al., 2008; Xia et al., 164 2019a). The FY-4 AOD retrievals uncertainty ranged from 0.0001 to 0.96, 165 with an average of 0.039. Quality control (QC) and accurate 166 quantification of the systematic bias and random errors of observations 167 arising from the errors in numerical models and instruments are key 168 components in assimilating satellite aerosols. We only assimilated the 169 highest quality AOD retrievals and thinned to the same resolution as the 170 model grid. As suggested by Saide et al. (2014), pixels adjacent to 171 missing values were discarded and only AOD values below 2.5 were used. 172 The observations with innovations exceeding 3 times of the observation 173 error were rejected before minimization to reduce cloud contamination 174 and noise in the data (Xia et al., 2019b; Wang et al., 2018). All QC 175 decisions, observation thinning, and observation error assignments for 176 aerosol observations were made within GSI for all experiments. 177

To allow comparison of simulation results from FY-4 data, the latest

version of the Himawari-8 AOD (Yoshida et al., 2018; Kikuchi et al., 2018; Dai et al, 2019), which are freely available at the website of the Japan Aerospace Exploration Agency (JAXA) Himawari Monitor (http://www.eorc. jaxa.jp/ptree/index.html) were regridded resolution of our model and then the Himawari-8 AOD were assimilated this study. The Aerosol Robotic Network (AERONET; http://aeronet.gsfc.nasa.gov/) provides **AOD** measured by sun photometers (Holben et al., 1998). Fig 1 shows the locations of Aerosol Robotic Network (AERONET) sites whose data were served as the independent observations in this study.



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Fig. 1. The experimental domain. Blue labeling indicates the ground-based AOD data

191 acquired by the AERONET sites used in this study.

3. Experimental Setup

3.1 The Severe Dust Storm Event

The study focused on an extreme dust storm event that occurred over East Asia in the year 2019 over northern China. During 12-14 May 2019, there was a large range of persistent heavy air pollution in large areas of China. The dust storm caused poor visibility, air quality, and flights' cancelation or delays over eastern China. AOD (500 nm) values observed at the Beijing AERONET site varied from less than 1 on 11 May to a maximum of 2 on 12 May, when China was severely affected by the dust storm.

3.2 Aerosol DA system

The Gridpoint Statistical Interpolation (GSI) 3DVAR algorithm is explained by Wu et al. (2002), and its expansion for aerosol DA using the GOCART aerosol module is described by Liu et al. (2011). We added a new interface to the FY-4A AOD data in this study as a follow on work of Liu et al. (2011) and Xia et al. (2019b). The variational method is a well-established approach which combines model background information with observations to obtain the "best" forecast possible. This approach is widely used in many NWPs centers. The method is based on minimization of a cost function which measures the distance between observations and their model equivalent, subject to a background

constraint usually provided by the model itself (Morcrette et al., 2009). 213 Associated with the background and observations are their error 214 characteristics. Given the background, observations, and errors, the 215 analysis vector (x) can be determined by minimizing a scalar 216 cost-function J(x). The cost function J(x) to be minimized with respect to 217 the bias parameters and model state becomes,

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$$J(x) = \frac{1}{2}(x - x_b)^T B^{-1}(x - x_b) + \frac{1}{2}[y - H(x)]^T R^{-1}[y - H(x)]$$
 (1)

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The covariance matrices determine how closely the analysis is 220 weighted toward the background and observations. H is the potentially 221 nonlinear "observation operator" that interpolates model grid point values 222 to observation locations and transforms model-predicted variables to 223 observed quantities. where Band Rare the background and 224 observation error covariance matrices of dimensions $m \times m$ and $p \times p$, 225 respectively. x_h denotes the background vector, y is a vector of 226 observations and B and R determine the relative contributions of the 227 background and observation terms to the final analysis. 228

The 3DVAR algorithm requires background error covariance (BEC) statistics for each analysis variable. GSI uses recursive filters and permits spatially inhomogeneous BECs (Wu et al., 2002). The BEC were computed for each aerosol species via the "National Meteorological Center (NMC) method" (Parrish and Derber, 1992) by taking the differences of 24 and 12 h Weather Research and Forecasting with

Chemistry (WRF/Chem) forecasts of the 14 aerosol species (including hydrophobic and hydrophilic organic carbon [OC] and black carbon [BC]; sulfate; sea-salt in four particle-size bins (effective radii of 0.3, 1.0, 3.25, and 7.5 μm for dry air); and dust particles in five particle-size bins (effective radii of 0.5, 1.4, 2.4, 4.5, and 8.0 μm)) valid at the same time for 62 pairs valid at either 0000 or 1200 UTC from 11 April 2019 to 11 May 2019.

3.3 Experimental design

In this study, we performed three numerical experiments to evaluate the impact of FY-4 AOD DA on aerosol analyses and forecasts over the East Asia. One experiment ("CNT") served as the control and did not employ any DA. The other two experiments all performed 3DVAR DA but assimilated different observations. The FY-4 AOD data were assimilated by one experiment ("FY DA"), while the other assimilated Himawari-8 AOD observations ("Hima DA"). It is necessary to assimilate high-frequency data at short intervals so that the initial field can contain the information of the dust storm system as much as possible (Xia et al.,2019b). So observations taken within 1 hour of each analysis time were assimilated. In our domain, Himawari-8 AOD provided coverage only 11–12 h d (daytime) (Yumimoto et al., 2016). Thus, the DA experiments only cycled during daylight hours and short forecasts at night.

All observations were assumed to be valid at the analysis time for each experiment.

The model used to simulate the transport of aerosols and chemical 259 species was the WRF-Chem (Grell et al., 2005). The data assimilation 260 and forecast system constructed in this study is based on version 3.8.1 of 261 the WRF-Chem model. As in Liu et al. (2011) and Schwartz et al. (2012), 262 Goddard Chemistry Aerosol Radiation and Transport (GOCART) aerosol 263 scheme was chosen as the aerosol option within WRF-Chem. And then 264 forecasts of 3D mass mixing ratios of 14 aerosol species are produced: 265 hydrophobic and hydrophilic OC and BC; sulfate; sea-salt in four 266 particle-size bins (effective radii of 0.3, 1.0, 3.25, and 7.5 um for dry air); 267 and dust particles in five particle-size bins (effective radii of 0.5, 1.4, 2.4, 268 4.5, and 8.0 µm). The experiments were all run over the same domain, 269 which covers main area of east China (Figure 1) and the three 270 experiments were implemented with the same model configuration: $200 \times$ 271 150 horizontal mesh grid using 20-km spacing and 40 vertical levels up to 272 50 hPa. The NCAR's MOZART-4 model (Pang, 2012), the Rapid 273 Radiative Transfer Model longwave radiation scheme (Mlawer et al., 274 1997), the Dudhia shortwave radiation scheme (Dudhia, 1989), Noah land 275 surface model, the Yonsei University (YSU) planetary boundary layer 276 scheme (Hong et al., 2006; Hu et al., 2010) are used for all deterministic 277 forecasts. 278

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The three experiments all initialized a new WRF-Chem forecast 279 every 1 h between 0100 UTC 12 May and 1100 UTC 14 May. The initial 280 aerosol fields were produced by a 5-day forecast, which was needed to overcome the unrealistic of the WRF/Chem forecasting, similar to 282 Pagowski et al. (2010) and Schwartz at al. (2012). Then the newly 283 developed 3DVAR aerosol DA system uses 14 individual aerosol species 284 of the WRF/Chem built-in GOCART module as control variables, 285 including hydrophobic and hydrophilic organic carbon [OC] and black 286 carbon [BC]; sulfate; sea-salt in four particle-size bins (effective radii of 287 0.3, 1.0, 3.25, and 7.5 µm for dry air); and dust particles in five 288 particle-size bins (effective radii of 0.5, 1.4, 2.4, 4.5, and 8.0 µm). The 289 two experiments that assimilated observations performed a new aerosol 290 analysis every 1 h to update the control variables before initializing WRF-Chem forecasts. The first analyses used the 0100 UTC 12 May 292 fields as backgrounds, while subsequent analyses used the previous 293 cycle's 1h aerosol forecasts as backgrounds. The experiment without DA 294 initialized its first forecast from the 0100 UTC 12 May fields, while 295 initial aerosol fields for future forecasts were simply taken from the 296 previous cycle's 1h forecast. No meteorological DA was performed. 297 Every initialization, all experiments' meteorological fields were updated 298 by interpolating FNL analyses onto the computational domain. 299

The analyses and forecasts from the three experiments were

- compared to AOD observations from FY-4, Himawari-8 and AERONET.
- The results of these comparisons are now described.

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4. Results

4.1. Comparison to Fengyun-4 AOD and Himawari-8 AOD

A variety of metrics were used to verified the WRF-Chem forecast. 306 Many statistics assessing forecast accuracy can be defined based on the 307 correspondence between the model and observations (Schwartz at al., 308 2012). The mean Bias is simply the difference of the mean and observed 309 values and can be considered a measure of systematic model error. The 310 root mean square error (RMSE) could quantify how individual 311 forecast-observation pairs agree and can be interpreted as a measure of 312 non-systematic model error. To evaluate the experiments' ability to 313 discriminate between events, we calculate the average RMSE and Bias 314 between those model fields and the observations. In order to evaluate the 315 overall performance of the DA system, time series of the hourly pollutant 316 concentrations from the CNT forecast, the DA analysis, and the DA 317 forecast of the three experiments were compared with the independent 318 observations in the domain (Fig. 2). In Fig 2, the X-axis presents analysis 319 time, Y-axis presents values of AOD values, Bias and RMSEs. 320 Furthermore, Table 1-3 summarized the statistical analysis of the 321 simulated and observed AOD in the experiments. 322

323	In the CNT experiment, it did not perform very well, although it was
324	able to capture the synoptic variability when there was a severe dust event
325	In order to see the evolution process of aerosol more directly, Time series
326	of the absolute AOD values, Bias and RMSE of the simulated AOD in the
327	FY DA systems and in the Hima DA systems during the period of 12-14
328	May 2019 were included in Fig. 2. Based on the model performance
329	evaluation statistical metrics, both two assimilation experiments are
330	superior to the CNT experiment, indicating that all of the analyses can
331	better represent the temporal evolution characteristics of aerosols than the
332	CNT experiment. It can be seen from the Fig. 2a and Fig. 2b that after
333	assimilation the AOD value has been significantly improved in both FY
334	DA and Hima DA experiments. These statistics indicate that the DA
335	system was able to adjust the analysis and forecast fields. As shown in the
336	Table 1 and Fig. 2c-2f, there were larger systematic biases and RMSEs
337	for the CNT experiment. The biases were lower than the corresponding
338	observed concentrations, suggesting a significant systematic
339	underestimation of the WRF-Chem simulation. It can be seen that the
340	Bias and RMSE generally reduced with forecast range, indicating that the
341	DA system was well calibrated. After the assimilation of AOD
342	observations, the magnitudes of the bias and the RMSEs decreased. The
343	Biases were near -0.46 in the CNT and near -0.37 after DA; the RMSE
344	were near 0.45 in the CNT and near 0.36 after DA. Thus the Bias and

RMSE reduced about 20% compared to the CNT experiment, indicating that the analysis fields were very close to the observations. Interestingly, in the Hima DA experiment, the Bias and RMSE values were reduced a little more than in the FY DA systems at sometime. These statistics indicate that the DA system was able to adjust the analysis and forecast fields.

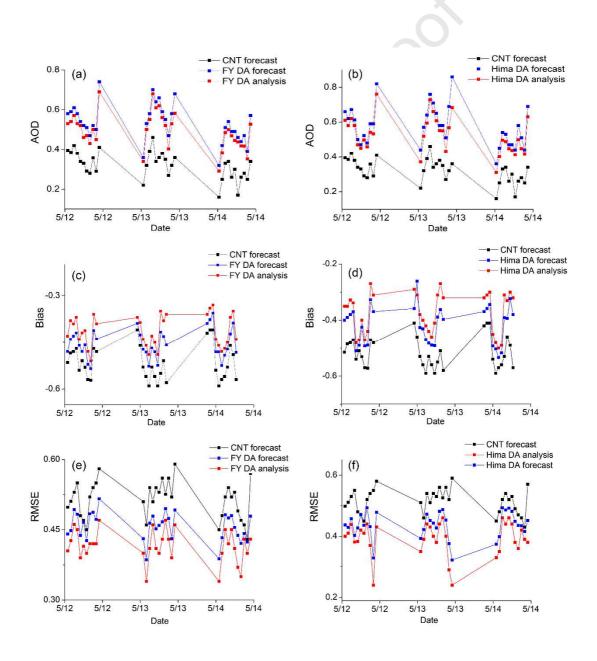


Fig. 2. Time series of the absolute AOD values (a, b), Bias (c, d) and RMSE (e, f) of the simulated AOD. The DA systems involved FY DA AOD data (a, c, e) and Hima DA AOD data (b, d, f) during the period of 12-14 May 2019.

Table 1 Statistical analysis of the simulated and observed AOD in the experiments on 12 May 2019.

Date	FYDA				HimaDA	·			CNT	
	\mathbf{B}_{b}	\mathbf{B}_{a}	R_b	R_a	\mathbf{B}_{b}	B_a	R_b	R_a	В	R
051201	-0.482	-0.431	0.441	0.405	-0.402	-0.355	0.437	0.407	-0.515	0.498
051202	-0.443	-0.387	0.448	0.427	-0.396	-0.352	0.429	0.413	-0.485	0.511
051203	-0.485	-0.392	0.493	0.461	-0.389	-0.328	0.457	0.438	-0.487	0.534
051204	-0.423	-0.376	0.483	0.452	-0.374	-0.337	0.403	0.381	-0.473	0.551
051205	-0.466	-0.447	0.438	0.391	-0.512	-0.486	0.427	0.383	-0.546	0.488
051206	-0.481	-0.428	0.466	0.415	-0.492	-0.471	0.471	0.426	-0.512	0.472
051207	-0.458	-0.412	0.427	0.409	-0.425	-0.404	0.436	0.415	-0.534	0.453
051208	-0.512	-0.483	0.484	0.428	-0.492	-0.473	0.493	0.442	-0.570	0.526
051209	-0.535	-0.516	0.487	0.426	-0.489	-0.443	0.431	0.375	-0.572	0.540
051210	-0.412	-0.368	0.472	0.423	-0.327	-0.272	0.329	0.249	-0.474	0.557
051211	-0.439	-0.395	0.516	0.477	-0.369	-0.316	0.478	0.433	-0.481	0.589

Table 2 Statistical analysis of the simulated and observed AOD in the experiments on 13 May 2019.

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Date	FYDA		1		HimaDA				CNT	
	B_b	B_a	R_b	R_a	B_b	\mathbf{B}_{a}	R_b	R_a	В	R
051301	-0.389	-0.373	0.431	0.406	-0.358	-0.292	0.356	0.393	-0.422	0.515
051302	-0.268	-0.255	0.386	0.342	-0.261	-0.211	0.422	0.391	-0.417	0.468
051303	-0.473	-0.442	0.464	0.415	-0.426	-0.383	0.472	0.443	-0.416	0.543
051304	-0.482	-0.461	0.479	0.463	-0.431	-0.406	0.451	0.434	-0.543	0.515
051305	-0.527	-0.490	0.411	0.452	-0.469	-0.427	0.443	0.408	-0.597	0.544
051306	-0.469	-0.438	0.459	0.408	-0.482	-0.441	0.427	0.387	-0.575	0.536
051307	-0.481	-0.454	0.467	0.433	-0.465	-0.489	0.481	0.442	-0.564	0.569
051308	-0.524	-0.496	0.495	0.477	-0.466	-0.491	0.487	0.466	-0.533	0.527
051309	-0.417	-0.355	0.474	0.430	-0.312	-0.389	0.453	0.401	-0.460	0.561
051310	-0.412	-0.431	0.431	0.399	-0.277	-0.392	0.376	0.293	-0.494	0.527
051311	-0.458	-0.367	0.388	0.492	-0.184	-0.357	0.322	0.245	-0.571	0.593

Table 3

Statistical analysis of the simulated and observed AOD in the experiments on 14 May	2019.
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Date	FYDA	HimaDA	CNT

		B_b	B_a	R_b	R _a	B_b	\mathbf{B}_{a}	R _b	R_a	В	R
	051401	-0.389	-0.365	0.388	0.346	-0.369	-0.325	0.374	0.338	-0.423	0.452
	051402	-0.377	-0.346	0.433	0.409	-0.358	-0.317	0.399	0.354	-0.417	0.489
	051403	-0.356	-0.332	0.482	0.453	-0.344	-0.306	0.493	0.462	-0.416	0.525
	051404	-0.481	-0.443	0.475	0.428	-0.492	-0.452	0.486	0.443	-0.542	0.540
	051405	-0.483	-0.469	0.480	0.451	-0.503	-0.489	0.492	0.461	-0.593	0.523
	051406	-0.525	-0.486	0.455	0.412	-0.534	-0.502	0.481	0.447	-0.571	0.537
	051407	-0.492	-0.472	0.438	0.375	-0.517	-0.491	0.449	0.386	-0.566	0.496
	051408	-0.472	-0.451	0.421	0.356	-0.391	-0.314	0.435	0.365	-0.534	0.471
	051409	-0.423	-0.378	0.442	0.432	-0.393	-0.332	0.434	0.429	-0.468	0.466
	051410	-0.388	-0.353	0.424	0.406	-0.323	-0.306	0.416	0.393	-0.490	0.438
_	051411	-0.483	-0.440	0.479	0.431	-0.382	-0.328	0.451	0.380	-0.575	0.575

In order to evaluate the improvement of AOD spatial distribution in the initial field by AOD assimilation, we conducted a comparative analysis of the background field and the analysis field, in which the simulation value of AOD was calculated by the extinction coefficient of 550nm wavelength output of the model (Fig. 3). Observations of 1-hour DA windows from FY4, distributions of simulated AOD based on DA the background fields, the analysis fields and the 1-h forecast fields for the FY DA system were displayed in Fig 3a, 3b, 3c and 3d, respectively.

The simulation of the dust storm in the CNT experiment was inadequate (Fig 3b). After the assimilation of the satellite AOD data, the analysis field added the aerosol distribution of the main sources of dust in the dust storm weather - the middle and eastern regions, the north plain and the southern regions of East Asia. The assimilation experiments all reflected the maximum positive increment center of AOD in the region, which was consistent with the distribution of the high value region of

AOD in the satellite observation data. After the assimilation adjustment, the analysis field (Fig 3c) has more abundant observation information, and the aerosol distribution of the analysis field is closer to the satellite observation. The assimilation experiments all reflected the maximum positive increment center of AOD in the region, which was consistent with the distribution of the high value region of AOD in the satellite observation data. After the assimilation adjustment, the analysis field has more abundant observation information, and the aerosol distribution of the analysis field is closer to the satellite observation. As the AOD represents contributions from all aerosol types, so we also plotted the distributions of the dust field (Fig. 4), which gives a direct indication of the dust storm event.

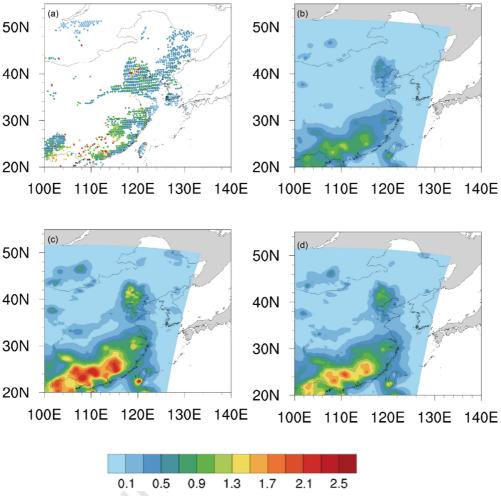


Fig.3. Observations of 1-hour DA windows from FY4(a), The result of the control experiment (b), the analysis fields (c) and the 1-h forecast fields (d) for the FY DA system, valid at 0600 UTC 12 May 2019.

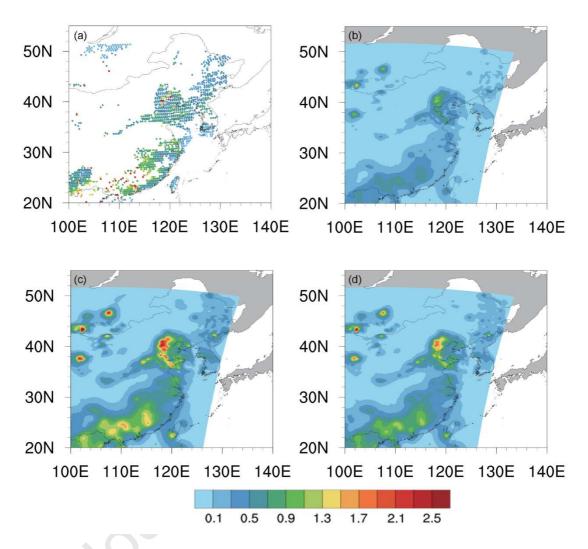


Fig. 4. Observations of 1-hour AOD DA windows from the FY4 (a). The dust fields from CNT experiments (b). Distributions of simulated dust field based on FY DA analysis(c), and the forecast dust fields from FY DA experiments (d), valid at 0600 UTC 12 May 2019. The dust fields here included dust particles in five particle-size bins (effective radii of 0.5, 1.4, 2.4, 4.5, and $8.0~\mu m$).

Similar to Fig. 3, it can be seen that the assimilation of AOD data can predict the distribution of large dust centers in Beijing and Tianjin more accurately, which improving the underestimation of dust values in Beijing and Tianjin in the CNT experiment. At the same time, the

addition of observation data also strengthened the distribution of large 409 AOD values in southeast China, Vietnam, northeast China region, making 410 up for the underestimation of CNT and reflecting the real aerosol 411 diffusion and distribution. In general, AOD data assimilation can 412 significantly improve the forecast field, which can improve the intensity 413 and distribution of AOD model, and correct the location of strong center 414 of dust area. 415 In order to determine the contributions of major aerosol components 416 to the distribution of large value AOD, we showed the spatial 417 distributions of the analysis field of the individual component (i.e., dust, 418 sulfate, organic carbon [OC], black carbon [BC], sea salt, P25 and P10) in 419 the FY-4 assimilation experiments (Fig 5) and in the Himawari-8 420 assimilation experiments (Fig 6). As shown in Fig. 5 and Fig. 6, the 421 increase of the AOD over the Gobi Desert and Beijing areas for all the 422 two assimilations are all mainly dominated by the dust aerosols, 423 indicating that the underestimation of AOD over the Gobi Desert and 424 Beijing areas in the CNT experiment is mainly caused by the 425 underestimation of dust aerosols. In contrast with the north China, in the 426 south China the large underestimation of AOD of the CNT experiment, 427 which are mainly caused by the underestimation of P25 component, are 428 also removed by the assimilation of the FY-4 AODs and the Himawari □8 429

AODs. The large P25 component mainly dominated to the AOD over

- south China in all the two assimilation experiments. All of the analyses
- correctly increase the AODs to better match the observations.

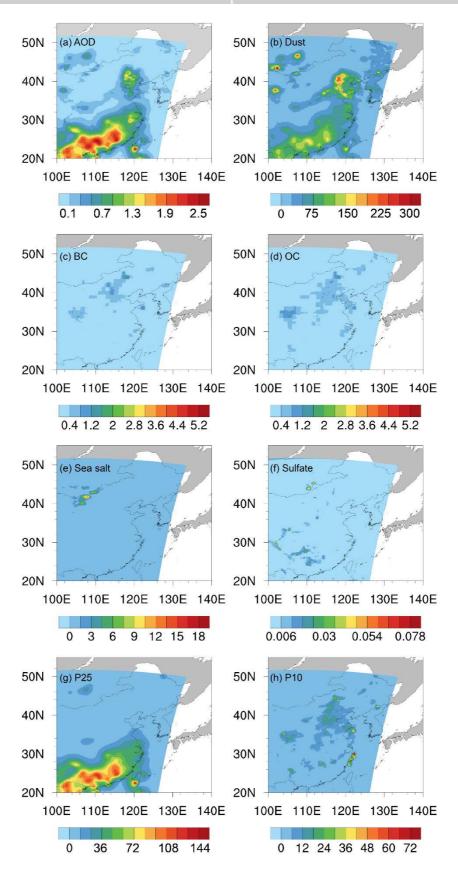


Fig 5. The horizontal distributions of the AOD and the analysis field of individual component (μ gkg⁻¹) (i.e., dust (b), sulfate (c), organic carbon [OC] (d), black carbon [BC] (e), sea salt (f), P25 (g) and P10 (h)) in the FY-4 assimilation experiments at 0600 UTC 12 May 2019.

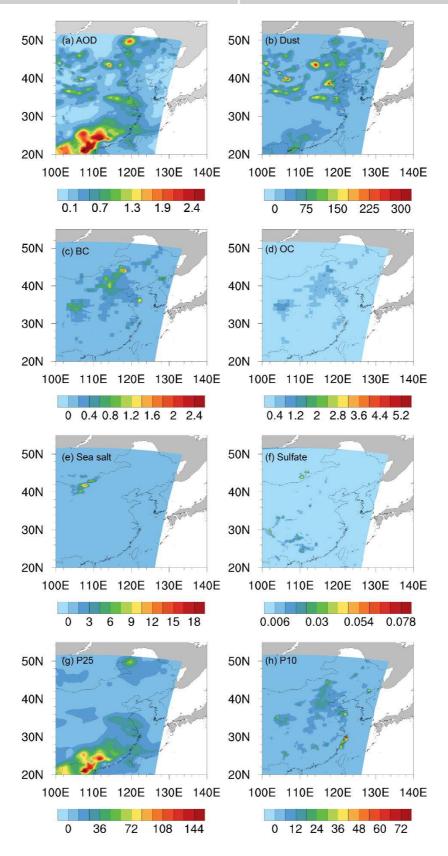


Fig 6. The horizontal distributions of the AOD (a) and the analysis field of individual component (μ gkg⁻¹) (i.e., dust (b), sulfate (c), organic carbon [OC] (d), black carbon [BC] (e), sea salt (f), P25 (g) and P10 (h)) in the Himawari-8 assimilation experiments at 0600 UTC 12 May 2019.

4.2. Comparison with AERONET AOD

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NASA's Aerosol Robotic Network (AERONET) (Holben et al., 443 1998; O'Neill et al., 2003) is a well-established ground-based network of 444 models and measurements due to its small wavelength-independent 445 measurement error, with calibration error on the order of 0.015 (Holben 446 et al., 1998; Eck et al., 1999). AERONET data plays an important role in 447 the validation of aerosol satellite remote sensing products and model 448 products. 449 The ground-based AOD data acquired by AERONET were used to 450 evaluate the assimilation results. Fig 7 shows comparisons of modeled 451 results in the three experiments with AOD from AERONET. The dust 452 storm from 12 May to 14 May 2019 was detected by AERONET at the 453 Anmyon, Beijing, EPA_NCU, Xuzhou, Hong_kong, Socheongcho, Xitun 454 and Yonsei_university sites. At all sites, DA increased AOD values 455 compared to the control and usually agreed better with observations. In 456 some cases, the increase due to DA still was insufficient (e.g., Beijing, 457 EPA_NCU and Xuzhou; Fig 7b, Fig 7c and Fig 7d). As shown in Fig. 7, 458 the hourly assimilation of AOD improves the model predictions. The 459 background model predictions were generally under predicting the 460 observations, which the assimilation corrects for and brings the AOD 461 values after assimilation closer to the AERONET observations. Quite 462 generally speaking, results for AOD in FY-4A and Himawari-8 DA 463

experiments agree better with the observations because of the 464 assimilation. At Anmyon site, we can also see the assimilation improved 465 the underestimation by the CNT experiment. The values in Hima DA 466 were usually larger than those in FY DA and the Hima assimilation 467 overestimate AOD on 13 May. At Beijing site, where more observations 468 are available, there was higher improvements presented. At EPA_NCU 469 and Xuzhou sites, we clearly see a strong positive impact of the 470 assimilation on AOD. The AOD in two DA experiments matched 471 AERONET observations very well in most of the time except when the 472 air pollution reached a high level and the observed AOD value exceeded 473 1.0. At Hong kong, Socheongcho, Xitun and Yonsei university sites, we 474 see little difference on AOD for both FY DA and Hima DA experiments 475 and both agree reasonably well with the independent observations. In 476 general, the result presents good assimilation efficiency to improve the 477 capability of the model to simulate AOD over Eastern Asia. 478

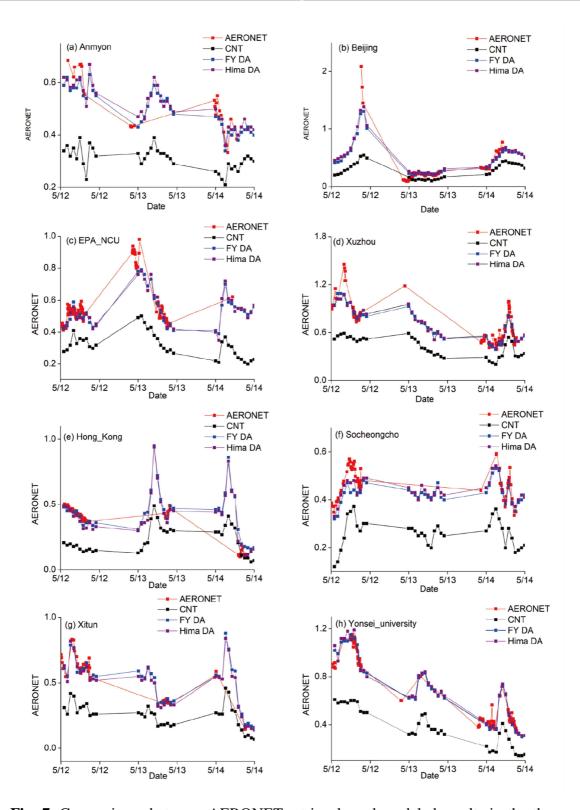


Fig. 7. Comparisons between AERONET retrievals and modeled results in the three experiments from 12 May to 14 May 2019, at the AERONET sites of (a) Anmyon, (b) Beijing, (c) EPA-NCU, and (d) Xuzhou, (e) Hong_Kong, (f) Socheongcho, (g) Xitun, (h) Yonsei_university.

5. Summary and discussion

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In this study, the ability of assimilating hourly averaged AOD observations from the FY-4 was added into the aerosol DA framework within the GSI 3DVAR analysis system. Parallel experiments assimilated AOD from FY-4A and Himawari-8 separately and a control experiment that did not employ DA was also performed.

Assimilation of satellite AOD data improved the initial field of the model obviously. AOD assimilation provides the analysis field with more abundant aerosol observation information and more accurate description of the model initial field. AOD data assimilation showed a positive effect on prediction, reasonably improved the intensity and distribution of each variable, and made it more consistent with the observation reality. The statistical results indicated that this assimilation system agreed better with the observations. In addition, aerosol analyses and forecasts with FY-4A AOD DA was substantially improved when compared to independent AOD observations from AERONET sites. The comparison with independent AERONET observations showed that the assimilation can substantially improves modelled AOD. The Bias and RMSE reduced about 20% of the corresponding observed concentrations, indicating that the analysis fields were very close to the observations. The results presented here showed that the FY-4A AOD data assimilation system has a broad development prospect in the application of air quality prediction.

508	Future developments of the assimilation system will include more
509	advanced DA techniques, high resolution emissions. In the future work,
510	the AOD and PM2.5, PM10 could try to be assimilated together to
511	improve the capability of air pollution forecast.
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Highlights

- 1) It is the first attempt to assimilate FY-4A AOD with a rapid-update DA system.
- 2) It was investigated for the dust storm occurred over East Asia on 12-14 May 2019.
- 3) General positive impacts were achieved from assimilating high-frequency data.

Declaration of Inter	ests
	are that they have no known competing financial interests or personal relationships beared to influence the work reported in this paper.
□The authors decla as potential compet	are the following financial interests/personal relationships which may be considered ting interests:

Sincerely,

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